Subseasonal Midlatitude Prediction Skill Following QBO-MJO Activity

Kirsten J. Mayer [1] and Elizabeth A. Barnes [1]

[1]Department of Atmospheric Science, Colorado State University, Fort Collins, CO, USA.

Correspondence: Kirsten J. Mayer (kjmayer@rams.colostate.edu)

Abstract. The Madden-Julian Oscillation (MJO) is known to force extratropical weather days-to-weeks following an MJO event through excitation of stationary Rossby waves, also referred to as tropical-extratropical teleconnections. Prior research has demonstrated that this tropically forced midlatitude response leads to increased prediction skill on subseasonal to seasonal (S2S) timescales. Furthermore, the Quasi-Biennial Oscillation (QBO) has been shown to possibly alter these teleconnections through modulation of the MJO itself and the atmospheric basic state upon which the Rossby waves propagate. This implies that the MJO-QBO relationship may affect midlatitude circulation prediction skill on S2S timescales. In this study, we quantify midlatitude circulation sensitivity and prediction skill following active MJOs and QBOs across the Northern Hemisphere on S2S timescales through an examination of the 500 hPa geopotential height field. First, a comparison of the spatial distribution of Northern Hemisphere sensitivity to the MJO during different QBO phases is performed for ERA-Interim reanalysis and ECMWF and NCEP hindcasts. Secondly, differences in prediction skill in ECMWF and NCEP hindcasts are quantified following MJO-QBO activity. In both hindcast systems, we find that regions across the Pacific, North America and the Atlantic demonstrate an enhanced MJO impact on prediction skill during strong QBO periods on Week 1-4 lead times compared to MJO events during neutral QBO periods.

1 Introduction

Previous research has focused on the impact of the Madden-Julian Oscillation (MJO) on the extratropical circulation in order to extend midlatitude prediction skill (e.g. Henderson et al., 2016; Baggett et al., 2017; Tseng et al., 2018; Zheng et al., 2018). The MJO is a 20-90 day tropical intraseasonal convective oscillation (Madden and Julian, 1971, 1972, 1994), and through its convective heating, initiates an extratropical response through the excitation of quasi-stationary Rossby waves. These waves modulate the mid-latitude circulation days to weeks following MJO activity and have been shown to provide coherent and consistent modulation of midlatitude circulation into subseasonal-to-seasonal (2-5 Weeks; S2S hereafter) timescales (e.g. Hoskins and Karoly 1981; Sardeshmukh and Hoskins 1988; Henderson et al. 2016; Tseng et al. 2018).

More recent research has demonstrated a dependence of the MJO on a stratospheric phenomenon known as the Quasi-biennial Oscillation (QBO). The QBO is an approximately 28 month, downward propagating zonal mean, zonal wind oscillation in the tropical stratosphere and has many subsequent impacts such as modulation of the upper tropical troposphere (e.g. Collimore et al. 2003; Garfinkel and Hartmann 2011b; Son et al. 2017), the subtropical jet (e.g. Simpson et al. 2009; Garfinkel

and Hartmann 2011a) and the stratospheric polar vortex (e.g. Holton and Tan 1980; Garfinkel et al. 2018). The QBO is typically divided into two phases, easterly and westerly (EQBO and WQBO, respectively), determined by the direction of the anomalous zonal wind in the lower tropical stratosphere (Baldwin and Dunkerton, 2001). Recent work has shown that the MJO convective envelope tends to be stronger and have slower eastward propagation and longer path lengths during EQBO compared to WQBO

30 (Son et al., 2017; Nishimoto and Yoden, 2017; Densmore et al., 2019; Zhang and Zhang, 2018). Son et al. (2017) hypothesize that this slower MJO propagation during EQBO is a consequence of strengthened MJO convection, as stronger MJO events tend to propagate more slowly across the Maritime Continent. However, Zhang and Zhang (2018) argue that stronger MJO wintertime events during EQBO are a consequence of a greater number of MJO days instead of larger amplitudes of individual MJO events. While there are still uncertainties regarding the exact impacts of the QBO on the MJO, these studies demonstrate

35 the importance of considering the QBO in MJO research.

Much of the recent MJO-QBO research has focused on the direct impacts of the QBO on the tropical tropopause, and thus, MJO activity, while only a handful of studies have examined how the QBO subsequently impacts MJO teleconnections (e.g. Baggett et al., 2017; Mundhenk et al., 2018; Wang et al., 2018). Baggett et al. (2017) and Mundhenk et al. (2018) emphasize the impact of the QBO on MJO teleconnections through its modulation of MJO-induced Rossby waves, and consequently,

40 changes in the steering and frequency of atmospheric rivers. Wang et al. (2018) found that when accounting for the phase of the QBO, the amplitude of the North Pacific storm track shift in response to MJO activity is greater during EQBO compared to WQBO, which they hypothesize to be from increased MJO strength during EQBO.

An MJO-QBO relationship has also been found in dynamical models. For example, Abhik and Hendon (2019) recently demonstrated that hindcast simulations, initialized with observations during active MJOs, capture the increase in MJO ampli-

45 tude and maintenance during EQBO events after about 5 days. In addition, this strengthened MJO amplitude during EQBO has been shown to translate to increased MJO prediction skill (Marshall et al., 2017; Lim et al., 2019), suggesting that the prediction skill of the subsequent midlatitude teleconnections may also increase following the MJO under EQBO conditions. Baggett et al. (2017) further shows that prediction skill of atmospheric rivers on Week 1-2 timescales varies with QBO phase within ECMWF hindcasts over North America following MJO activity. This highlights the potential for an MJO-QBO relationship to

50 modulate midlatitude prediction skill on S2S timescales.

Since hindcast models capture the increase in MJO amplitude during EQBO as well as exhibit enhanced prediction skill of the MJO in Weeks 1-3 under strong QBOs, this raises the question as to whether the MJO-QBO relationship also translates to enhanced prediction skill of MJO teleconnections under specific QBO phases. This paper explores this question through an analysis of the influence of the QBO on midlatitude prediction skill following active MJOs on S2S timescales within the

55 ECMWF and NCEP hindcasts.

2 Data and Methodology

2.1 Data

We utilize daily mean 500-hPa geopotential height (z500; years 1979-2017) from the European Centre for Medium-Range Weather Forecasts Interim reanalysis (ERA-I; Dee et al. 2011) as well as the ECMWF and NCEP hindcasts obtained from the S2S database established by the World Weather Research Program/World Climate Research Program (WWRP/WCRP; Vitart 2017). The ECMWF hindcasts are composed of 11 ensemble members with hindcasts initialized 4 times a week (years 1995-2016). The NCEP hindcasts are composed of 4 ensemble members with hindcasts initialized daily (years 1999-2010). In the following analysis, the ensemble mean for both models was used, and so, the different number of members between ECMWF and NCEP may contribute to differences in results between models.

