How an uncertain short-wave perturbation on the North Atlantic wave guide affects the forecast of an intense Mediterranean cyclone (Medicane Zorbas)

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Abstract. Mediterranean cyclogenesis is known to be frequently linked to ridge building over the North Atlantic and subsequent anticyclonic Rossby-wave breaking over Europe. But understanding of how this linkage affects the medium-range forecast uncertainty of Mediterranean cyclones is limited, as previous predictability studies mainly focused on the relatively rare cases of Mediterranean cyclogenesis preceded by upstream extratropical transition of tropical cyclones. This study exploits a European Centre for Medium-Range Weather Forecast (ECMWF) operational ensemble forecast with an uncertain PV streamer position over the Mediterranean that, three days after initialization, resulted in an uncertain development of the Mediterranean tropical-like cyclone (Medicane) Zorbas in September 2018. Later initializations showed substantially lower forecast uncertainties over the Mediterranean. An ad-hoc clustering of the ensemble members according to the PV streamer position in the Mediterranean is used to study the upstream evolution of the synoptic to mesoscale forecast uncertainties. Cluster differences show that forecast uncertainties amplified at the stratospheric side of a jet streak over the North Atlantic during the first day of the ensemble prediction. Subsequently, they propagated downstream and further amplified within a short-wave perturbation along the wave guide, superimposed to the large-scale Rossby-wave pattern. After three days, the uncertainties reached the Mediterranean, where they resulted in a large spread in the position of the PV streamer. These uncertainties further translated into uncertainties in the position and thermal structure of the Mediterranean cyclone. In particular, the *eastward* displacement of the PV streamer in more than a third of the ensemble members resulted in a very different cyclone scenario. In this scenario, cyclogenesis occurred earlier than in the other members in connection to a pre-existing surface trough over the Levantine Sea. These cyclones did not develop the deep warm core typical for medicanes. It is proposed that the *eastward* shifted cyclogenesis resulted in reduced values of low-level equivalent potential temperature in the cyclogenesis area. As a result, latent heating was not intense and deep enough to erode the upper-level PV anomaly and allow the formation of a deep warm core. The *westward* displacement lead to surface cyclones than were too weak and a medicane formed in only half of the members. The *central*, i.e. correct, PV streamer position resulted in the most accurate forecasts with a strong medicane in most members. These results extend current knowledge of the role of upstream uncertainties for the medium-range predictability and unsteady forecast behavior of Mediterranean cyclones in general, and medicanes in particular. Moreover, this study is the first that explicitly investigates the role of an uncertain PV streamer position for medicane development.

1 Introduction

Mediterranean cyclones are major causes of heavy precipitation and strong winds all across the Mediterranean basin and are therefore the main meteorological threat in this region (e.g. Jansa et al., 2001; Pfahl and Wernli, 2012; Raveh-Rubin and Wernli, 2015). Hence, it is of high societal relevance to study uncertainties in their forecasts. It is widely known that Mediterranean cyclones mostly form from elongated equatorward intrusions of stratospheric high potential vorticity (PV), so-called PV streamers, that are the result of anticyclonic Rossby-wave breaking over the North-Atlantic/European region (e.g. Tafferner, 1990; Fita et al., 2006; Tous and Romero, 2013; Flaounas et al., 2015). Often, these PV streamers break up into PV cutoffs, which are typically identified as quasi-circular PV anomalies isolated from the main stratospheric reservoir on an isentropic surface (Appenzeller and Davies, 1992; Wernli and Sprenger, 2007). The dynamical forcing and destabilization by the upper-level PV streamer or cutoff are known to favor strong convection and cyclogenesis. Indeed, a range of case studies established that mesoscale errors in the structure of the upper-level PV streamer or cutoff are relevant sources of errors in the low-level development of Mediterranean cyclones and heavy precipitation events (Fehlmann and Davies, 1997; Fehlmann and Quadri, 2000; Romero, 2001; Homar and Stensrud, 2004; Argence et al., 2008; Chaboureau et al., 2012). Other case studies have shown that intense latent heating in upstream cyclones over the North Atlantic can enhance anticyclonic Rossby-wave breaking and thereby affect the PV streamer formation and, as a result, Mediterranean cyclogenesis and heavy precipitation (e.g. Massacand et al., 2001; Grams et al., 2011; Pantillon et al., 2015). Wiegand and Knippertz (2014) showed that far equatorward reaching PV streamers over the North Atlantic and Mediterranean are systematically preceded by latent heating upstream. The basic mechanism behind this link is that intense latent heating results in the transport of low-PV air from low levels to the upper troposphere where it contributes to the build-up or amplification of a ridge and, hence, the formation of a PV streamer downstream. Similar situations also occur in other regions, e.g. the North Pacific (Grams and Archambault, 2016). Such substantial latent heat release often occurs for instance in warm conveyor belts (WCBs), which are coherent, rapidly ascending airstreams within extratropical cyclones (Browning, 1990; Madonna et al., 2014). In a systematic analysis, Raveh-Rubin and Flaounas (2017) found that more than 90% of the 200 strongest Mediterranean cyclones were preceded by North Atlantic cyclones with substantial WCB activity, out of which five where tropical cyclones that underwent extratropical transition (ET). Situations with ET result in especially strong WCBs, ridge building, and downstream flow amplification (e.g. Riemer et al., 2008; Torn, 2010; Archambault et al., 2013; Quinting and Jones, 2016). Therefore, in recent years, particular attention has been devoted to cases where the genesis of intense Mediterranean cyclones or heavy precipitation over Europe was preceded by ET in the North Atlantic, thereby establishing a causal link s between the two (e.g. Grams et al., 2011; Chaboureau et al., 2012; Pantillon et al., 2013, 2015; Grams and Blumer, 2015).

Mediterranean cyclone Zorbas in September 2018, whose formation due to an upper-level PV streamer is the subject of this study, led to considerable damage through severe winds, torrential rainfall, major flooding, and even tornadoes [Wikipedia and references therein]. The main affected regions were Tunisia and Lybia in the early stages and southern Greece and Turkey later on [Wikipedia and references therein]. A specialty of this cyclone was that it belonged to a special class of Mediterranean cyclones that develop visual and structural similarities to tropical cyclones in a process similar to tropical transition [for further

details on tropical transition see Davis and Bosart (2004)]. These cyclones are often called medicanes (Mediterranean hurricanes; Emanuel, 2005; Tous and Romero, 2013; Cavicchia et al., 2014) or Mediterranean tropical-like cyclones (Miglietta et al., 2013). Medicanes are often associated with strong horizontal pressure gradients, wind, and rainfall, which may result in large damage, although medicanes rarely attain hurricane intensity. Sometimes they acquire the typical appearance of a hurricane, with convective cloud bands wrapped around a cloud-free central eye and a typical size of the associated cloud clusters on the order of 300 km in diameter, i.e. smaller than typical hurricanes (Emanuel, 2005). There is not yet a clear definition of medicanes. A common property seems to be the development of a (symmetric) warm-core structure throughout the whole troposphere, often called deep warm core. The deep warm core of medicanes is not necessarily produced directly by convection but can, especially at lower levels, be promoted by horizontal advection and the seclusion of warm air in the cyclone centre (Fita and Flaounas, 2018). The relative role of the positive upper-level PV anomaly and air-sea interaction for the intensification of medicanes is currently debated. In most cases the positive upper-level PV anomaly seems to be important for the initial intensification (Miglietta et al., 2017). This study focuses on uncertainties related to the PV streamer that lead to genesis and initial intensification of Medicane Zorbas.

Using ECMWF operational ensemble forecasts, Pantillon et al. (2013) investigated the predictability of a medicane that was preceded by upstream ET. They found that ensemble members that appropriately forecasted the cyclone track during ET, correctly reproduced ridge building and the resulting PV streamer over the Mediterranean, while ridge building was reduced and the PV streamer was absent in the ensemble members that wrongly predicted a southward deflection of the tropical cyclone. At the same time, only few of the ensemble members with a PV streamer forecasted a downstream medicane, showing that the predictability of the medicane was still strongly limited, even when the representation of upstream ET was roughly correct and a PV streamer was forecasted. This indicates that other synoptic-scale aspects related to the PV streamer, for example uncertainty in its position, shape, or intensity limited the predictability of the medicane (similar to Maier-Gerber et al., 2019). The existence of several 'stages' in cyclone predictability is consistent with the systematic analysis of the predictability of recent medicanes by Di Muzio et al. (2019), which showed that such "forecast jumps" occur at different lead times for different forecast parameters. For example, rapid increases in the probability of cyclone occurrence can be found at longer lead times (4-7 days before the mature phase of the medicane) than rapid decreases of the spread of cyclone position (2-5 days) and of the strength of the upper-level warm core (initializations after cyclogenesis). While the dynamical link between uncertain upstream ET and uncertainties in the occurrence of Mediterranean cyclones and heavy precipitation is well established, Raveh-Rubin and Flaounas (2017) showed that only a small percentage of the most intense Mediterranean cyclones is preceded by ET in the North Atlantic. Therefore current understanding of the role of upstream processes for uncertainties and jumps in the forecasts of intense Mediterranean cyclones in general, and medicanes in particular, is limited.

The present study addresses this knowledge gap by presenting a chain of dynamical processes linking ensemble forecast uncertainties in the North Atlantic to an uncertain PV streamer position over the Mediterranean, and subsequently to an uncertain position and thermal structure of Medicane Zorbas. It is also one of the first studies that explicitly documents the role of an uncertain PV streamer position for uncertainties in the forecast of a medicane. To this aim, it investigates the ECMWF ensemble prediction initialized 84 h before cyclogenesis of Medicane Zorbas, which was the last initialization before a rapid decrease in

ensemble spread of geopotential height at 500 hPa over the Mediterranean (i.e. a forecast jump) the day before cyclogenesis.
This suggests that this particular ensemble forecast provides valuable information about the dynamics related to the forecast jump.

Many previous medicane studies have used high-resolution, convection permitting models to study small-scale processes like air-sea interaction and convection involved in the transition to and during the tropical-like phase (Fita et al., 2007; Davolio et al., 2009; Mazza et al., 2017; Fita and Flaounas, 2018; Miglietta and Rotunno, 2019). In this study, forecasts of the comparably coarse ECMWF operational ensemble forecast model with parametrized convection are used to study synoptic- to mesoscale forecast uncertainties before cyclogenesis and during the initial 24-hour period of cyclone intensification. This data is valuable because the model is able to reproduce medicanes adequately (Pantillon et al., 2013; Di Muzio et al., 2019). Additionally, other studies have shown that a detailed analysis of ensemble forecasts, especially of different scenarios in a particular ensemble prediction, can be highly rewarding for better understanding the involved dynamics and predictability limits. As an example, such scenarios have already been used to identify key dynamical elements limiting the predictability of medicanes (Pantillon et al., 2013), but also tropical cyclones (Torn et al., 2015; Pantillon et al., 2016; Gonzalez-Aleman et al., 2018; Maier-Gerber et al., 2019) and atmospheric blocking (Quandt et al., 2017).

The remainder of this article is structured as follows. After a description of the data and methods in Sect. 2, an overview of the large-scale situation prior to cyclogenesis and the synoptic evolution of Zorbas is given in Sect. 3. In Sect. 4, uncertainties of Zorbas in the operational ensemble forecast are discussed and a pragmatic clustering is introduced that uses the uncertainty in the PV streamer position to separate the ensemble forecast into three distinct PV streamer scenarios. Sect. 5 discusses the origin and the dynamical pathway of the forecast uncertainty that lead to the three PV streamer scenarios and, as a result, uncertain position and thermal structure of the Mediterranean cyclone. Finally, the main conclusions are summarized in Sect. 6.

