Responses to the Editor:

1) Now that you have added tracks to figures 3-5, you need to adapt the captions accordingly. I would also discuss the tracks explicitly in the text, as you do in the reply.

Thank you for the suggestion, the figure captions have been modified and the tracks are now explicitly discussed in the text.

2) In response to Point 6 of Reviewer 1, you computed correlations for detrended time series. As far as I can see this new result has not been included in the text and I would suggest to do so.

Thank you for the suggestion, we have added this sentence to the results section “The MLR models were also done with detrended decadal series of off-season TCs, and detrended decadal series of the different climate variability and change factors, and similar results were found.” Since are the results are quite similar, we think that writing about those results would be repetitive.

3) I would suggest that you spend a few sentences on discussing limitations of ICOADS.
(Reviewer 1, point 4)

Thank you for the suggestion, we have added this sentence in the data section: “It is important to note that the ICOADS data set has some key limitations, like data coverage been sparse and limited corrections that account for changes in observing practices.”

4) I would suggest that you make the reader aware of the issues with detecting tropical depressions in the pre-satellite era (Reviewer 1, point 2).

Thank you for the suggestion, we have added this sentence to the methods section: “Another issue related to pre-satellite era TC track data is the undercount of weaker tropical depressions, since the detection and classification of those weaker storms that showed poor organization was probably more difficult before 1966 (Moon et al., 2019).”

Non-public comments to the Author:

Next time you submit a paper to this journal I would ask you to clearly state in your reply what you have changed in the manuscript rather than just answering the questions and leaving it to the editor to check what you changed. Also providing a version of the manuscript with track changes helps. Overall, I feel you could try a little harder to make changes in response to the reviewers (in most cases useful) suggestions. There is quite a number of points you rejected without even making minor changes to the text. It is not a reason for me not to accept the paper in the end, but it would seem a better use of everyone’s time to be more responsive.

Thank you for those suggestions. We want to apologize for nor submitting a manuscript with track changes, that would have made things easier for you. Next time I submit a paper to WCD I
will make sure to address all of the minor comments with changes in the text and I will also provide a track changes version of the manuscript.
Increasing Frequency in Off-Season Tropical Cyclones and its relation to Climate Variability and Change

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Correspondence: José J. Hernández Ayala (jose.hernandezayala@sonoma.edu)

Abstract. This article analyzes the relationship between off-season tropical cyclone (TC) frequency and climate variability and change for the Pacific and Atlantic Ocean basins. TC track data was used to extract the off-season storms for the 1900-2019 period. TC counts were aggregated by decade and the number of storms for the first six decades (pre-satellite era) was adjusted. Mann-Kendall non-parametric tests were used to identify trends in decadal TC counts and multiple linear regression models (MRL) were used to test if climatic variability or climate change factors explained the trends in off-season storms. MRL stepwise procedures were implemented to identify the climate variability and change factors that explained most of the variability in off-season TC frequency. A total of 713 TCs were identified as occurring earlier or later than their peak seasons, most during the month of May and in the West Pacific and South Pacific basins. The East Pacific (EP), North Atlantic (NA) and West Pacific (WP) basins exhibit significant increasing trends in decadal off-season TC frequency. MRL results show that trends in sea surface temperature, global mean surface temperature, and cloud cover explain most of the increasing trend in decadal off-season TC counts in the EP, NA, and WP basins. Stepwise MLR results also identified climate change variables as the dominant forces behind increasing trends in off-season TC decadal counts, yet they also showed that climate variability factors like El Niño-Southern Oscillation, the Atlantic Multidecadal Oscillation, and the Interdecadal Pacific Oscillation also account for a portion of the variability.

Keywords: Tropical Cyclones; Hurricane Season; Climate Variability; Climate Change

1. Introduction

Increasingly, scientific evidence has shown a link between tropical cyclones (TC) and global warming, especially following the dramatic rise in both the intensity and frequency of storms during the first two decades of the present century (Goldenberg et al., 2001; Holland and Webster, 2007). Scientific studies (Landsea, 2005; Emanuel, 2005; Trenberth and Shea, 2006; Trenberth, 2007) are not in agreement as to whether sea surface temperatures have a measurable effect on the frequency of tropical cyclones and other studies (Camargo and Sobel, 2005; Nogueira and Kim, 2007; Mahala et al., 2015; Zhao et al., 2018) have evaluated cyclonic activity on a time scale longer than interannual and have associated it with variability in the El Niño Southern Oscillation (ENSO), the Atlantic Multidecadal Oscillation (AMO) and the Interdecadal Pacific Oscillation (IPO). However, little is known about the changes in the frequency of off-season TCs, storms that occur before and after the peak TC season months, and their connections to climate variability and change.