We focus on December, January and February (DJF) since MJO teleconnections are strongest during boreal winter (e.g. Madden 1986), and the relationship between the MJO and QBO is strongest during these months as well (e.g. Yoo and Son 2016; Son et al. 2017). The annual cycle is removed from the ERA-I reanalysis by subtracting the daily climatology of z500 across 1979-2017 from the z500 field. For the hindcast models, a daily, lead-dependent climatology is subtracted from each models' z500 field. To do this, we calculate the daily climatology for each lead time independently. Since the ECMWF model is not initialized daily, two (forward and backward moving) 31-day running means are applied to the climatology at all lead times to reduce noise, following Sun et al. (2018). These smoothed lead-dependent daily climatologies are then subtracted from the z500 field of the corresponding model to remove the annual cycle.

There is presently no definitive understanding of the impact of the El Nino Southern Oscillation (ENSO) on the QBO-MJO relationship. Some earlier research indicates that ENSO has a limited impact on the QBO-MJO interaction (e.g. Yoo and Son 2016; Nishimoto and Yoden 2017); however, recent work on QBO-MJO teleconnections has shown a possible dependency of results on ENSO (Son et al., 2017; Wang et al., 2018; Sun et al., 2019). In addition, other research suggests that the QBO affects ENSO teleconnections (Garfinkel and Hartmann, 2010; Richter et al., 2015; Hansen et al., 2016), which may consequently impact the MJO and its teleconnections. Thus, in an attempt to ensure our results are not somehow biased by ENSO, we use the Nino3.4 Index (climatedataguide.ucar.edu/climate-data) to remove strong ENSO winter seasons from our analysis. Specifically, when the amplitude of the NINO3.4 index for a month within DJF is greater than 1°C (signifying El Nino) or less than -1°C (signifying La Nina), that DJF season is excluded from the analysis. With that said, we have repeated our analysis with ENSO seasons included and find that our STRIPES conclusions remain the same. The prediction skill conclusions also remain the same during EQBO, but the WQBO results appear more sensitive to ENSO (see Supplemental Figures S6-S9). The impact of ENSO on the MJO-QBO relationship and their teleconnections still remains an active area of research.

2.2 MJO and QBO Indices

The real-time multivariate MJO (RMM) index is used to define the amplitude and phase of the MJO in the ERA-I reanalysis (Wheeler and Hendon, 2004). This index uses empirical orthogonal function (EOF) analysis applied to anomalous outgoing longwave radiation (OLR) and 200- and 850-hPa zonal wind, near-equatorially averaged (15°S to 15°N), to determine the first

two principal components (RMM1 and RMM2). A day is considered to have an active MJO when the RMM amplitude for that day (defined as $\sqrt{(RMM1^2 + RMM2^2)}$) is greater than 1.0. The MJO phase is then defined as $tan^{-1}(RMM2/RMM1)$ and largely corresponds to the longitudinal location of the convective envelope. Active MJO dates within ERA-I that correspond to initialization dates in ECMWF and NCEP are determined from this index. The RMM index is not separately calculated for each hindcast model because we do not aim to quantify the ability of the models to forecast the MJO directly (e.g. Vitart 2017). Rather, we use the index calculated from reanalysis to see how the hindcast models initialized on observed active MJO days ultimately forecast MJO teleconnections.

Identical to the definition of Yoo and Son (2016), the QBO index is calculated within ERA-I using monthly standardized zonal wind at 50-hPa, area-averaged between 10°S to 10°N. Westerly QBO (WQBO) and Easterly QBO (EQBO) events are defined as when the standardized value is greater than 0.5σ or less than -0.5σ, respectively. Absolute values less than 0.5σ are considered neutral QBO (NQBO) events.

2.3 Methods

Quantification of each models' ability to represent MJO teleconnections under different QBO phases is conducted using the Sensitivity to the Remote Influence of Periodic Events (STRIPES) index (Jenney et al., 2019). STRIPES is an index recently developed to determine regions of extratropical sensitivity to remote periodic events such as the MJO. As used here, the STRIPES index quantifies the strength and consistency of MJO teleconnections in z500 through average phase and 0-28 day lead information at individual grid points for a variety of observed phase speeds (5-8 days/phase; Wheeler and Hendon 2004). Specifically, a composite of average z500 anomalies for each MJO phase and lead (phase-lead diagram) is created for each grid point in the Northern Hemisphere (example shown in Figure 1a). For further intuition of the phase-lead diagram, Figure 1b and 1c show composite z500 anomalies for the domain around 45°N and 5°W (marked by the white X) 12 days following phase 6 and phase 2, respectively. The value of the box in the phase-lead diagram is the same as the value plotted at the X in Figure 1b,c. In a phase-lead diagram, MJO induced quasi-stationary rossby waves are apparent as slowly alternating-sign z500 anomalies with lead following a specific phase of the MJO (e.g. Figure 1a). In addition, the MJO is a propagating phenomenon with a phase speed of approximately 5-8 days/phase. Therefore, if there is a teleconnection signal 10 days following phase 2, this signal is likely also present 5 days following phase 3 in the same region, in a composite sense. On a phase-lead diagram, this is seen as a diagonal line or 'stripe' slanted at the phase speed of the MJO (Figure 1a). Therefore, if a region is sensitive to the MJO, we expect alternating z500 anomaly stripes approximately sloped at the average phase speed of the MJO, as in Figure 1a, which we refer to as the 'stripey-ness'.

To calculate STRIPES, averages along the slopes in the phase-lead diagram corresponding to the MJO phase speed are calculated, and if there are alternating stripes (i.e. sensitivity to the MJO), the resulting averages concatenated together will oscillate between positive and negative z500 anomalies as a sine wave, for which the amplitude can be calculated. The amplitude of this oscillatory vector is the STRIPES index (Jenney et al. 2019). The more sensitive the region is to MJO teleconnections, the larger the STRIPES index. Therefore, the STRIPES index allows us to regionally quantify the strength, consistency and