2 Data and methods

2.1 Data

The basic data for this study are from the ECMWF Integrated Forecasting System (IFS, Cycle 45r1; ECMWF, 2018). Operational ensemble forecasts with 50 perturbed members initialized at 0000 UTC 22 - 27 Sep 2018 and the operational analysis are used. The ECMWF operational ensemble forecast is based on perturbed initial conditions as well as stochastic perturbations of model physics (for details see ECMWF, 2018). The spectral resolution of the operational ensemble is TCO639 (about 18 km) on 91 model levels, and the resolution of the operational analysis is TCO1279 (about 9 km) on 137 model levels. The data are available every 6 h and have been interpolated to a regular grid with a horizontal resolution of $1°$. For the initialization at 0000 UTC 24 Sep 2018, PV on isentropic surfaces (every 5 K) and equivalent potential temperature (θ_e) on pressure levels (every 25 hPa) are additionally computed from the standard variables. As a measure for forecast skill, anomaly correlation coefficients (ACC) are calculated for geopotential height at 500 hPa for each ensemble member of the forecasts initialized at 0000 UTC 24 Sep 2018 and 0000 UTC 27 Sep 2018. As a reference, the daily mean ERA-Interim climatology from 1979-2014 is used

(for details see supplementary material S4). Anomalies of θ_e at 900 hPa were computed with respect to the September/October ERA-Interim climatology from 1979-2017.

In addition, observational data are used from satellite imagery: the infrared channel 9 ($10.8\,\mu$m) of MSG SEVIRI provided by the European Organisation for the Exploitation of Meteorological Satellites (EUMETSAT), and the satellite-based 3-hourly rainfall intensity estimate Tropical Rainfall Measuring Mission (TRMM) level 3 version 7 on a 0.25 degree regular grid. Accumulated rainfall over a nine-hour period is computed using the average intensity (in $\text{mm}\,\text{h}^{-1}$) over three subsequent 3-hourly time intervals, consistent with the computation of the daily accumulated TRMM rainfall estimate.

2.2 Trajectory computation and identification of WCBs

Computing trajectories provides insight into the Lagrangian history of air parcels. In this study, the Langrangian analysis tool LAGRANTO (Wernli and Davies, 1997; Sprenger and Wernli, 2015) is used to identify WCB trajectories in the operational analysis (ascent rate larger than 600 hPa in 48 h, see e.g. Madonna et al., 2014). Intersection points of WCBs with 322.5 and 327.5 K isentropic layer are identified. Additionally, LAGRANTO was used to compute 24 h forward trajectories of low-level air parcels in the 850-950 hPa layer (every 25 hPa) and in a circle with 250 km radius around cyclogenesis position.

2.3 Cyclone phase space and cyclone tracking

The cyclone phase space (CPS; Hart, 2003) is a useful tool to diagnose the thermal structure of cyclones throughout their life cycle. The CPS uses three parameters to define the thermal structure: lower-tropospheric horizontal thermal asymmetry (B), which measures the across-track 900-600 hPa thickness gradient, i.e. frontal nature; and thermal winds in the lower ($-V_T^L$; 900-600 hPa) and upper troposphere ($-V_T^U$; 600-300 hPa), which measure the vertical thermal structure. In this three-dimensional parameter space, cyclones can be classified as frontal ($B > 10$) or non-frontal ($B \leq 10$), cold-core ($-V_T^L < 0$ and $-V_T^U < 0$), hybrid ($-V_T^L > 0$ and $-V_T^U < 0$), or deep warm-core ($-V_T^L > 0$ and $-V_T^U > 0$). In this study, cyclones that at least once in their life cycle fulfill the deep warm-core criterion are classified as medicanes. As the frontal nature of the cyclone is not in the focus of this study, the symmetry parameter B is not considered.

Cyclone tracks at six-hourly temporal resolution are obtained for each of the 50 ECMWF ensemble members and the operational analysis using the cyclone detection and tracking method described by Picornell et al. (2001). This method was specifically designed to study mesoscale cyclones in the Mediterranean Sea, including medicanes (Gaertner et al., 2018). More specifically, six-hourly SLP fields are used to identify pressure minima after applying a Cressman filter (radius of 200 km; Sinclair, 1997) to smooth out noisy features and small cyclonic structures. Weak cyclones are then filtered with a SLP gradient threshold of 0.5 hPa per 100 km. Cyclone tracks are identified with the aid of the horizontal wind field at 700 hPa, which is considered the steering level for cyclone movement. For one member (member 32) the cyclone track showed an unrealistically large jump during the first 6 h and therefore the first track point was removed. The CPS is calculated every 6 h based on the track positions and the CPS values at each time are smoothed using a running-mean filter with a 24-h window. Due to the small

size of medicanes, a radius of 150 km is used to calculate the CPS values, consistent with previous studies (e.g. Gaertner et al., 2018) .

2.4 Normalized PV differences

Ensemble members will be grouped into clusters (see Sect. 4.2). To compare PV of two ensemble clusters at different lead times, it is useful to compute normalized cluster-mean differences (see e.g. Torn et al., 2015):

$$\Delta \mathrm{PV}_{\mathrm{AB}} = \frac{\overline{\mathrm{PV}}_{\mathrm{A}} - \overline{\mathrm{PV}}_{\mathrm{B}}}{\sigma_{PV}} \tag{1}$$

where σ_{PV} is the standard deviation of the full ensemble. Hence, $\Delta \mathrm{PV}_{\mathrm{AB}}$ becomes large when the cluster-mean difference of PV between clusters A and B at a given location is much larger than the ensemble standard deviation at the same location, i.e. when the two clusters contain the members of the ensemble that are most different from each other. Additionally, it allows different lead times to be easily compared. For example, if $\Delta \mathrm{PV}_{\mathrm{AB}}$ increases with lead time, the cluster differences grow faster than the ensemble standard deviation, which means that the clusters become increasingly distinct from each other relative to the full ensemble.

2.5 Statistical significance

In order to be confident that the differences between two ensemble clusters are robust, a two-sided Wilcoxon rank-sum test (Wilks, D. S., 2011) for each cluster pair and considered field was applied. With such a test, the null hypothesis is investigated that it is equally likely that at a certain grid point the value of a randomly picked ensemble member in one cluster is larger or smaller than in a randomly picked ensemble member of the other cluster. When applying such a statistical test to a field, the false discovery rate should be controlled in order to avoid over-interpretation of the results (Wilks, D. S., 2016). This can be done by correcting the p-values of the statistical test taking into account the number of tests. For this study we are only interested in a domain covering the North Atlantic and the Mediterranean and consider a box from 30 to 70°N and 80°W to 30°E. Therefore, a number of 4400 tests are used for the correction. As suggested by Wilks, D. S. (2016), a Benjamini-Hochberg correction is used in this study and the false discovery rate is set to a rather conservative value of $\alpha_{\mathrm{fdr}} = 0.1$ for all analyses. Regions where the null hypothesis is rejected on this level of α_{fdr} can then be used to identify where and when robust differences in the clusters emerge in the ensemble forecast.

3 Synoptic overview

This section provides first an overview of the synoptic situation over the Euro-Atlantic region prior to cyclogenesis of Medicane Zorbas. Then cyclogenesis and the initial intensification of Zorbas are discussed, during which it acquired a deep warm core. A particular focus is given to the link between the evolution of low-level θ_e and upper-level PV, and how it affected the formation of a deep warm core. This is followed by an overview of the full cyclone track and its evolution in the CPS diagram. The section ends with a discussion of satellite-based observations of clouds and precipitation.

The period prior to cyclogenesis was initially, at 0000 UTC 24 Sep 2018 (Fig. 1a), characterized by a large-scale situation with a low-amplitude Rossby-wave pattern in the Euro-Atlantic region including: a large-scale trough (T), a zonally oriented jet maximum (J) over Newfoundland, and a weak ridge (R) and trough (T2) pattern over northern Europe. In the following three days, the wave pattern strongly amplified and anticyclonic Rossby-wave breaking resulted in a pronounced PV streamer (S2) over the Mediterranean at 0000 UTC 27 Sep 2018 (Fig. 1d). In greater detail, this transition can be split into different phases. First, the trough over Newfoundland and the weak ridge / trough pattern over Europe amplified and the jet maximum moved over the North Atlantic and became directed north-eastward towards Iceland (Fig. 1b). Subsequently, the trough (T) elongated into a PV streamer (S), then interacted with Hurricane Leslie (L) resulting in substantial WCB activity over the central North Atlantic (black crosses in Fig. 1c,d). At the same time, anticyclonic wave breaking downstream over Europe was accompanied by a strongly amplified ridge and a narrowing of the trough (T2) over eastern Europe. This narrowing trough finally elongated to a PV streamer (S2) reaching the central Mediterranean Sea at 0000 UTC 27 Sep 2018 (see Fig. 1d). Note that in this case the anticyclonic wave breaking over Europe occurred clearly before the inflow of WCB air masses into the upstream ridge, indicating that the ridge amplification was not driven by the WCB related to Leslie. Hence, even if the presence of an upstream tropical cyclone and enhanced WCB activity are reminiscent of situations with Mediterranean cyclogenesis documented in previous studies (e.g. Pantillon et al., 2013; Raveh-Rubin and Flaounas, 2017), the temporal sequence in this case differs from the evolution documented in these studies.

The formation of the PV streamer over the Mediterranean was followed by cyclogenesis of Zorbas at the PV streamer's south-eastern flank (Fig. 2a-c). At the time of cyclogenesis, the PV streamer broke up resulting in the formation of a PV cutoff (Fig. 2b). In the following 24 h, the cyclone intensified rapidly and at the same time the PV cutoff was eroded, indicating the presence of intense diabatic processes. Shortly before cyclogenesis, the large mean sea level pressure gradient between a surface high pressure system over eastern Europe and a surface low pressure area over the Levantine Sea resulted in strong low-level advection of air with low θ_e across the Aegean Sea (Fig. 2a,d). Air with anomalously high θ_e was present over Lybia and at cyclogenesis in immediate proximity of the cyclone centre (hatched regions in Fig. 2d,e). A detailed trajectory analysis (see supplementary material S1 and Figs. S1, S2, and S3) shows that the low-level air in the 850-950 hPa layer and within a radius of 250 km around the cyclone centre at the time of genesis (1200 UTC 27 Sep 2018) originated from the Aegean and Black Sea and was substantially moistened by sea-surface fluxes as it traveled across the Aegean Sea. This is consistent with the direction of low-level winds (Fig. 2d,e). The relevance of strong advection of cold and dry air and moistening by surface fluxes over the Mediterranean has already been pointed out for medicane formation in the western Mediterranean (Miglietta and Rotunno, 2019). A day after cyclogenesis, a low-level θ_e maximum was present in the cyclone centre (Fig. 2f) and, as diagnosed by the CPS, Zorbas had aquired a deep warm core at this time (Fig. 3b). The system aquired an equivalent barotropic structure, i.e. the surface cyclone, the minimum of geopotential height at 500 hPa, and the upper-level PV maximum were vertically aligned (Fig. 2c,f). The upper-level PV cutoff had decayed and the small-scale PV maximum with values around 2 PVU above the cyclone centre at 1200 UTC 28 Sep 2018 marked the upper end (at around 400 hPa) of a vertically extended diabatic PV maximum above the cyclone centre that was disconnected from the stratospheric high PV reservoir (not shown). The presence of a mid- and low-tropospheric PV maximum and an upper-tropospheric PV minimum are characteristic for many medicanes (Miglietta

et al., 2017).

225   In order for the latent heating to reach high enough and result in the erosion of the upper-level PV anomaly, sufficiently moist and warm air has to be present in the lower troposphere. In this case, values of θ_e at 900 hPa in the cyclogenesis area exceeded 330 K. This value provides the approximate maximum isentropic level to which these air masses can ascend upon the release of latent heat. In fact, more than half of the low-level air parcels with θ_e>322.5 K between 850 and 950 hPa and within a radius of 250 km around the cyclone centre experienced strong diabatic heating and rapid cross-isentropic ascent (average $\Delta\theta$ of about

230 30 K) to levels above 322.5 K during the first 24 h after cyclogenesis (black crosses in Fig. 2c). Hence, they contributed substantially to the erosion of the PV cutoff. Despite limitations of trajectories in convective situations, this is a strong indication that air with high θ_e at low levels was a crucial prerequisite for the erosion of the stratospheric PV anomaly and, as a result, the formation of a deep warm core. In fact, the difference between θ_e at 850 hPa and tropopause potential temperature is a measure of bulk stability and has shown to be a suitable predictor for the occurrence of tropical transition (so-called coupling index,

235 McTaggart-Cowan et al., 2015). The formation of a deep-warm core system with an equivalent barotropic structure as result of high θ_e values at low tropospheric levels and strong latent heating will become relevant again towards the end of this article when cyclone evolutions in the ensemble members are discussed.