propagation of the MJO impact on the extratropics and thus, allows us to quantify the ability of hindcast models to capture tropical-extratropical teleconnections on one to four week timescales in a single metric.
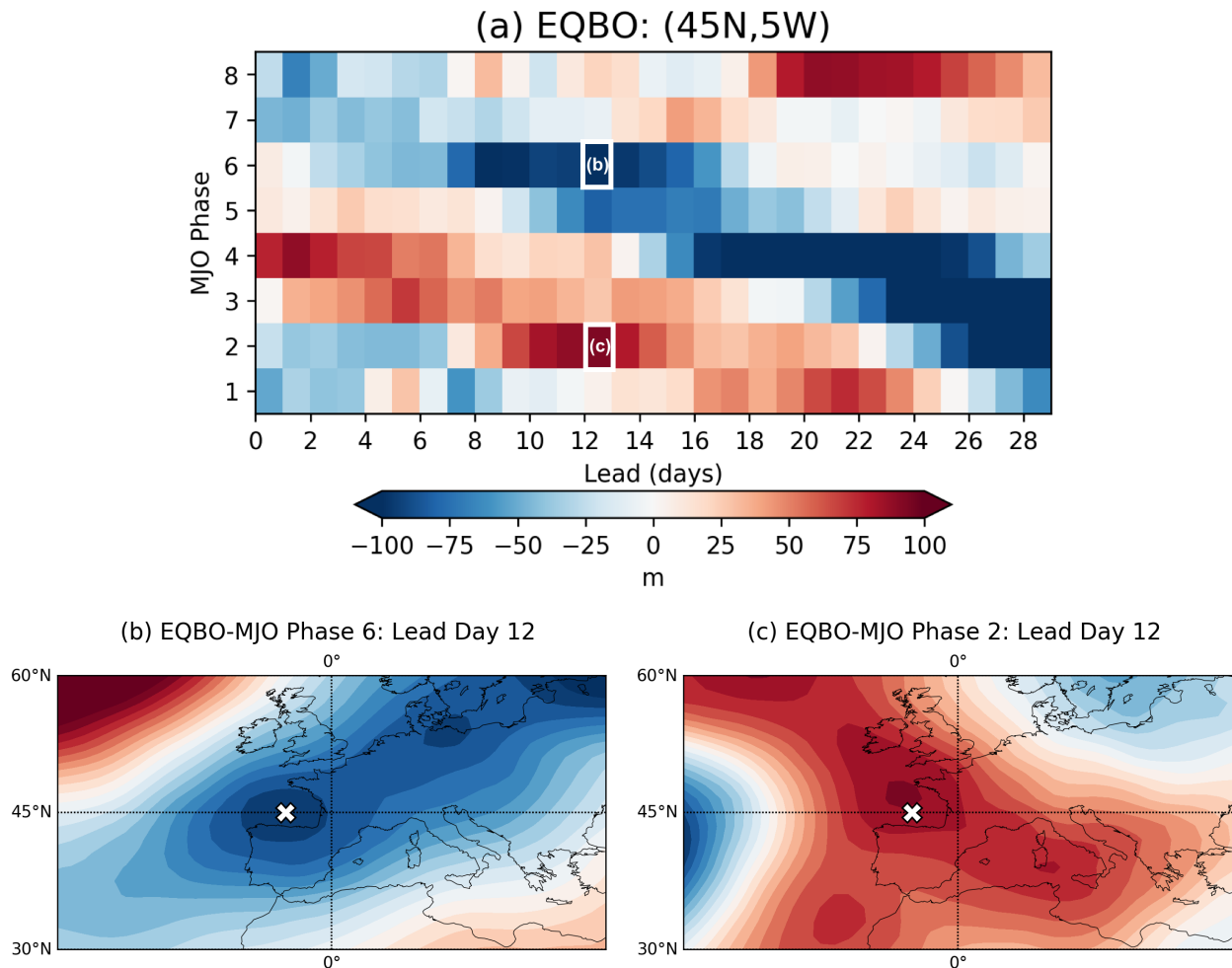
Figure 1. Boreal winter (DJF) composite ERA-I z500 anomalies subsampled to ECMWF initialization dates (1995-2016) for (a) each MJO phase during EQBO vs lead at 45N and 5W. White boxes and text denote the corresponding panels below. The bottom panels include composite ERA-I z500 anomalies subsampled to ECMWF initialization dates (DJF, 1995-2016) over Europe for (b) Phase 6 and (c) Phase 2 at lead day 12. The white X denotes 45°N and 5°W.

For equal comparison of STRIPES between the models and reanalysis, we calculate STRIPES for ERA-I only with dates that overlap with the hindcasts, thus, the ERA-I STRIPES figures differ for ECMWF versus NCEP dates. It should be noted that the westerly phase of the QBO has been documented to reduce the propagation speed of the MJO (Nishimoto and Yoden 2017), however, we find that our STRIPES results are robust to changes in phase speed of +/- 2 days/phase. Also note that

since our application focuses on extratropical sensitivity in z500, we use z500 anomalies in terms of meters instead of standard deviation for STRIPES, different from Jenney et al. (2019). Standardization may mute the extratropical signal due to the greater variability of z500 in the midlatitudes, which is of main interest here. In addition, we wish to retain any differences in z500 anomaly amplitudes between the QBO phases.

STRIPES values that are statistically larger than expected by chance are determined using a bootstrapping method. The number of random days grabbed corresponds to the observed number of days for the QBO-MJO event of interest. In order to retain autocorrelation within MJO events, we keep the day-of-year (DOY) and phase distribution information for each MJO event and randomly sample years (with replacement). Since the ECMWF hindcast data is not initialized on the same day each year, if the DOY needed is not available for a particular year, we instead use the date of initialization closest to this DOY. From this sample, we calculate STRIPES. This is repeated 250 times for each latitude and longitude. Any STRIPES value greater than the 90th percentile of these bootstrapped values are deemed significant. When the data is subdivided by QBO phase, we begin to see the effects of sample size on the uncertainty, leading to fewer points of significance. However, when the QBO phase is not considered or in other words, when all MJO events are included (see Figure S1), the statistical analysis shows significance in regions of large STRIPES values. This bootstrapping analysis is only conducted on ERA-I, as these are the 'observed' sensitivities and thus, the regions of interest.

To quantify midlatitude prediction skill, a daily area-weighted Pearson correlation is conducted between hindcast and ERA-I anomalous z500 (anomaly correlation coefficient; ACC). The data is separated into NQBO-, EQBO- and WQBO-MJO events in each hindcast dataset and the corresponding reanalysis data is obtained from ERA-I. The ACC between a given model day and the same day in ERA-I is calculated within a centered 60° longitude wide box extending from 30-60° N. Our conclusions are not affected by the latitudinal extent of the box when it is varied by +/- 10-30 ° N. This calculation is repeated for every initialization and subsequent lead time as well as every 5° longitude beginning at 0°E. ACCs are grouped and averaged by QBO phase to obtain average ACCs across the Northern Hemisphere at every lead for each QBO phase (see Supplemental Figure S3 for an example).

Statistically significant differences in ACCs across lead and longitude are also computed with a bootstrapping method. Specifically, all model data within DJF is shuffled and random dates are grabbed. The number of random dates corresponds to the number of observed dates for the particular QBO phase and MJO activity being tested. These dates are then found in ERA-I. The spatial correlations between the model and the observations are calculated and then averaged to get an average ACC. This is repeated for each QBO-MJO combination, and the differences between their ACCs is calculated. The above analysis is repeated 10,000 times for each longitude and lead time. Differences greater than the 90th percentile of the 10,000 bootstrapped differences are considered significantly greater from that expected by chance. In this bootstrapping analysis, we were able to repeat the calculations 10,000 times (instead of 250) because the calculation was less computationally expensive.

3 Results

 ### 3.1 Extratropical Sensitivity

The left column of Figure 2 shows the STRIPES analysis of ERA-I for days within the ECMWF hindcasts, split by QBO phase. Darker shading indicates regions of greater sensitivity to the MJO for each QBO state. Regions along the North Pacific and Atlantic storm tracks as well as over North America are highlighted by STRIPES following the MJO for all phases of the QBO (Figure 2a,c,e). This is consistent with previous research as these regions have been shown to be sensitive to MJO excited Rossby waves through, for example, their modulation of the North Atlantic Oscillation (Cassou, 2008), the Pacific North American Oscillation (Mori and Watanabe, 2008) and Northern Hemisphere wintertime blocking (Henderson et al., 2016). Interestingly, the Pacific and Atlantic sectors have similar STRIPES values. One may expect higher STRIPES values over the Pacific compared to the Atlantic since the Pacific is generally known to have a strong response to the MJO. We hypothesize that the Atlantic and European sectors also have similar STRIPES values to that of the Pacific due to enhanced blocking over the Atlantic and Europe at later leads following the MJO (Henderson et al. 2016). Since the STRIPES index accounts for all leads as well as the strength and consistency of the z500 anomalies, we therefore may expect STRIPES values over the Atlantic and European sectors to be large as well.

The right column of Figure 2 shows the STRIPES analysis of the ECMWF hindcasts for the same dates. ECMWF largely captures the spatial patterns and locations sensitive to the MJO under different QBO phases (spatial correlation with ERA-I: $r_{NQBO-MJO} = 0.92$, $r_{EQBO-MJO} = 0.93$, and $r_{WQBO-MJO} = 0.95$), but overall the model has smaller STRIPES values than ERA-I. This is likely a result of model forecast degradation at later lead times since the calculation of STRIPES utilizes z500 forecasts out to 28 days lead time.

An examination of the NCEP hindcasts shows that it also generally captures regions sensitive to the MJO under varying phases of the QBO (Figure 3b,d,f; spatial correlation with ERA-I: $r_{NQBO-MJO} = 0.96$, $r_{EQBO-MJO} = 0.95$, and $r_{WQBO-MJO} = 0.93$) and is also weaker than the corresponding ERA-I analysis (Figure 3a,c,e). The ERA-I STRIPES analysis for NCEP hindcasts largely has the same features as the ERA-I analysis for ECMWF hindcasts, but with larger values due to differences in sample size and dates of initialization between NCEP and ECMWF. From this STRIPES comparison (Figures 2 and 3), we conclude that the ECMWF and NCEP hindcast models generally capture Northern Hemisphere regions sensitive to the MJO as highlighted by large spatial correlations between each model and ERA-I.
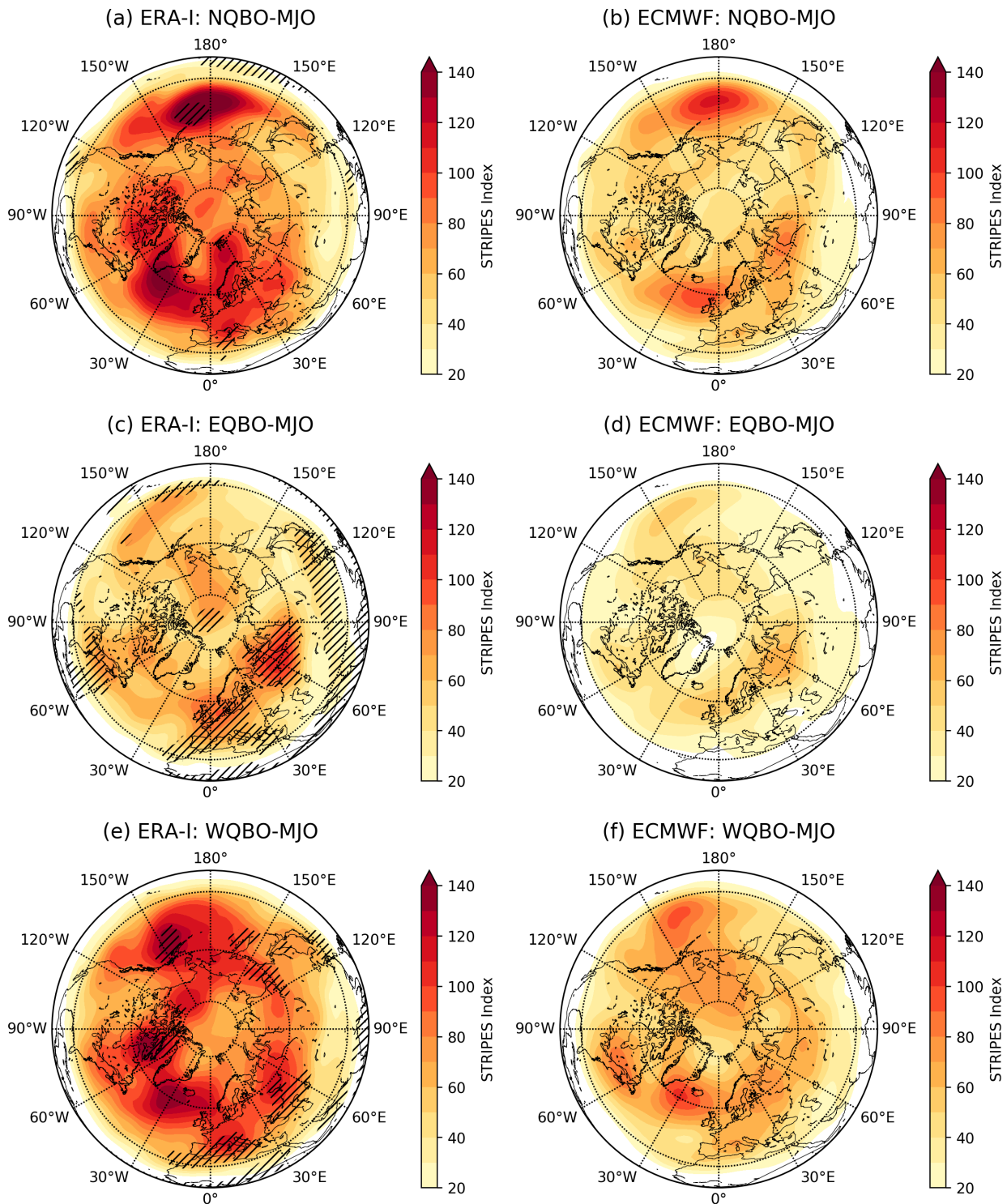
Figure 2. STRIPES values for (left) ERA-Interim and (right) ECMWF for all (top) NQBO-MJO, (middle) EQBO-MJO and (bottom) WQBO-MJO events. (a,c,e) Black hatches denote STRIPES values that are statistically larger than expected by chance at 90% confidence in ERA-I. Significance is only calculated for ERA-I.
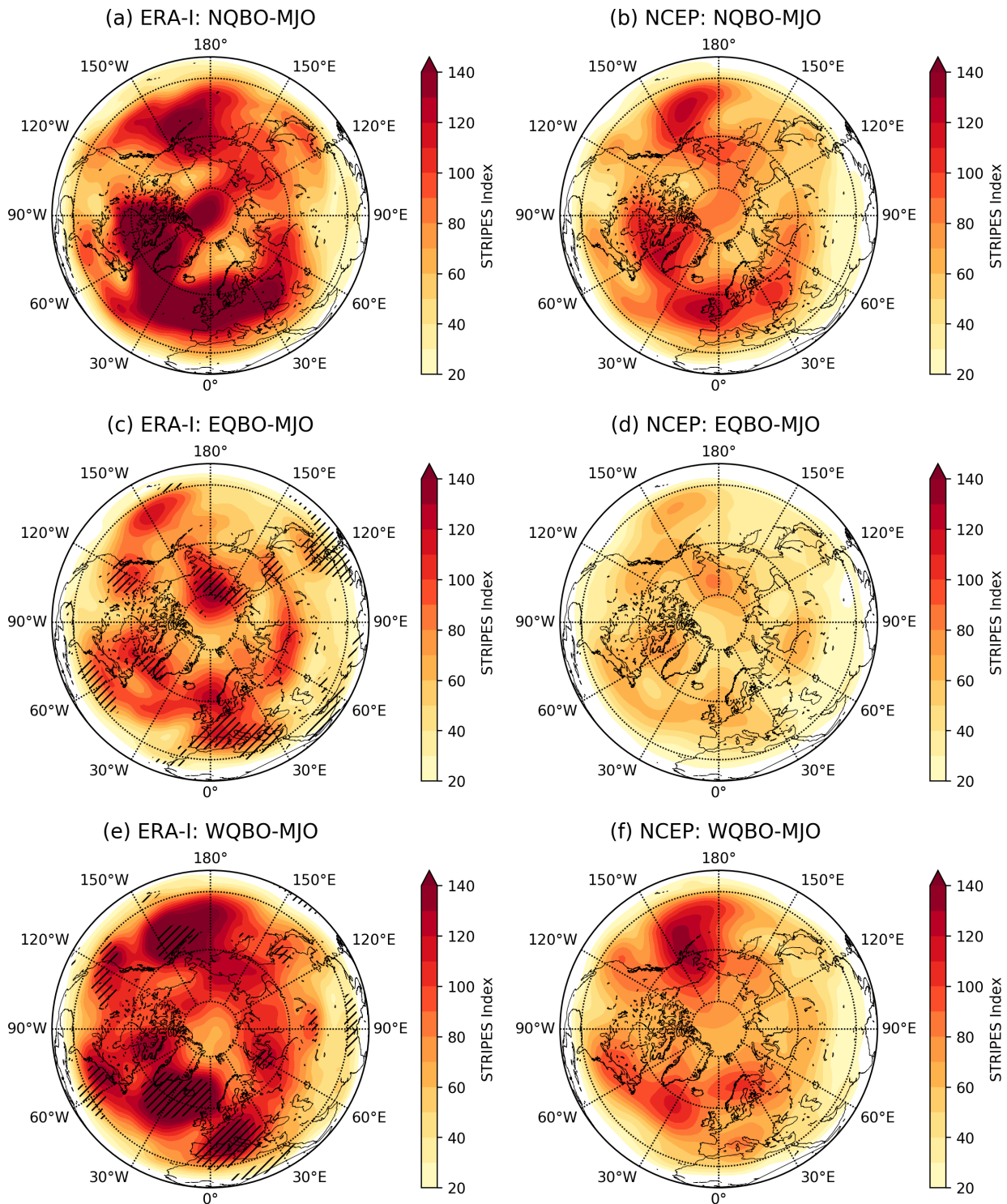
Figure 3. STRIPES values for (left) ERA-Interim and (right) NCEP for all (top) NQBO-MJO, (middle) EQBO-MJO and (bottom) WQBO-MJO events. (a,c,e) Black hatches denote STRIPES values that are statistically larger than expected by chance at 90% confidence in ERA-I. Significance is only calculated for ERA-I.