    The full track of Medicane Zorbas and the CPS diagram are shown in Fig. 3. Zorbas formed at 1200 UTC 27 Sep 2018 close to Benghazi, moved into the central Mediterranean and then sharply turned eastward and moved over Greece into the Aegean Sea,

240 where it finally decayed four days after its formation. According to the CPS (Fig. 3b), Zorbas formed as a cold-core cyclone and within 18 h acquired a deep warm core that was sustained for more than three days. This is comparable to the January 1982 case investigated by Picornell et al. (2014), but substantially longer than for all eight recent medicane cases investigated by Di Muzio et al. (2019) using the same dataset. Zorbas reached its maximum intensity (992 hPa) already 12 h after cyclogenesis. Note that the ECMWF reported a substantial underestimation of the cyclone intensity in the operational analysis in the later

245 stage of its life cycle and estimated the real central pressure at land fall to below 990 hPa (ECMWF, 2019). Satellite images indicate the formation of an eye-like feature shortly before Zorbas reached Greece on 29 Sep 2018. However, this aspect of the life cycle is beyond the scope of this study. Instead, we focus on the synoptic aspects prior to cyclogenesis, and the subsequent about one-day period of initial cyclone intensification and the formation of a deep warm core.

    An observational perspective on this period is given in Fig. 4a. Shortly after cyclogenesis, intense rainfall was observed north-

250 west of the cyclone centre (Fig. 4a). The dense cloud patch and lightning activity (see supplementary material S2 and Fig. S4) at this location are indicative of strong latent heating and the presence of deep convection, which helps to explain the rapid diabatic erosion of the PV cutoff (see e.g. Wirth, 1995; Portmann et al., 2018). At the time when the deep warm core was well established, precipitation was weaker and mainly located near the cyclone centre (Fig. 4b). The cloud structure indicates that well-defined fronts were absent, i.e. Zorbas acquired a more tropical-like appearance. According to TRMM, Zorbas lead to

255 extreme daily rainfall (>99th precentile, reference period 1998-2018) in several regions in Lybia, the central Mediterranean, and Greece (not shown).

    The evolution of Zorbas in the first day of its life cycle to some extent agrees well with the climatological evolution found for the strongest Mediterranean cyclones around their time of maximum intensity (Flaounas et al., 2015), in particular considering

the shape and cyclone-relative position of the upper-level PV anomaly. Remarkable in this case were the anomalously high values of low-level θ_e. Further, the evolution bears similarities to the first phase of tropical transition events of strong extratropical cyclones as discussed in Davis and Bosart (2003, 2004). In particular, this is true for the enhanced rainfall to the west of the cyclone centre, the diabatic erosion of the stratospheric PV anomaly, the emergence of a tropospheric PV maximum above the storm centre, and the formation of an equivalent barotropic vortex.

4   ECMWF ensemble forecasts

4.1   Ensemble spread and uncertainties in the formation of the Mediterranean cyclone

The synoptic situation over the Mediterranean was associated to substantial uncertainties in the operational ECMWF ensemble forecasts until three days before genesis of Zorbas. This is shown by the evolution of the ensemble spread of geopotential height at 500 hPa averaged over the Mediterranean at 0000 UTC 27 Sep 2018 for different initializations (Fig. 5). For initializations until 0000 UTC 24 Sep 2018, the spread decreases only marginally, but it does so very rapidly for later initializations, indicating a forecast jump. For the initialization at 0000 UTC 24 Sep 2018, all ensemble members develop a surface cyclone but uncertainties related to cyclone position and its thermal structure were substantial (Fig. 6, ignore colors for this paragraph). While several members forecasted the cyclone approximately at the correct location over Lybia at 1200 UTC 27 Sep 2018, some forecasted it too far to the west and some too far north-east in the Aegean Sea (Fig. 6a). This zonal position uncertainty was still present 24 h later, and in four members the cyclone already disappeared (Fig. 6b). At this time, the uncertainty in the position of the cyclone ranged from the coast of Tunisia to the Aegean Sea. Interestingly, in the members where the cyclone formed too far to the north-east, genesis occurred more than 12 h before cyclogenesis in the operational analysis. In addition, only 21 ensemble members predicted a medicane (markers with white centres), and among the medicanes differences in their strength (as defined by the maximum $-V_T^U$ value, see marker size) were substantial. Also there was a clearly preferred region for medicanes to form near the Lybian coast, indicating a link between cyclone position and thermal structure. Considering these large uncertainties associated with the prediction of the surface cyclone and the timing of the forecast jump we focus in the following on the 0000 UTC 24 Sep 2018 initialization. The rest of this article aims to show that these uncertainties in the Mediterranean cyclone evolution can be traced back to a short-wave perturbation in the wave guide over the North Atlantic initiated during the first day of the forecast, and that they were tightly linked to uncertainties in the position of the Mediterranean PV streamer.

4.2   Ensemble clustering according to the position of the PV streamer

In the 0000 UTC 24 Sep 2018 forecast, the ensemble spread at 0000 UTC 27 Sep was mainly enhanced because of a substantial uncertainty in the position of the Mediterranean PV streamer. To establish the dynamical link between, on the one hand, the uncertainties in the position and thermal structure of the cyclone and, on the other hand, upstream uncertainties along the

North Atlantic wave guide, a pragmatic clustering procedure is presented that allows separating the ensemble members of the 0000 UTC 24 Sep 2018 initialization into three clusters based on the position of the Mediterranean PV streamer (S2 in Fig. 1d), at day 3 of the forecast. The three identified PV streamer scenarios, i.e. clusters, are the basis for all remaining analyses. For the clustering, a box is defined around the PV streamer identified at 0000 UTC 27 Sep 2018 in the operational analysis (Mediterranean box, 5-30°E, 30-45°N, see black box in Fig. 7). The clustering uses vertically averaged PV between 320 and 330 K, hereafter called PV_{av}. Before averaging, all PV values with PV<2 PVU are set to zero to remove the contribution of the variability of tropospheric PV values. Hence, PV_{av} is high in areas where the stratospheric PV streamer is strong and deep, and low where it is weak and shallow. The pragmatic clustering is then based on two different steps: First, from all 50 ensemble members the ones are identified for which the area with $PV_{av} \geq 2$ PVU in the box has more than 75% overlap with the corresponding area in the analysis. In these 19 members, the streamer has a similar location as in the analysis, i.e. a central position in the ensemble, and is therefore referred to as cluster C (see blue shading in Fig. 7). The remaining members are separated into two clusters depending on whether the maximum PV_{av} is shifted to the west (cluster W, 12 members, green shading in Fig. 7) or east (cluster E, 18 members, red shading in Fig. 7) relative to the analysis. There is one ensemble member that cannot be attributed to one of the three clusters because its overlap is less than 75% but the maximum of PV_{av} is located at the same longitude as in the analysis. The histogram of the longitude where the maximum of PV_{av} occurs between 36-37°N (inset in Fig. 7) shows three clearly distinct peaks, one for each cluster, validating the simple clustering approach. There are a few borderline members but they do not affect the main results of this study.

The meaningfulness of this clustering for studying the predictability of this case is further supported by the fact that it helps to explain the temporal development of the ACC averaged in the Mediterranean box. As shown in Fig. 8a, the ACC of geopotential height at 500 hPa in the Mediterranean starts to decrease in the majority of the ensemble members at the time when the PV streamer reaches the Mediterranean on 26 Sep 2018 and even more after cyclogenesis occurs, while it remains high (close to 1) until 29 Sep 2018 for most members of cluster C (blue lines in Fig. 8a). After the decrease from 1 to around 0.8 the median ACC (red line) remains fairly constant until 29 Sep 2018. In comparison, for the ensemble forecast initialized at 0000 UTC 27 Sep 2018, i.e. at the time when the PV streamer has developed, the ACC remains high in all members during the intensification and deepest phase of Zorbas, decreasing only after 29 Sep 2018 (Fig. 8b), likely due to errors associated with a second PV streamer reaching the Mediterranean in the northern part of the box (not shown). It can be concluded that errors in the position of the PV streamer limited the large-scale predictability as measured by the ACC of geopotential height on 500 hPa in the Mediterranean, and that cluster C contains the members with the most accurate forecasts.

5 Propagation of forecast uncertainties along the North Atlantic wave guide to the Mediterranean cyclone

In the following the chain of dynamical processes is described that enabled the propagation and amplification of forecast uncertainties (here considered as significant differences between cluster means) from the North Atlantic to the Mediterranean cyclone. In a first step, the propagation and amplification along the North Atlantic wave guide to the Mediterranean PV streamer

is considered. Subsequently, the effect of the uncertain PV streamer position on the cyclone position and thermal structure is discussed.

5.1 Propagation from North Atlantic jet streak to Mediterranean PV streamer

325 To investigate the first part of the uncertainty propagation, normalized PV differences (see Sect. 2.4) on 325 K between clusters E and W (ΔPV_{EW}) are analyzed, as these are the clusters that deviate the most in terms of the PV streamer position.

During the first day after initialization a short-wave pattern of positive and negative PV differences emerges on the North Atlantic wave guide from an initially very spotty difference field (Fig. 9 and Fig. 10a). At 1800 UTC 24 Sep 2018 (Fig. 9d), a positive ΔPV_{EW} with an amplitude larger than 1.5 standard deviations appears on the stratospheric side of the strengthening jet

330 streak (yellow contour, marked as J in Fig. 1b six hours later) over the North Atlantic. This ΔPV_{EW} is statistically significant (as indicated by the white contours in Fig. 9d; see Sect. 2.5 for the testing procedure). Visually, it seems that this ΔPV_{EW} emerges from a small positive ΔPV_{EW} that propagates from the Gulf of Saint Lawrence at initial time (Fig. 9a) along the wave guide into the North Atlantic and amplifies (indicated by the grey boxes in Fig. 9). However, these differences early in the forecast are not statistically significant and other sources of the amplified ΔPV_{EW} at 1800 UTC 24 Sep 2018 could

335 be relevant. However, a substantial contribution of moist diabatic processes and WCBs can be excluded. There is no WCB activity in the area indicated by the grey boxes (as shown in Fig. 1a,b) and precipitation is weak (not shown). Furthermore, the positive ΔPV_{EW} is located in a region that is clearly stratospheric, i.e. where strong moist diabatic processes are mostly absent. Hence, it is most likely dry upper-tropospheric dynamics associated to the jet streak that drives the amplification of the initial condition differences until 1800 UTC 24 Sep 2018. This is in contrast to studies that showed the relevance of diabatic

340 low-PV outflow for the emergence of upper-tropospheric forecast uncertainties (e.g. Pantillon et al., 2013; Grams et al., 2018), but in agreement with studies emphasizing the important role of non-linear (barotropic) upper-tropospheric dynamics for the amplification of forecast error and ensemble spread (Baumgart et al., 2018; Baumgart and Riemer, 2019). An analysis of geostrophic and ageostrophic wind differences between the two clusters and PV gradients in the jet streak region suggests that differences in the ageostrophic circulation are important for this amplification (not shown). A physically plausible cause of the

345 strengthening of the jet streak and the initiation of a wave-like perturbation at the tropopause at 1800 UTC 24 Sep 2018 could be the band of high stratospheric PV that is approaching the tropopause region (green contours in Fig. 9d).

After the emergence of the prominent positive ΔPV_{EW} at 1800 UTC 24 Sep 2018, there is a clear downstream propagation as a short-wave perturbation along the wave guide into the Mediterranean (Fig. 10). First, the positive ΔPV_{EW} further amplifies and a negative ΔPV_{EW} emerges downstream over the North Sea (Fig. 10a). The deviation of the 2 PVU contours shows an

350 increasingly clear short-wave pattern superimposed onto the large-scale Rossby-wave pattern, which is particularly pronounced in cluster W (dashed contour). The positive (negative) ΔPV_{EW} are associated with cyclonic (anticyclonic) difference winds, resulting in a westward phase speed of the ΔPV_{EW} relative to the mean flow. During the propagation, the maximum amplitude of the ΔPV_{EW} pattern increases and moves downstream, consistent with an eastward group speed. For example, at 0000 UTC 25 Sep 2018, the negative ΔPV_{EW} has a smaller amplitude than the positive ΔPV_{EW} (Fig. 10a), while 12 hours later, the

355 opposite is the case (Fig. 10b). The downstream development of the ΔPV_{EW}-wave subsequently results in a more progressed

anticyclonic Rossby-wave breaking in cluster W compared to cluster E (Fig. 10c,d). Ultimately, this leads to a zonally shifted tip of the narrow trough (T2 in Fig. 1d), and later, the PV streamer (S2 in Fig. 1c), as seen in Fig. 10d and consistent with Fig. 7.

Similar wave-like patterns of PV errors that emerge from a localized PV error and result in downstream development have been shown for example by Davies and Didone (2013) and Baumgart et al. (2018) for time scales of several days. Here such a pattern is shown for significant PV differences between ensemble clusters with a particularly fast propagation from 40°W to 20°E in less than two days. The initiation of the short-wave perturbation by a stratospheric PV anomaly is consistent with results of idealized studies looking at the effect of an isolated vortex approaching a strong isentropic PV gradient from the stratospheric side (e.g. Schwierz et al., 2004) and of a climatological analysis of precursor perturbations of Rossby-wave initiation events (Röthlisberger et al., 2018, their Fig. 13). It is also reminiscent of the situation depicted by Davies and Didone (2013), in which a lower-stratospheric PV anomaly approaches the wave guide, initiates a wave train at the tropopause, and results in forecast error growth. In our case, this initiation occurs around 1800 UTC 24 Sep 2018, consistent with the large drop of ensemble spread over the Mediterranean from the 0000 UTC 24 Sep to the 0000 UTC 25 Sep 2018 initializations (Fig. 5). This indicates that, once the short-wave perturbation was captured properly, the forecast uncertainties associated with the formation of the PV streamer were substantially reduced.

5.2 How the uncertain PV streamer position affects the cyclone's position and thermal structure

In the previous section the dynamical pathway leading to the uncertainty in the position of the PV streamer at day 3 of the ensemble forecast was analyzed. In the following, it is investigated how this uncertainty affects the position and thermal structure of the resulting Mediterranean cyclone.

First, the effect of the PV streamer position on the subsequent PV cutoff formation and surface cyclogenesis is discussed based on synoptic composites for the three clusters (Fig. 11). At 1200 UTC 26 Sep 2018, the position and shape of the PV streamer in cluster C is still very close to the analysis, whereas in cluster W the tip of the streamer is thinner and extends more to the west, and in cluster E it is shifted to the east (Fig. 11a,d,g), consistent of course with the previous discussion of Fig. 10d. In these regions, clusters W and E significantly differ from cluster C. This is not surprising as the clustering was specifically designed to focus on these differences. One day later, at the time of cyclogenesis in the analysis, a PV cutoff has formed in all clusters (Fig. 11b,e,h) and the differences in the scenarios over the Mediterranean are very prominent. While in cluster C the cutoff is located south of Italy in the Central Mediterranean (in excellent agreement with the analysis), cluster W exhibits a much weaker cutoff further to the west over Tunisia, and cluster E shows a stronger cutoff shifted to the east. In all clusters the developing surface cyclones are located slightly east of the cutoff (cyclone centres of individual ensemble members are shown by black dots). Hence, in cluster C the cyclones are located close to Benghazi (as in the analysis, indicated by the teal star), in cluster W too far west close to Tripoli, and in cluster E too far east over Crete. One day later (Fig. 11c,f,i), the PV cutoff in the analysis has decayed into smaller patches. In cluster C, the cutoff has clearly weakened (PV values < 3 PVU), in cluster W it has fully decayed, and in cluster E it is still very prominent and strong (PV values > 6 PVU), indicating substantial differences in the diabatic modification of the PV cutoff. In cluster W, the absence of clear structures at 1200 UTC 28 Sep 2018

in both upper-level PV and SLP is partially due the large variability within this cluster at this time. Note again that the cyclone positions in cluster C agree well with the analysis whereas cyclones in the other clusters have typical position errors of about 400 km.

The surface cyclones in cluster E show a very different behavior than in clusters W and C. First, cyclogenesis occurs earlier and takes place in a pre-existing low pressure area over the Levantine Sea (Fig. 10g). At 1200 UTC 26 Sep 2018, 9 out of 18 members in this cluster have a cyclone identified in the Levantine Sea close to Cyprus, whereas in clusters W and C most cyclones form later in the southern part of the Central Mediterranean Sea, in immediate proximity to the PV cutoff. Second, cyclones in cluster E are on average much weaker than in clusters W and C (see box plots in the individual panels of Fig. 11). The pre-existing cyclones over the Levantine Sea deepen slightly when they interact with the PV cutoff but – with the exception of two cases – weaken again afterwards. It can be concluded that the eastward shift of the PV streamer (cluster E) leads to a particularly strong non-linear response of the cyclogenesis process and the intensity evolution of the surface cyclone. The westward shift (cluster W) also results in weaker surface cyclones than in cluster C and in the analysis, some with very short life times. Hence, the uncertain surface cyclone positions after 84 h and 106 h of the forecast, i.e. at the time of cyclogenesis and after one day of intensification, can be clearly attributed to the uncertain PV streamer position and the cluster with the best representation of the PV streamer position (cluster C) results in the best forecast of the cyclone position and intensity. This is also supported by Fig. 6, now referring the colors. They indicate the cluster to which each cyclone belongs, in full agreement with the aforementioned conclusion.

Interestingly, cluster E also results in the lowest fraction of medicanes (i.e. cyclones with a deep warm core, 2 out of 18 members), followed by cluster W (6 out of 19 members) and cluster C (15 out of 19 members). Hence, uncertainties in the cyclone's position are also tightly linked to uncertainties in its thermal structure. Because this link is not straightforward, the final part of this section sketches a potential pathway of how the uncertainties of the positions of the PV streamer and cyclogenesis affect the cyclone's thermal structure. This analysis will not be complete, because, as uncertainties reach the mesoscale, we approach the limits of the dataset and methods used in this study. Nonetheless, key elements of this link can be identified by focusing on mesoscale differences at upper- and lower levels between cyclones in the three clusters during the 24 h period after cyclogenesis in the operational analysis. As shown in Fig. 11f,i, in both clusters C and E, the upper-level PV cutoff and the surface cyclone are vertically aligned at 1200 UTC 28 Sep 2018, i.e. they have an equivalent barotropic structure. Despite this similarity, the vertical thermal structure of the vortices strongly differs between the clusters. As diagnosed by the CPS, most cyclones in cluster E have an upper-level cold core, whereas most cyclones in cluster C have an upper-level warm core. In essence, this means that the stratospheric PV anomaly in cluster C is eroded (as in the operational analysis, see Sect. 3), while it remains strong in cluster E (consistent with Fig. 11f,i). In the operational analysis, the erosion of the stratospheric PV anomaly was linked to low-level air masses with high θ_e values near the cyclone centre. Here, we investigate the hypothesis that the differences in the modification of the stratospheric PV anomaly in cluster C and E, i.e. the differences in the thermal structure of the cyclones, are linked to differences in low-level θ_e values. To this aim, θ_e between 850 and 950 hPa is averaged within a 250 km radius around the cyclone centre (as shown for the analysis in Fig. 2e), to quantify low-level θ_e values for each cyclone. Then, the 50% grid points with the highest PV values on 325 K within a radius of 750 km around the cyclone centre are averaged to quantify the

amplitude of the upper-level PV anomaly. The larger radius is chosen because the highest values of upper-level PV are usually expected west of the cyclogenesis area. Each cyclone can now be positioned in a diagram of low-level θ_e vs. upper-level PV. Figure 12a,b shows such diagrams with the same markers and for the same times as the geographical maps in Fig. 6. At cyclogenesis time in the operational analysis, cluster C and W are positioned at moderate upper-level PV values around 2-3 PVU and high low-level θ_e values between 320 and 330 K (Fig. 12a). Cluster W cyclones tend to have slightly lower upper-level PV and low-level θ_e values than cluster C. The operational analysis (black marker) is positioned fairly well within cluster C, albeit with relatively high upper-level PV. Cluster E cyclones exhibit higher upper-level PV values around 3-5 PVU and much lower low-level θ_e values around 310-315 K. These values are most likely lower because cyclogenesis occurred closer to the northern coast of the Mediterranean and as a result, the low-level air parcels were less exposed to the sea surface and therefore less strongly moistened by latent heat fluxes compared to the operational analysis (for a discussion of the Lagrangian history of the low-level air parcels at cyclogenesis in the operational analysis, see supplementary material S1 and Figs. S1, S2, and S3). Within the following 24 h period of initial intensification and deep warm core formation in the operational analysis, cyclones in cluster C and W are associated with substantial precipitation in their centres (Fig. 12c) and a reduction of upper-level PV by about 1 PVU and of low-level θ_e by about 3-5 K (Fig. 12b), similar to the operational analysis. This behavior is consistent with the crucial role of low-level air with high θ_e values for strong latent heating and cross-isentropic upward transport and the resulting erosion of the upper-level PV anomaly as observed in the operational analysis (see Sect. 3). Cyclones in cluster E experience an increase in low-level θ_e and, most importantly, little change in upper-level PV. The slight increase in upper-level PV values in some members might be related to the vertical alignment of the upper-level PV anomaly and the surface cyclone, such that a larger fraction of the upper-level PV anomaly is located within the 750 km radius around the cyclone. Precipitation in the cyclone centre is much lower for the cyclones in cluster E compared to those in clusters C and W and the operational analysis, indicating much lower latent heating. As an interesting side remark, we note that also cluster C underestimates the precipitation in the cyclone centre compared to TRMM rainfall along the cyclone track in the operational analysis.

In summary, the results of this analysis are consistent with the following pathway of the forecast uncertainties: In some members, the eastward displacement of the PV streamer position resulted in a strongly north-eastward shifted cyclogenesis position, leading to a reduced supply of low-level air with high θ_e in the cyclogenesis region. This resulted in weaker and shallower latent heating and no diabatic PV destruction on the 325 K level. Hence, the PV anomaly was not eroded and no upper-level warm core could form. There are a few ensemble members that deviate from this storyline and they are mostly the ones where the PV streamer position is on the borders between the clusters.

6 Summary and conclusions

The basis of this study was an ECWMF operational ensemble forecast, which, after three days, showed large uncertainties in the position of a PV streamer over the Mediterranean and the subsequent development of Medicane Zorbas in September 2018. These uncertainties were substantially smaller for later forecast initializations, motivating the use of this forecast to gain insight into the dynamics behind this rapid decrease in forecast uncertainty. The ensemble members were clustered into three

distinct scenarios according to the position of the PV streamer at day 3 of the forecast. The differences between these scenarios were used to present a dynamical pathway of forecast uncertainties from a short-wave perturbation in the North Atlantic wave guide to the PV streamer and, as a result, the position and thermal structure of the Mediterranean cyclone.