Recent research has shown that during EQBO, the MJO amplitude is larger and the convective envelope propagates slower compared to MJO activity during WQBO (Son et al., 2017; Nishimoto and Yoden, 2017; Zhang and Zhang, 2018). If direct impacts to the MJO (e.g. through changes in upper tropospheric tropical static stability) lead to changes in MJO teleconnection sensitivity across the Northern Hemisphere, we might expect EQBO-MJO events to have larger midlatitude sensitivity to the MJO compared to WQBO-MJO. Instead, we find that Northern Hemisphere sensitivity to the MJO is significantly reduced during EQBO-MJO events compared to WQBO-MJO events (compare Figure 2c,e and Figure 3c,e; significance of difference not shown). We explored this further and found that this difference can largely be explained by the tendency for WQBO to have larger magnitude z500 anomalies compared to EQBO, not more distinct stripes. This is likely due to the larger sample size during EQBO (Table S1) leading to reduced noise in the average. Therefore, when the amplitude differences between the z500 anomalies are accounted for through normalization, the difference in Northern Hemispheric sensitivity to the MJO between QBO phases is greatly reduced (Figure S3).

3.2 Prediction Skill

3.2.1 Regional Prediction Skill

Knowing that the ECMWF and NCEP hindcasts generally capture regional sensitivity to the MJO, we next address whether strong QBOs enhance MJO impacts on midlatitude skill. As mentioned in the introduction, EQBO has been found to impact the MJO in ways that may enhance MJO teleconnections (e.g. Son et al. 2017; Nishimoto and Yoden 2017). Since enhanced activity may provide a prominent signal above model noise and uncertainty, and thus, hypothetically lead to enhanced prediction skill, we focus here on only improved prediction skill (see Supplemental Figure S4 for regions of decreased prediction skill). Note that prediction skill at one week lead times is not likely to be significantly different following active MJOs compared to inactive MJOs since forecast models already have relatively good prediction skill for these early leads. Where we would expect the MJO to provide additional prediction skill is on timescales longer than one week. Here, skill is calculated as an anomaly spatial correlation coefficient (ACC) between z500 from the hindcasts and ERA-I (see Section 2.3), and we compare this skill over QBO-MJO combinations to skill during inactive MJOs. Figure 4 shows z500 ACC as a function of lead time for the North Pacific (165°W, 30-60°N), North Atlantic (30°W, 30-60°N), and Europe regions (0°E, 30-60°N).