Significant PV differences between clusters appeared first after 18 h forecast time at the stratospheric side of a strengthening jet streak over the North Atlantic. At the same time, a short-wave perturbation superimposed to the large-scale Rossby-wave pattern was initiated. The initiation of the short-wave perturbation and the strengthening of the jet streak was likely linked to a stratospheric high-PV filament approaching the tropopause region. A significant moist diabatic contribution to the emergence of these significant differences between ensemble clusters could be excluded. To some extent, they could be traced back to differences in the initial conditions on the stratospheric side of the upper-level jet streak over the Gulf of Saint Lawrence. Subsequently, the differences further amplified and propagated along the wave guide across the Atlantic into the Mediterranean, where large-scale anticyclonic Rossby-wave breaking occurred and triggered the formation of Zorbas. The propagation and amplification of the upstream PV differences resulted in substantial differences in the position of the PV streamer between the three clusters.

The shift in the position of the PV streamer resulted in a shift of cyclogenesis and it affected cyclone intensity and the diabatic modification of the upper-level PV anomaly. Ensemble members with the central (i.e. correct) PV streamer position produced cyclones most similar to the operational analysis, while the westward shift of the PV streamer resulted in slightly weaker cyclones too far to the west. In both cases, the upper-level PV anomaly was eroded. The eastward shift of the PV streamer led to a particularly different evolution. In this scenario, the cyclones formed too far north-east and more than 12 h earlier from a pre-existing surface trough in the Levantine Sea. The cyclones remained much weaker and the upper-level PV anomaly much stronger than in the operational analysis and most other ensemble members. The non-linear response of the eastward shift of the PV streamer is particularly interesting as a comprehensive analysis of the predictability of PV streamers over the North Atlantic and the Mediterranean in the ECWMF ensemble forecasts showed that there is a tendency for eastward displacement in the forecasts compared to the analysis (Wiegand and Knippertz, 2014).

The central PV streamer position provided the best synoptic conditions for the formation of a strong medicane, while the eastward shift resulted in conditions that were most unfavorable for medicane formation. A reason for this was that the eastward shifted cyclogenesis resulted in significantly reduced low-level equivalent potential temperatures in the cyclogenesis region, which prevented latent heating to be strong enough and reach high enough to erode the upper-level PV anomaly and form an upper-level warm core.

These results extend our understanding of the role of upstream uncertainties for the formation of Mediterranean cyclones in general, and medicanes in particular. It also provides new insight into the dynamics behind forecast jumps as identified by Di Muzio et al. (2019). Previous studies investigating medium-range predictability of Mediterranean cyclones mainly focused on the role of upstream uncertainties related to extratropical transition over the North Atlantic and the associated strong diabatic outflow (e.g. Pantillon et al., 2013). This study showed that the uncertainties limiting the predictability of Medicane Zorbas a few days ahead were mainly linked to dry dynamics in the tropopause region. It highlights the key role of a short-wave perturbation superimposed to the large-scale Rossby-wave pattern that was initiated in a jet streak over the North Atlantic.

Further research is needed to quantify how frequently this mechanism leads to an upstream influence on the medium-range predictability of Mediterranean PV streamers and, as a result, Mediterranean cyclogenesis.

495 Building on studies that document the relevance of upper-level PV anomalies for medicane development (e.g. Miglietta et al., 2017), this study also provides the first analysis of the relevance of an uncertain PV streamer position for uncertain medicane formation and supports the hypothesis that certain regions are more conducive to medicane development than others (Di Muzio et al., 2019). However, once the uncertainties reach the mesoscale, other factors can play an important role that were not analyzed in this study, such as for example vertical wind shear and mid-tropospheric humidity (Tous and Romero, 2013), or

500 the details of the convective processes. In a subsequent study we plan to investigate how these factors limit the predictability of Zorbas' life cycle after cyclogenesis has occurred.

Data availability. All data is available from the authors upon request

Author contributions. JJGA sparked this work by discovering the uncertain forecast of Zorbas. RP and JJGA designed the basic idea for the study. JJGA carried out the analyses that required the computation of the CPS and RP all remaining analyses. MS and HW helped with the

505 data access and handling and gave important guidance and useful inputs during the whole project. RP and JJGA prepared the manuscript and all authors gave critical feedback that helped to improve the article.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. We thank the reviewers Ron McTaggart-Cowan, Florian Pantillon, and one anonymous reviewer for their detailed feedback that helped to improve the manuscript. RP acknowledges funding from the ETH research grant ETH-0716-2. JJGA has been funded

510 through the PhD-grants BES-2014-067905, EEBB-I-18-12841 for short research stays, and by grant CGL2017-89583-R by the Spanish Ministry of Science, Innovation and University, and co-funded by the European Social and Regional Development Funds. We thank Andries de Vries (ETH Zürich) for providing support with the TRMM data and assessing the extremeness of precipitation. RP thanks Emmanouil Flaounas (ETH Zürich) for insightful discussions about medicanes.

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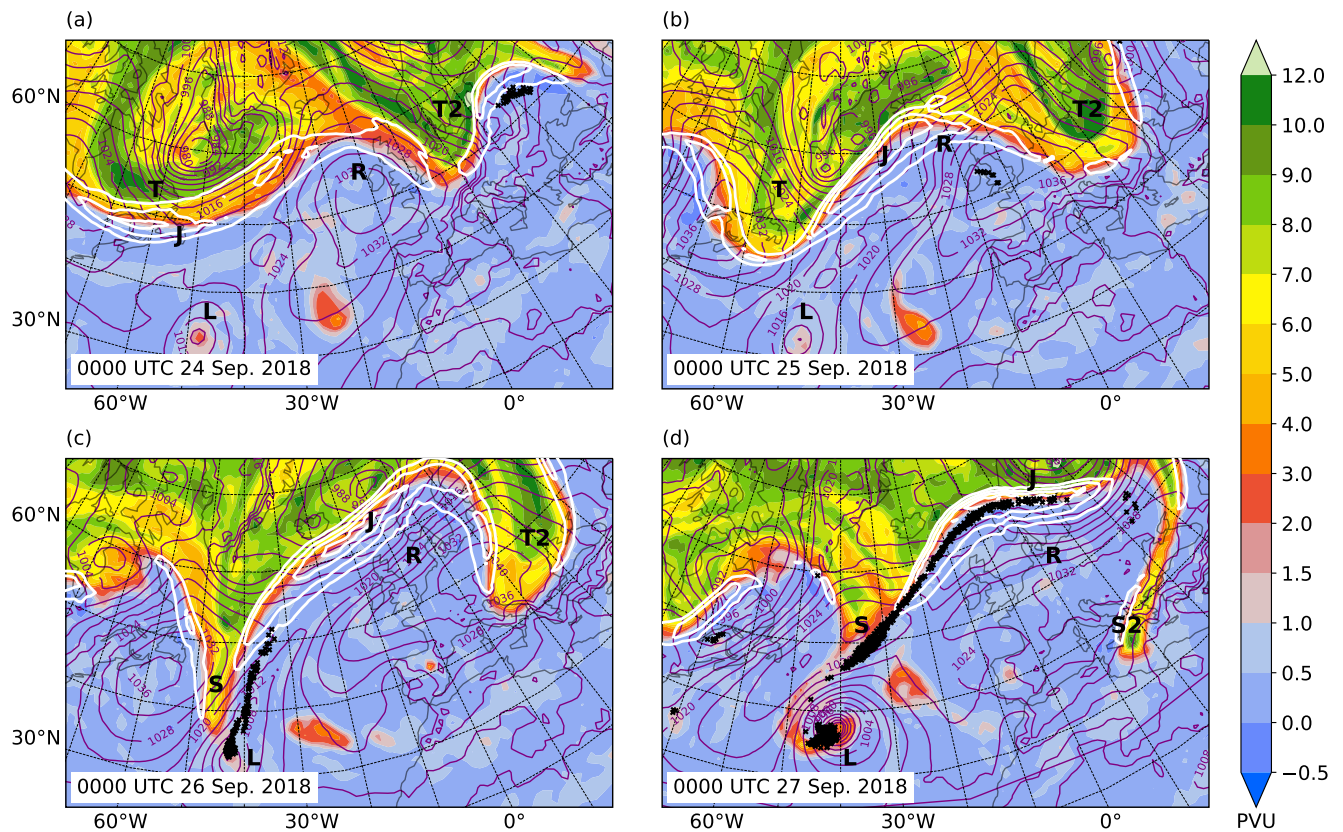
Figure 1. Synoptic situation over the Euro-Atlantic sector before the formation of the PV streamer over the Mediterranean. PV (shaded, in PVU) and wind speed (white contours, in m s^{-1}) on 325 K, intersection points of warm conveyor belts (ascent rate of more than 600 hPa in 48 h) with the 325 K isentrope (black crosses), and sea level pressure (purple contours, in hPa) at (a) 0000 UTC 24 Sep 2018, (b) 0000 UTC 25 Sep 2018, (c) 0000 UTC 26 Sep 2018, and (d) 0000 UTC 27 Sep 2018.
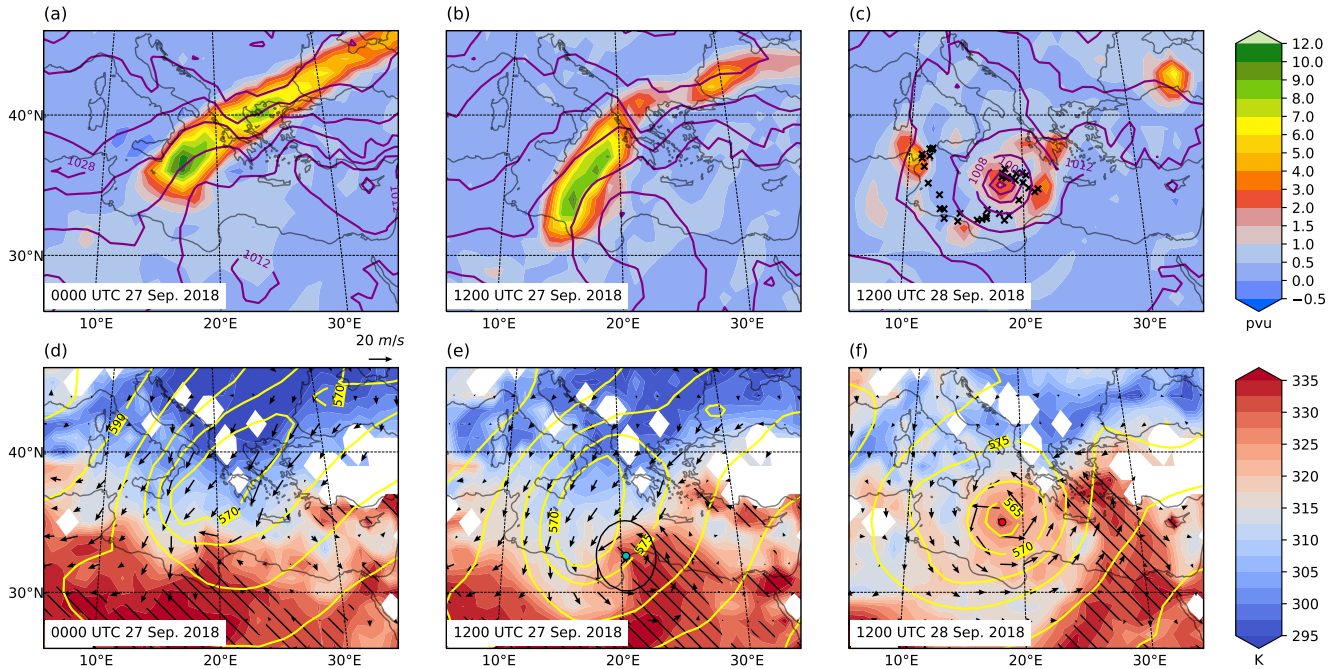
Figure 2. Synoptic situation over the Mediterranean after the formation of the PV streamer at (a,d) 0000 UTC 27 Sep 2018, (b,e) 1200 UTC 27 Sep 2018, and (c,f) 1200 UTC 28 Sep 2018. (a-c) PV on 325 K (shaded, in PVU; this level corresponds approximately to the 300-350 hPa pressure levels in this region), sea-level pressure (purple contours, in hPa); (d-f) equivalent potential temperature (θ_e, shaded, in K) and wind vectors (black arrows, reference vector above panel d) on 900 hPa, and geopotential height on 500 hPa. The hatched areas in (d-f) show regions where θ_e on 900 hPa is anomalously high (at least one standard deviation larger than climatology) with respect to the Sep-Oct ERA-Interim climatology for the period 1979 – 2017. In panel (e), the black circle denotes the circle with radius 250 km from which forward trajectories were started for air parcels with θ_e>322.5 K and black crosses in panel (c) show the position of these air parcels 24 h later if they reached above the 322.5 K level.
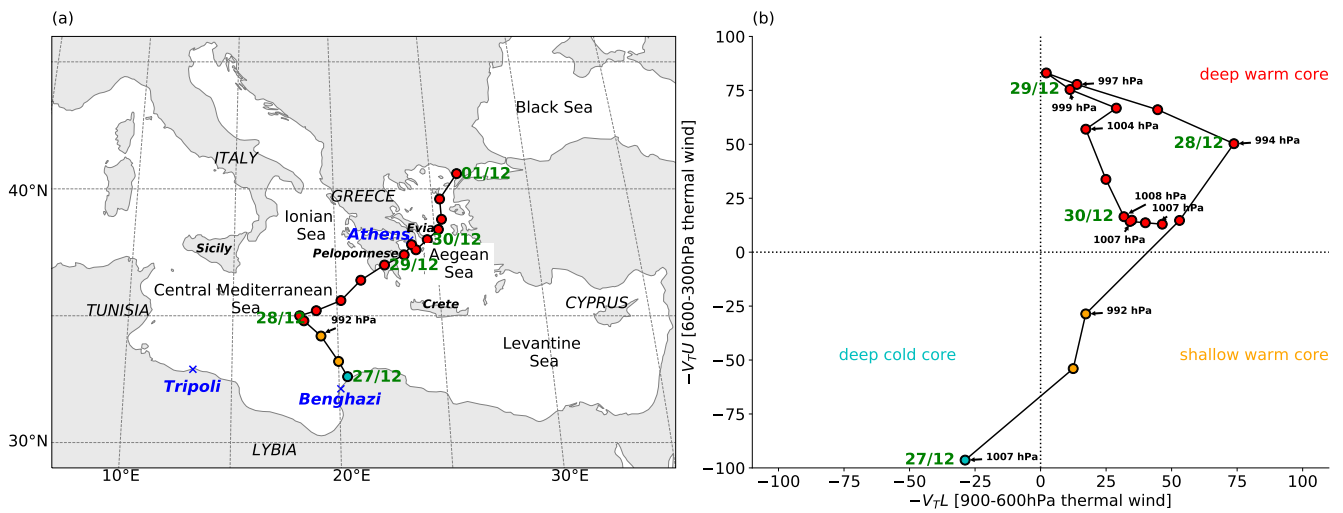
Figure 3. (a) Track of Medicane Zorbas (circles and black line) and (b) cyclone phase space diagram derived from the ECMWF operational analyses at six-hourly intervals in each panel. Cyclone positions are colored according to the quadrant in the CPS diagram (blue: cold core, orange: shallow warm core, red: deep warm core). Black numbers indicate the minimum sea-level pressure (hPa) of the cyclone at this particular time of its life cycle and green numbers the day (in Sep 2018) and time (hours UTC) .
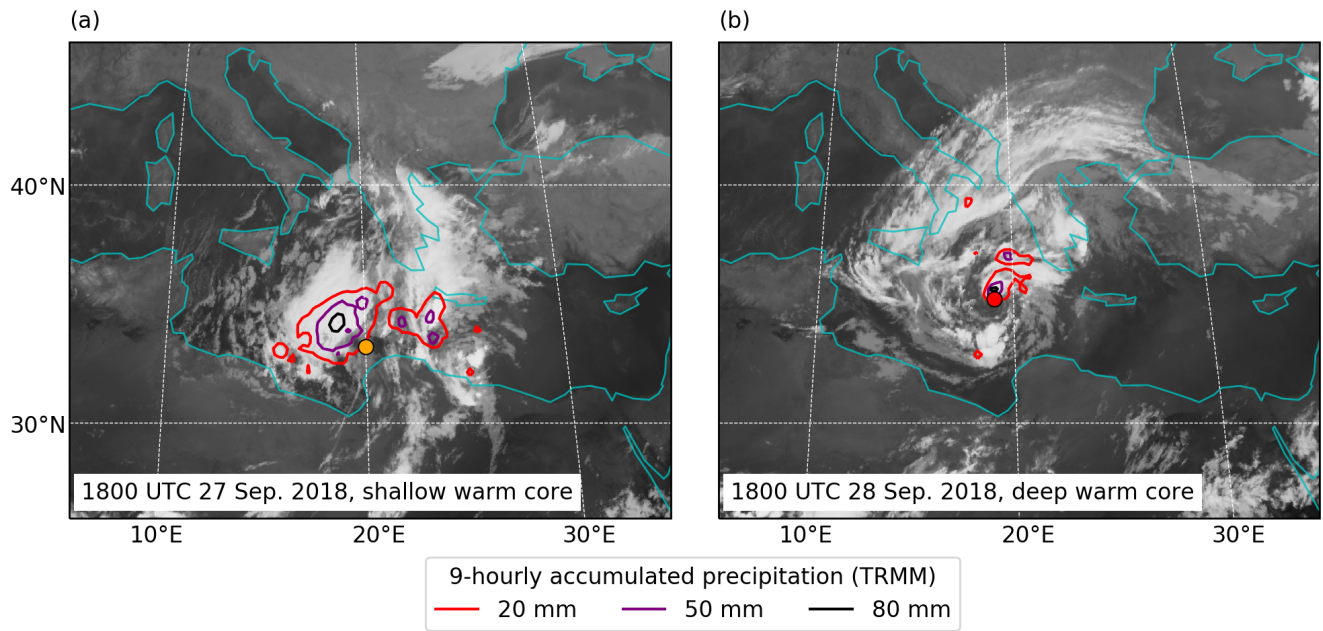