We invoke two requirements to address the question of whether a particular strong QBO (EQBO or WQBO) enhances the MJO impact on midlatitude prediction skill compared to an inactive QBO (NQBO), and a third requirement to answer whether the MJO leads to enhanced midlatitude prediction skill under a strong QBO compared to an NQBO. Each requirement builds on the previous requirement. For example, we can only examine requirement two if requirement one has been passed. These requirements are summarized below:

1. A significant MJO impact

2. A significant MJO impact during a strong QBO that is *significantly greater than* an MJO impact during NQBO

3. Enhanced prediction skill following an MJO during a strong QBO that is *significantly greater than* that during NQBO
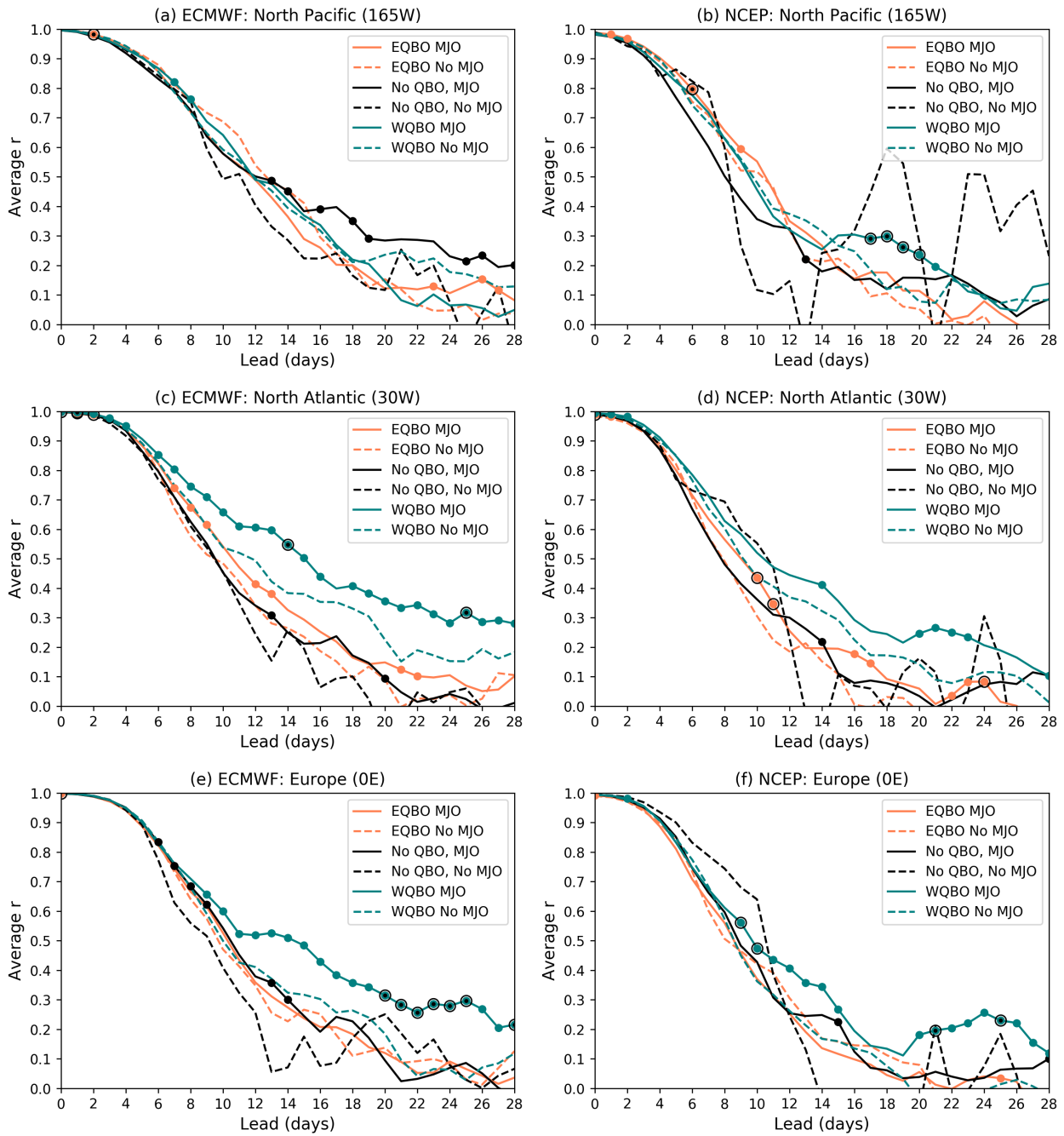
Figure 4. Anomalous spatial correlation coefficient at (top) 165°W, (middle) 30°W and (bottom) 0°E for (left) ECMWF and (right) NCEP. Solid lines correspond to active MJOs while dashed lines correspond to inactive MJOs. Colors refer to the phase of the QBO. Colored dots denote regions/leads where requirement 1 is passed at 90% confidence for the corresponding QBO. Black circles indicate regions/leads where requirement 2 is passed at 90% confidence and small black dots on orange/teal lines indicate regions/leads where requirement 3 is passed at 90% confidence. See text for details.

The first requirement is the presence of an MJO impact on midlatitude prediction skill during specific phases of the QBO. An 'MJO impact' on midlatitude prediction skill is defined as a significant difference in midlatitude ACC between active MJO and inactive MJO events and is denoted by colored dots in Figure 4. In other words, where the solid line (EQBO-/WQBO-/NQBO-MJO) is significantly above the corresponding colored dashed line (EQBO-/WQBO-/NQBO-noMJO).

Where there is an MJO impact we continue to the second requirement. The second requirement is that the magnitude of the significant MJO impact under strong QBOs is larger than the significant MJO impact under NQBOs. This second requirement is denoted by black circles around the colored dots in Figure 4. These two requirements together ensure that (1) there is an MJO impact and (2) that this impact is enhanced during strong QBOs compared to neutral QBOs.

Requirement three specifies significantly enhanced prediction skill following an MJO during strong QBOs compared to NQBO. In Figure 4, this is when a colored line (EQBO-/WQBO-MJO) is significantly above the black line (NQBO-MJO) and is denoted as a small black dot on a teal/orange dot. We applied this requirement to ensure that regions with enhanced MJO impacts during strong QBOs also have overall greater prediction skill following active MJO events compared to NQBO-MJO events.

For requirement one, we see that there is an MJO impact in the North Atlantic and Europe during WQBO out to Week 4 in ECMWF (Figure 4c,e; teal dots). In NCEP during WQBO, we see an MJO impact on Weeks 2 and 4 over Europe and Weeks 3-4 in the North Pacific and North Atlantic (Figure 4b,d,f; teal dots). For requirement 2, there is an enhanced MJO impact during WQBO in ECMWF over Europe during Weeks 3-4 and in NCEP over the North Pacific in Week 3. For all of the regions requirement 2 is passed, requirement 3 is also satisfied. Therefore, where WQBO enhances the MJO impact on midlatitude prediction skill, WQBO also leads to increased prediction skill following MJO events compared to NQBO.

While Figure 4 shows results for three specific regions, we extend these results to all longitudes in Figure 5. Specifically, the four panels show the difference in ACC between EQBO-MJO and EQBO-noMJO (Figure 5a,b; orange solid and dashed lines in Figure 4) and WQBO-MJO and WQBO-noMJO (Figure 5c,d; teal solid and dashed lines in Figure 4). The left column of Figure 5 shows the differences within ECMWF and the right column shows differences within NCEP. Shading specifies increased prediction skill following the MJO compared to inactive MJOs during the specific phase of the QBO and grey dots denote a significant MJO impact (Requirement 1), as denoted in Figure 4 by the colored dots. Regions where the MJO impact is significantly enhanced during a strong QBO compared to NQBO is denoted with a black circle around the grey dots (Requirement 2), and when the MJO leads to enhanced prediction skill during strong QBOs compared to NQBO (Requirement 3), a small black dot is plotted, as in Figure 4.