Figure 4. Infrared channel 9 (10.8 μm) of MSG SEVIRI provided by EUMETSAT (grey shading) and 9-hourly accumulated TRMM precipitation during the period centered at the indicated time (colored contours) at (a) 1800 UTC 27 Sep 2018 and (b) 1800 UTC 28 Sep 2018. Cyclone positions based on the ECMWF operational analysis are marked with circles and colored according to the thermal structure of the cyclone (as in Fig. 3).
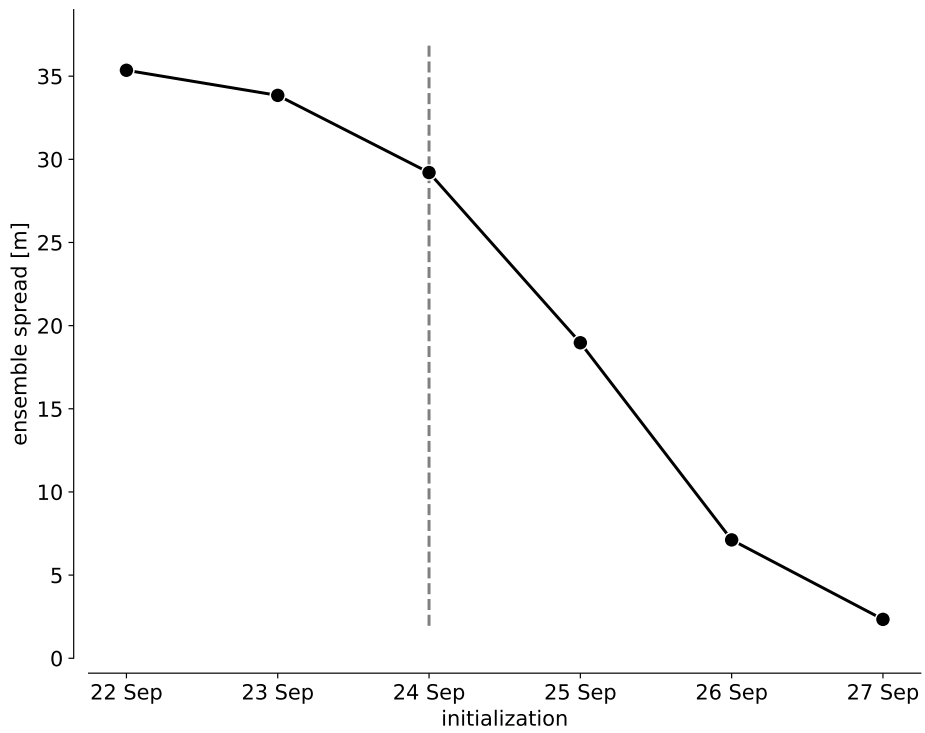
Figure 5. Area-averaged ensemble spread of geopotential height at $500\,\text{hPa}$ at 0000 UTC 27 Sep 2018 in a box over the Mediterranean (5-30°E, 30-45°N, see black box in Fig. 7) for initializations at 0000 UTC 22 Sep to 27 Sep 2018 (black dots and line). The initialization used in this study (0000 UTC 24 Sep 2018) is indicated by the dashed line.
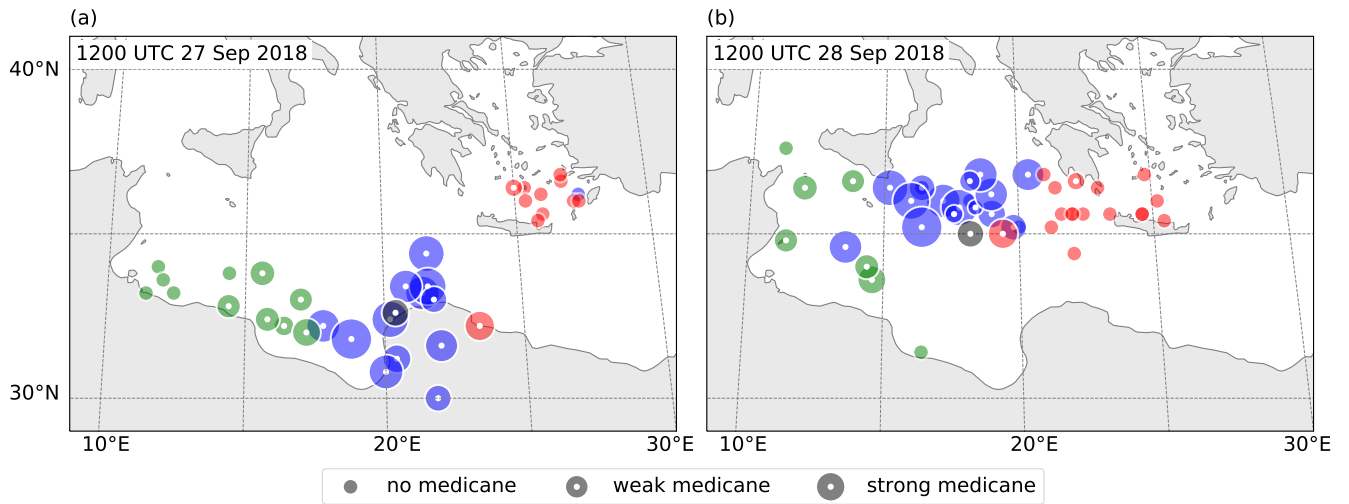
Figure 6. Geographical maps of cyclone positions in all ensemble members and the operational analysis at (a) 1200 UTC 27 Sep 2018 and (b) 1200 UTC 28 Sep 2018. For medicanes (markers with white centre), the size of the markers is proportional to the maximum intensity of the upper-level warm core (see legend for comparison, no medicane: $-V_T^U <= 0$, weak medicane: $-V_T^U = 50$, strong medicane: $-V_T^U = 100$). Marker colors indicate the cluster to which the cyclone belongs (W: green, C: blue, E: red, analysis: black; discussed in Sect. 5.2)
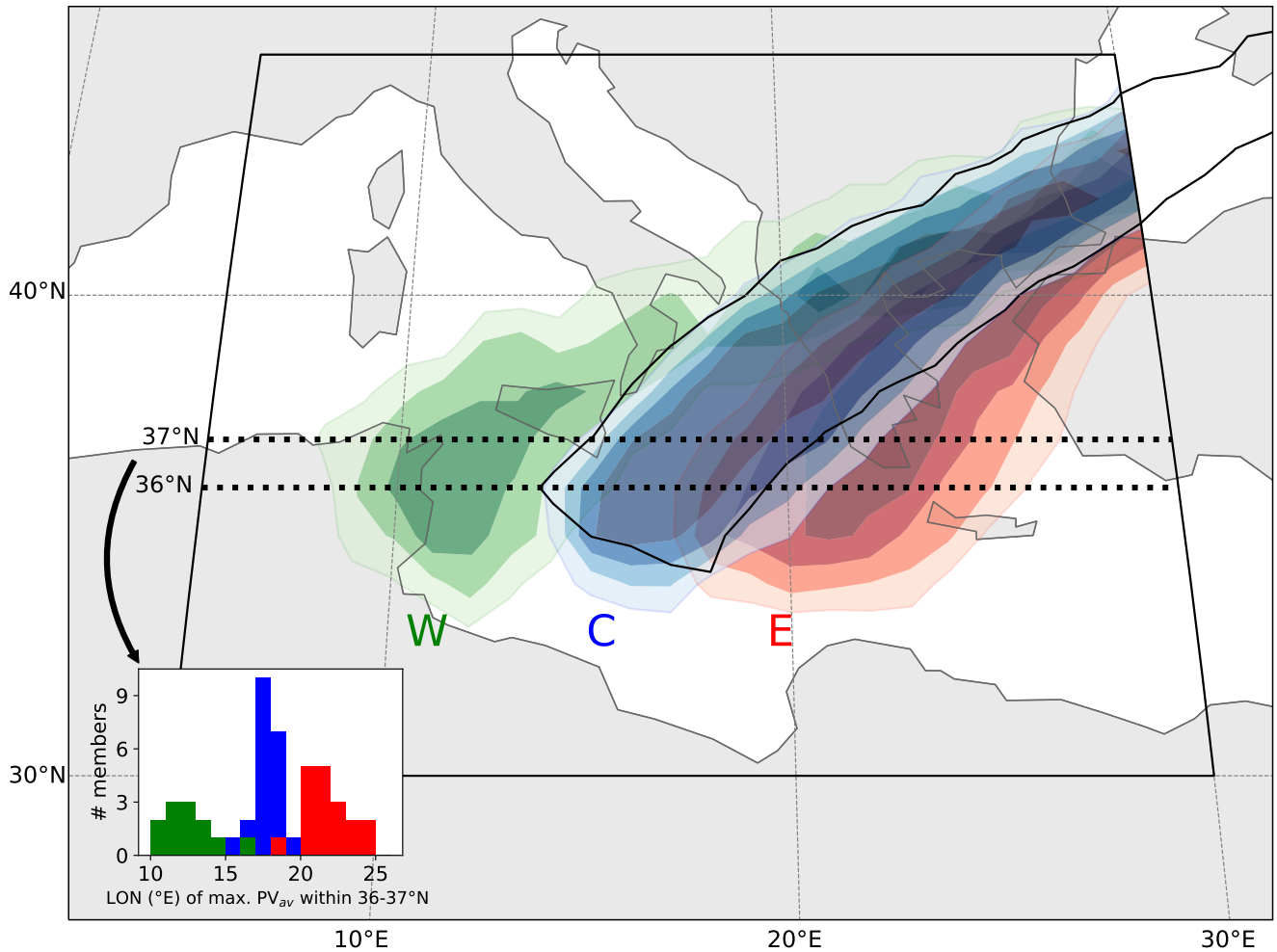
Figure 7. Clustering of ensemble members (initialized at 0000 UTC 24 Sep 2018) according to the position of the upper-level PV streamer in the Mediterranean at 0000 UTC 27 Sep 2018. Colors show frequencies of $PV_{av} \geq 2$ PVU (shading, every 20%) for each cluster (blue: cluster C, green: cluster W, red: cluster E) and the black line the contour where $PV_{av} = 2$ PVU in the operational analysis. The region considered for the clustering is shown by the black box (see text for details).
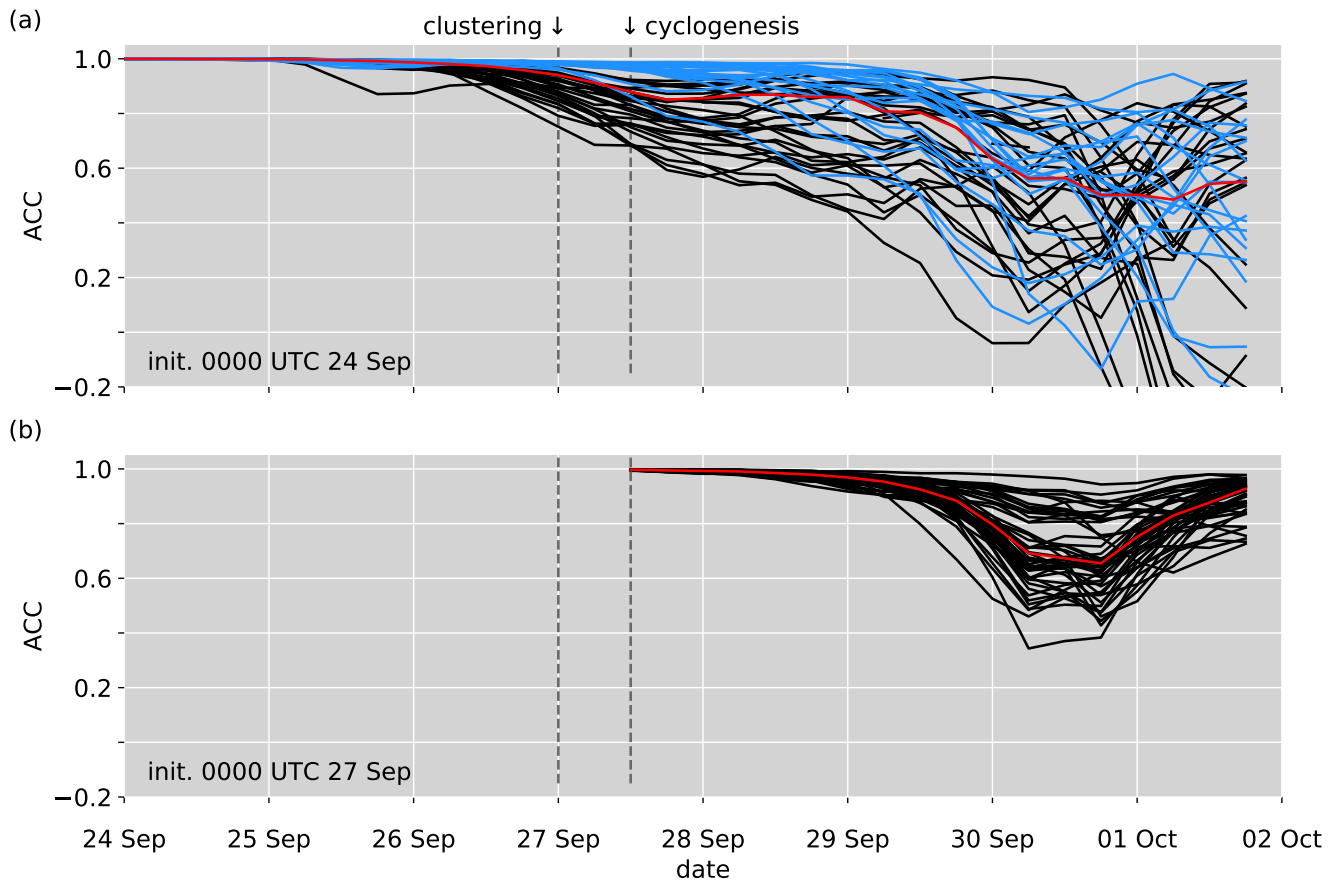
Figure 8. Temporal evolution of the anomaly correlation coefficient of geopotential height at 500 hPa in the Mediterranean box (see Fig. 7) for each ensemble member (black lines) and the median (red line) of the ensemble forecasts initialized at (a) 0000 UTC 24 Sep 2018 (blue lines show members of cluster C) and (b) 0000 UTC 27 Sep 2018.
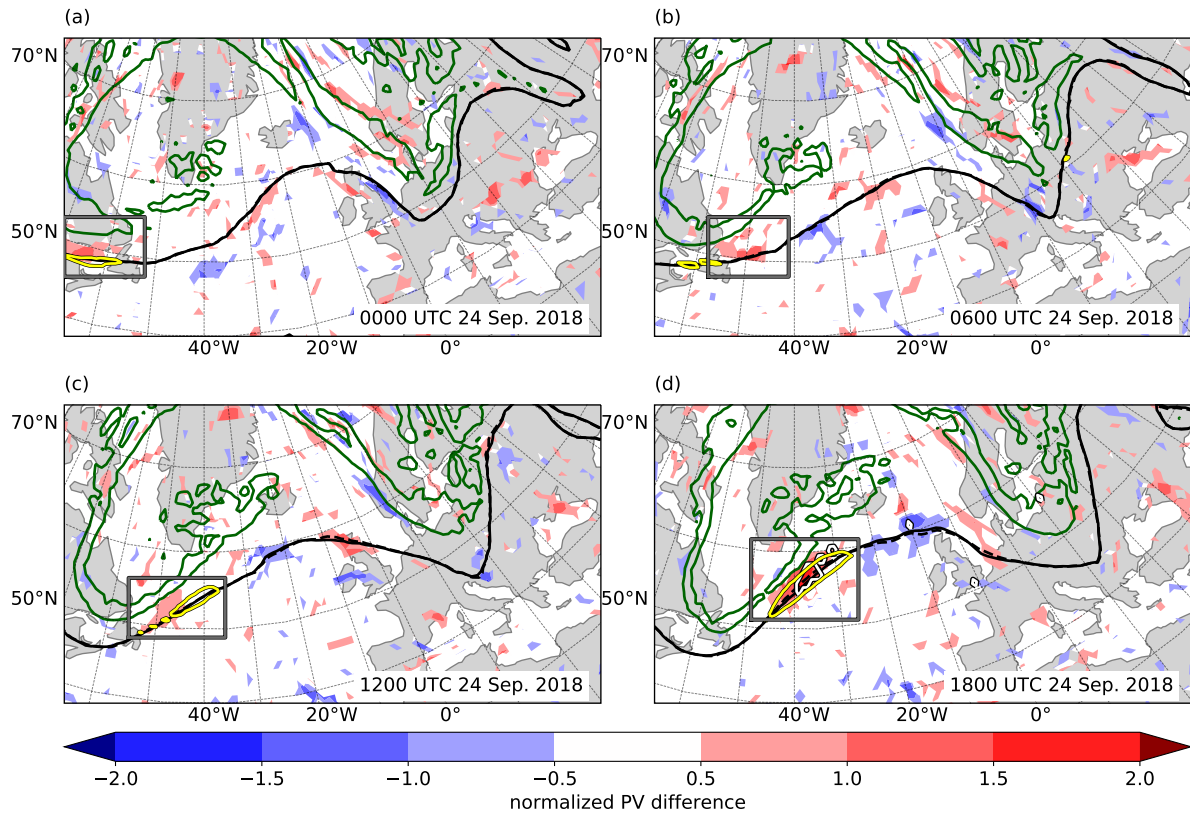
Figure 9. Normalized PV differences between clusters E and W ($\Delta\text{PV}_\text{EW}$, shaded), 2 PVU contour (black lines) of clusters E (solid) and W (dashed), 9 PVU contour (green lines) and wind speed (yellow lines, 70 m s^{-1}) in the operational analysis at 325 K from 0000 UTC 24 Sep to 1800 UTC 24 Sep 2018, every 6 hours. Regions with statistically significant PV differences ($\alpha_\text{fdr} < 0.1$) are marked with white contours and grey boxes mark the regions of interest (for details see text).
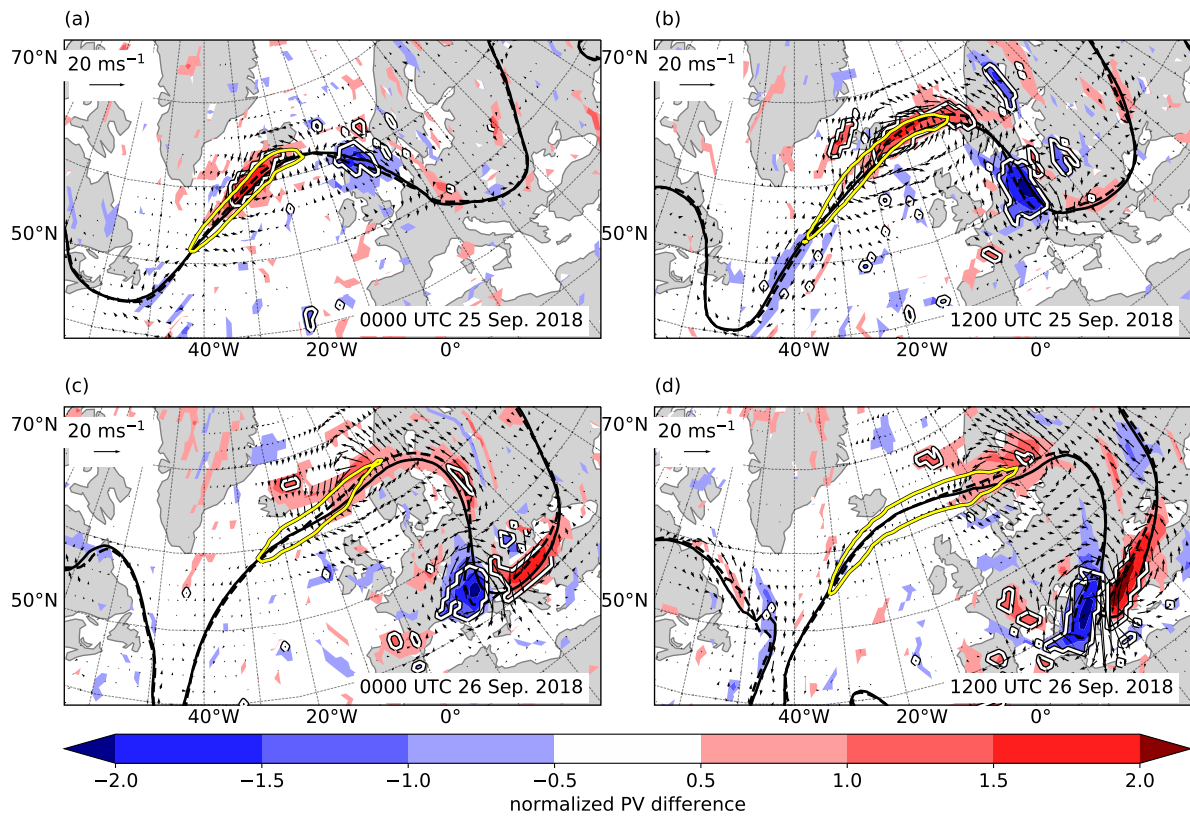
Figure 10. As Fig. 9 but with difference winds (arrows, only if larger than $1\,\mathrm{m\,s^{-1}}$, reference vectors is in top left of panels), without 9 PVU contours and from 0000 UTC 25 Sep to 1200 UTC 26 Sep every 12 hours.
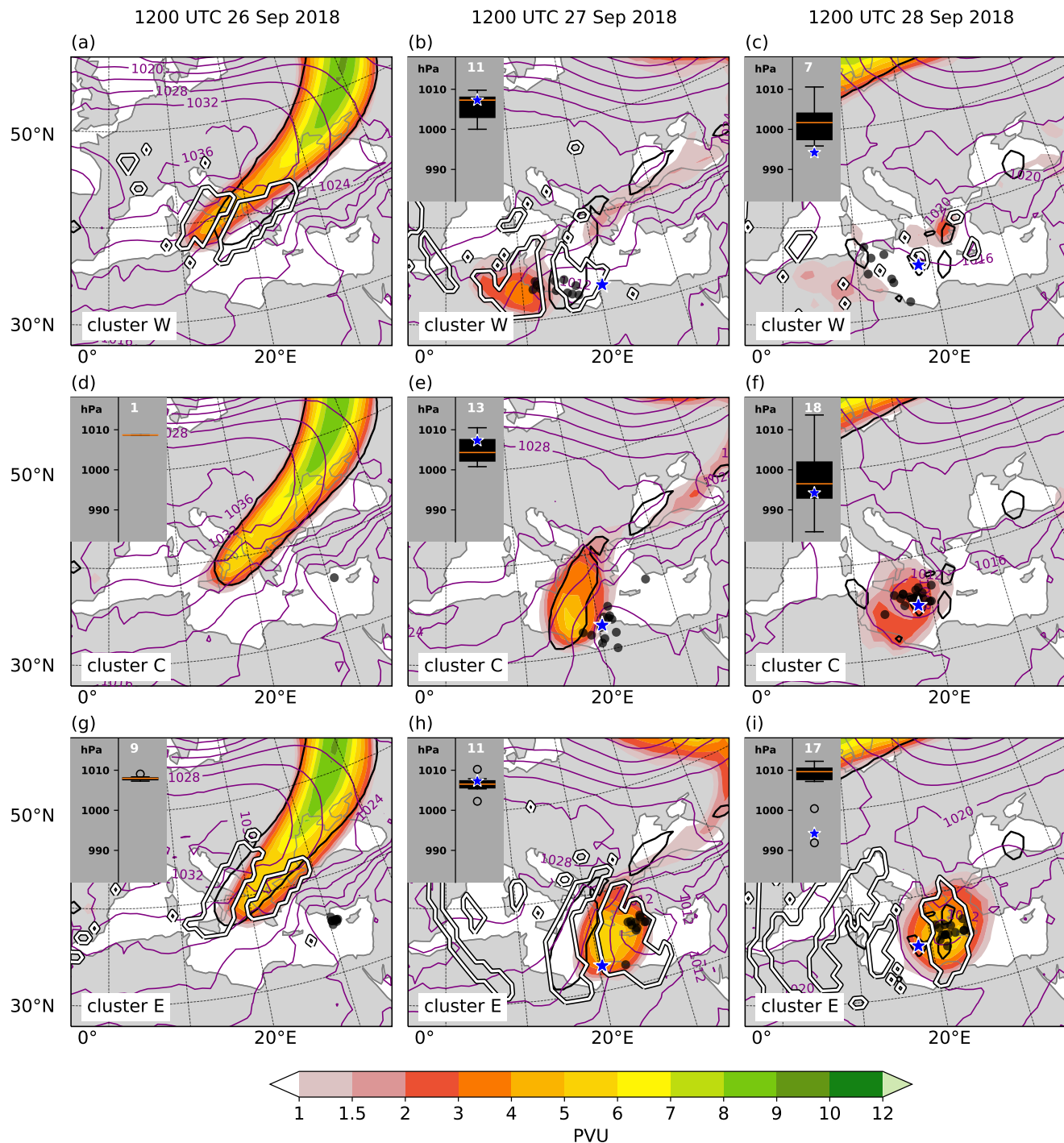
Figure 11. Cluster-mean PV on 325 K (shaded, in PVU), cluster-mean sea level pressure (purple contours, every 4 hPa), analysis 2 PVU contour on 325 K (black), cyclone positions (as identified with the method described in Sect. 2.3) in each ensemble member (black dots), and in the operational analysis (blue star) for clusters W, C, and E (panels a-c, d-f, and g-i) from 1200 UTC 26 Sep to 1200 UTC 28 Sep 2018 every 24 h. Insets at the top left of the panels show box plots of minimum sea level pressure of the cyclones in each cluster and the value in the operational analysis (blue star); white numerals indicate the number of cyclones. Regions where the differences to cluster C of the PV field on 325 K are statistically significant on the α_{fdr}=0.1 level are shown for clusters W and E as white contours (panels a-c, g-i).
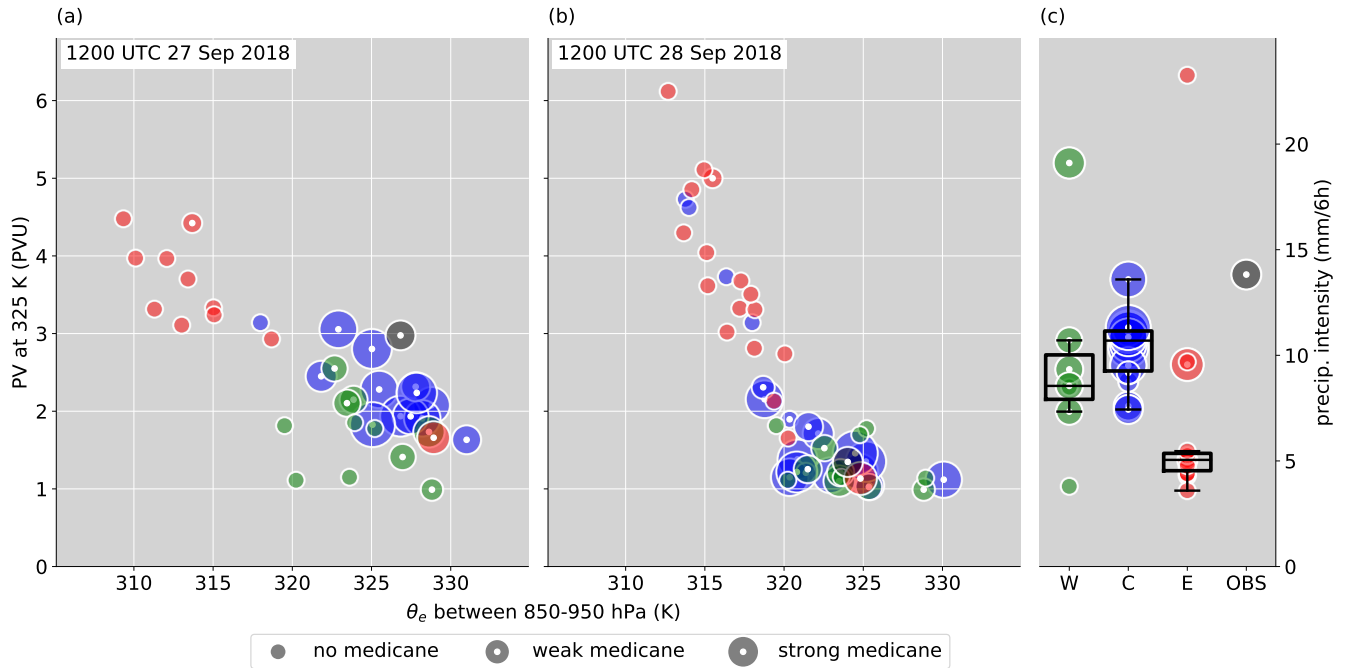
Figure 12. (a,b) Diagram of low-level θ_e vs. upper-level PV for all ensemble members with a cyclone and for the operational analysis at (a) 1200 UTC 27 Sep 2018 and (b) 1200 UTC 28 Sep 2018; markers are as in Fig. 7. (c) Average precipitation intensity within a 250 km radius around the cyclone centre between 1200 UTC 27 Sep and 1200 UTC 28 Sep 2018 for all cyclones present during the full 24 h period as individual data points (colored markers) and standard box plots (black). The observation (OBS) is based on TRMM rainfall data along the track in the operational analysis.