Focusing on the first requirement (grey dots), during EQBO (Figure 5a,b), there is an MJO impact on midlatitude prediction skill in North America at Week 2 leads for NCEP and ECMWF and extending into Asia at Week 2-3 leads for ECMWF and at Week 4 leads for NCEP. During WQBO in ECMWF and NCEP (Figure 5c,d), there is an MJO impact in the East Pacific into North America through Week 1. This impact continues through Week 2 into the North Atlantic and Europe and continues over Europe for Weeks 3-4 (Figure 5c,d). In NCEP, the MJO impact also occurs in the Pacific during Week 3 (Figure 5d). From Figure 5, we see that in both models, there is an MJO impact on midlatitude prediction skill during EQBO from North America to East Asia and during WQBO from the North Pacific through Europe on subseasonal timescales.
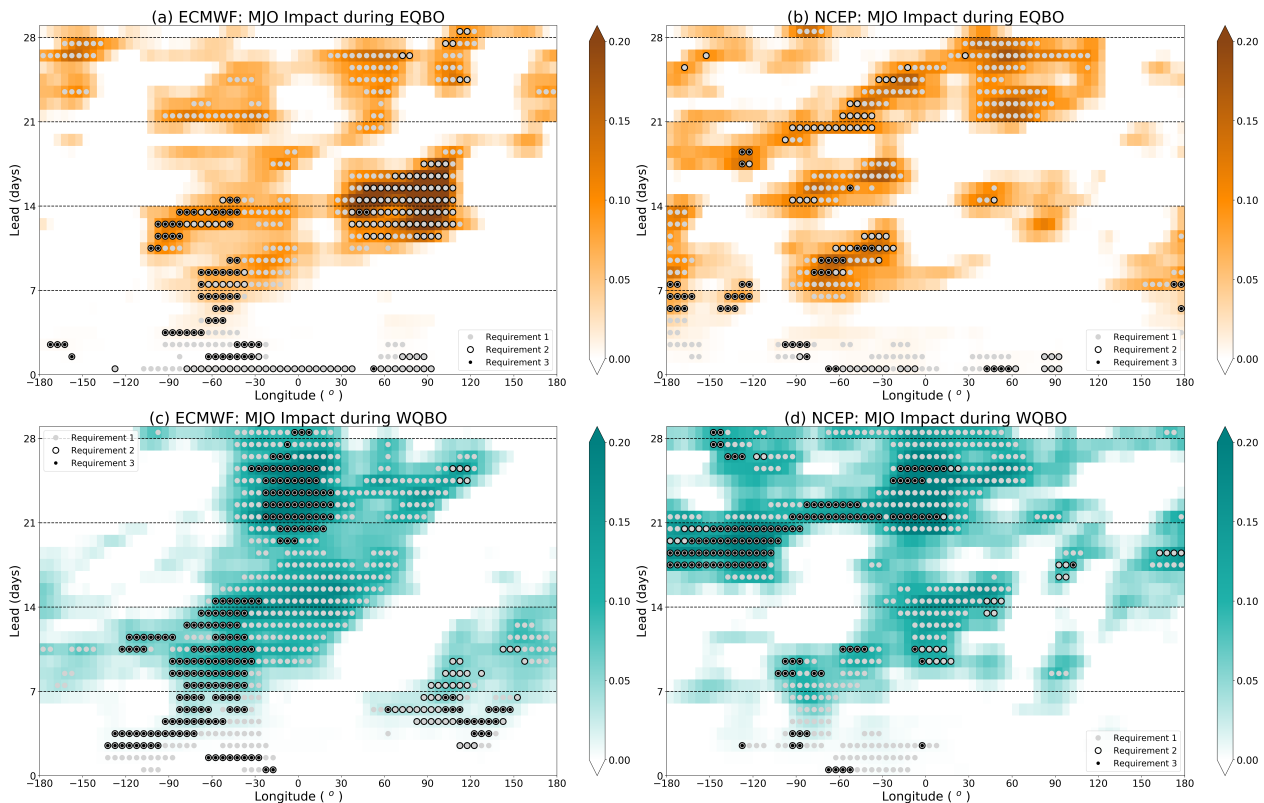
Figure 5. Anomalous correlation coefficient between (top) EQBO-MJO and EQBO-noMJO and (bottom) WQBO-MJO and WQBO-noMJO for (left) ECMWF and (right) NCEP at each longitude and lead from model initialization. Correlations are calculated within a 60°wide box, centered on each longitude, extending from 30-60°N. Shading denotes the phase of the QBO. Grey dots denote regions/leads where requirement 1 is passed at 90% confidence. Black circles indicate regions/leads where requirement 2 is passed at 90% confidence and small black dots indicate regions/leads where requirement 3 is passed at 90% confidence. See text for details.

For the second requirement (black circles), we see that during EQBO in ECMWF (Figure 5a) the MJO impact is greater than during NQBO over North America and the North Atlantic on Week 1-2 and over Asia on Week 2-3. For NCEP (Figure 5b), this occurs over the North Pacific to the Atlantic on Week 1-2 timescales, and again over the Atlantic on Week 3-4. During

255   WQBO in ECMWF (Figure 5c), there is an enhanced MJO impact in the East Pacific into the North Atlantic through Week 2. This enhanced MJO impact reemerges over the North Atlantic by Week 4. In NCEP (Figure 5d), there is an enhanced MJO impact over the Pacific on Week 3 and over North America and the Atlantic on Week 4. Thus, Figure 5 suggests that strong QBOs enhance the MJO impact on midlatitude subseasonal prediction skill from the Pacific to Europe, although we remind the reader once again of the small NQBO sample sizes.

260   For the third requirement (small black dots), during EQBO this requirement is satisfied over North America and the North Atlantic on Week 1-2 timescales in ECMWF and NCEP. For WQBO in ECMWF, this requirement is satisfied over the East Pacific through the Atlantic on Week 1-2 leads and reemerges over the North Atlantic and Europe during Week 4. For NCEP,

during WQBO the third requirement is satisfied on Week 3-4 timescales over the North Pacific to the Atlantic. Interestingly, for WQBO in both models, the third requirement is almost always satisfied over the regions/leads where requirement two is satisfied. This suggests that WQBOs enhance the MJO impact on midlatitude prediction skill as well as enhance overall prediction skill compared to NQBOs following active MJOs from the Pacific to Europe on Week 1-4 timescales in ECMWF and on Week 3-4 timescales in NCEP.

Since EQBO is thought to increase the amplitude of the MJO as well as help to propagate the MJO further into the Pacific Ocean compared to WQBO (Son et al., 2017; Nishimoto and Yoden, 2017; Zhang and Zhang, 2018), one may have expected that active MJOs during EQBO conditions would lead to stronger MJO teleconnections and thus, act to enhance subseasonal prediction skill in the midlatitudes. However, from Figure 5 we see that while EQBO has an enhanced MJO impact on Week 1-4 leads from North America to East Asia, WQBO has an enhanced MJO impact as well as enhanced overall prediction skill compared to NQBO on subseasonal timescales, specifically from the North Pacific through Europe. Perhaps most striking is the enhanced MJO impact and increased prediction skill following active MJO events during WQBO over the North Atlantic in Weeks 3-4 in both ECMWF and NCEP. This result is supported by recent research showing that the North Atlantic Oscillation and MJO connection is stronger during WQBO (Feng and Lin 2019; Song and Wu 2020). Furthermore, enhanced prediction skill of Atmospheric Rivers over Alaska is also found following active MJOs during WQBO (Baggett et al., 2017). Previous research has shown that WQBO alone leads to enhanced midlatitude prediction skill over the North Atlantic (Boer and Hamilton, 2008); however, we specifically look at the difference between active and inactive MJOs under WQBO and therefore, have removed any possible WQBO background skill. We also extended this analysis beyond Week 4 (not shown). We find that while there is still an MJO impact through 40 days, this MJO impact is seldom significantly larger than the MJO impact during NQBOs. This suggests that strong QBOs do not lead to an enhanced MJO impact beyond 4 weeks in these models.

It should be noted that inactive MJOs during NQBO events with ENSO removed only occur 12 times in ECMWF and 3 times in NCEP. When this is the case, there is shading across all longitudes (Figure S5). If ENSO events are not removed, the sample sizes increase to 47 and 52, for ECMWF and NCEP respectively (see Table S1). When we calculate the MJO impact during NQBO when ENSO is included (Figure S6), we see that much of the shading east of 0° is not apparent. The presence of skill east of 0° when ENSO is not included may be due to small sample sizes of the NQBO events. Thus, when comparing MJO impacts between strong and neutral QBOs, it is important to keep sample size in mind. That being said, the statistical analysis we have applied here for requirements 1-3 account for the small sample sizes in the analysis.

4   Conclusions

The MJO is the dominant mode of intraseasonal variability in the tropics (Madden and Julian, 1971; Adames and Kim, 2016), and through its convective heating, modulates midlatitude weather, days to weeks after an MJO event (e.g. Vecchi 2004; Zhou et al. 2012; Henderson et al. 2016; Tseng et al. 2019). Recent research has shown that the QBO impacts MJO amplitude, propagation, and prediction skill (Son et al., 2017; Nishimoto and Yoden, 2017; Zhang and Zhang, 2018; Marshall et al., 2017; Lim et al., 2019) as well as modulates MJO teleconnections (e.g. Baggett et al. 2017; Mundhenk et al. 2018; Wang et al. 2018).

This raises the question as to whether the QBO also affects the prediction skill of MJO teleconnections. The goal of this study is to address this question through an examination of differences in Week 1-4 prediction skill between different combinations of QBO-MJO activity.

Through a STRIPES analysis (Jenney et al., 2019), we show that ECMWF and NCEP hindcasts are capable of simulating midlatitude MJO sensitivity, in a composite sense, out to Week 4 under different phases of the QBO. We then use these hindcasts to study enhanced S2S prediction skill following QBO-MJO activity. Increased prediction skill is determined from significant increases in spatial correlations of z500 for various QBO-MJO combinations. First, comparing active MJOs to inactive MJOs during different QBO phases (Requirement 1), we find that there is an MJO impact on midlatitude prediction skill during EQBO from North America to East Asia and during WQBO from the North Pacific through Europe. Second, when comparing the MJO impact between strong QBOs and NQBO (Requirement 2), we see that strong QBOs enhance the MJO impact on midlatitude subseasonal prediction skill from the Pacific to Europe. Lastly, to ensure that regions with enhanced MJO impacts during strong QBOs also have overall greater prediction skill following active MJO events compared to NQBO-MJO events (Requirement 3), and find that WQBOs enhance the MJO impact on midlatitude prediction skill as well as enhance overall prediction skill compared to NQBOs from the Pacific to Europe on Week 1-4 timescales in ECMWF and on Week 3-4 timescales in NCEP.

This study provides insight on improved prediction skill following different MJO-QBO combinations; however, more research is needed to determine the causal link between the MJO-QBO, midlatitude teleconnections and prediction skill. It is unclear whether enhanced midlatitude prediction skill is a consequence of the QBO's direct effects on the tropical environment in which the MJO forms and/or through the modulation of the atmospheric basic state through which Rossby waves propagate.

We motivated this study by suggesting that enhanced MJO prediction following EQBO (Marshall et al., 2017; Lim et al., 2019; Abhik and Hendon, 2019) may also lead to enhanced midlatitude prediction skill following MJOs during EQBO. However, we find that *both* EQBO and WQBO lead to an enhanced MJO impact in these hindcasts rather than only EQBO. Enhanced skill following MJOs during both EQBO and WQBO may partially be explained by Kim et al. (2019) who find no significant impact of the QBO on MJO prediction skill within the Subseasonal Experiment database (SubX; Pegion et al. 2019), which suggests these models do not differentiate between the two phases of the QBO. In addition, we find that there is an enhanced MJO impact as well as increased prediction skill following MJO events during WQBO compared to NQBO on S2S timescales, specifically over the North Atlantic out to Week 4 lead times. This result is supported by a growing body of work suggesting the importance of WQBO on MJO teleconnections. For example, recent work has shown that the North Atlantic Oscillation and MJO relationship is stronger during WQBO (Feng and Lin, 2019; Song and Wu, 2020) and prediction skill of Atmospheric Rivers over Alaska is enhanced following WQBO-MJO (Baggett et al., 2017). Finally, while strong ENSO events were removed from our analysis in an attempt to separate the effects of the QBO from those of ENSO, an ENSO influence may still remain. In addition, the sample sizes for MJO-QBO activity are not large (Table S1), although we attempt to account for this through statistical analysis. Even so, this work suggests that both phases of the QBO may impact prediction skill of MJO teleconnections and should be considered in future studies.

330 *Data availability.* ERA-I Reanalysis data are provided by the European Centre for Medium Range Forecasts (https://www.ecmwf.int/en/forecasts/-datasets/reanalysis-datasets/era-interim/). The ECMWF and NCEP hindcasts are provided by the World Weather Research Program/World Climate Research Program (ECMWF: https://apps.ecmwf.int/datasets/data/s2s/ and NCEP: http://iridl.ldeo.columbia.edu/SOURCES/.Models/.SubX/). RMM Index data are provided by the Australian Bureau of Meteorology (http://www.bom.gov.au/climate/mjo/graphics/rmm.74toRealtime.txt). The Nino3.4 Index data was provided by NOAA/OAR/ESRL Boulder, CO (climatedataguide.ucar.edu/climate-data/).

335 *Author contributions.* Kirsten Mayer led the collection of the data, the data analysis and writing of the manuscript. Elizabeth Barnes aided with the experimental design, analysis techniques and writing of the manuscript.

Competing interests. The authors declare that they have no competing interests.

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