A number of recent papers (Wang and Lee, 2008; Knutson et al., 2010; Emanuel, 2013) have documented global increases in the proportion of very intense cyclones as well as latitudinal trends in maximum tropical cyclone (TC) intensity, which are consistent with future climate projections. A detailed review of the behavior of TCs (Walsh et al., 2019) concluded that it remains uncertain whether past changes in TC activity have exceeded the variability expected from natural causes,
while concerns remain about the temporal homogeneity of the best record (Landsea et al., 2006; Mann et al., 2007). Another study (Mann et al., 2009) found that recent increases in the frequency of intense TCs in the North Atlantic (NA) were the product of reinforcing effects, such as La Niña-like climate conditions and relative tropical Atlantic warming. Yet, no study has focused on examining changing trends in off-season TC frequency and its relation to natural variability or climate change.

A synthesis (Christensen et al., 2013) of the then-available regional projections of future TC climatology for 2081–2100 in relation to 2000–2019, for a business as usual emissions scenario similar to A1B, shows that worldwide the consensus projection was for decreases in TC numbers by approximately 5–30%, increased frequency of Category 4 and 5 storms between 0 and 25%, an increase by a small percentage in the typical maximum intensity of life, and an increase in TC rainfall amounts by 5–20%. Nevertheless, it is clear that there is great uncertainty about these projections. Such projections do not consider changes in off-season TC development in any of the basins where TCs form.

Several climatic reconstructions have been performed (Bradley et al., 2006; Mann et al., 2009) using proxy data by collecting sediments from the impact of hurricanes in the period 500–1850 and then calculated estimates from the statistical model of the activity of tropical cyclones based on modern instrumental weather indexes for the period (1851–2006). In analyzing these results and comparing them with the cyclone seasons fixed by the World Meteorological Organization, the hurricane season (tropical depressions, tropical storms and hurricanes) in the Atlantic Ocean was fixed as June 1 to November 30 in 1960, yet we observe a significant variability in off-season TC occurrence before/after the hurricane season after the 1960s.

Even though it not uncommon for TCs to develop outside of their peak TC season months, there is a need to examine trends in the number of storms that are forming during low activity months. The formation of the extratropical storm Andrea on May 20, 2019 marks the decade of 2010 as that with the greatest number of tropical cyclones in the Atlantic Ocean before or after the hurricane season dates established by official bodies like the World Meteorological Organization (WMO) and the National Oceanic and Atmospheric Administration (NOAA). The frequency of TCs in the North Atlantic basin has been found to be influenced by fluctuations in teleconnections such as ENSO and AMO (Trenberth et al., 2006; Nogueira et al., 2013). However, human-induced climate change manifested as higher sea surface temperatures (SST) and increasing evaporation rates in the tropical and sub-tropical North Atlantic basin could also be related to the higher frequency of off-season tropical or extratropical cyclone occurrences in more recent decades. That increasing trend in SSTs in the Atlantic and other ocean basins and its relation to out-ff-season TC occurrences during the last century has not been thoroughly examined by the scientific community.

This study aims to determine if off-season TCs have increased in their frequency since the 1900 and if that increment in the number pre and post off-season storms could be associated with normal climatic variability or climate change. The total number of out off-season TCs per decade for the North Atlantic (NA), West Pacific (WP), East Pacific (EP) and South Pacific (SP) ocean basins where analyzed to determine if any of the basins experienced an increase in the number of off-season tropical/extratropical cyclones over time that could be associated to climatic variability or climate change. The Indian Ocean basins were not included in this analysis due to limited data availability. Ocean basins that were found to have statistically significant trends were then analyzed further with multiple linear regression models (MRL) and regression stepwise procedures to determine if climate variability or change factors could explain increasing trends in off-season TC frequency over time.
Six-hourly TC track data for all storms across all ocean basins were obtained from the International Best Track Archive for Climate Stewardship (IBTrACS) (Knapp et al., 2018) and all TCs that occurred at or after 1900 were extracted. The TC tracks were then extracted for the northern hemisphere basins that include the East Pacific (EP), the North Atlantic (NA) and the West Pacific (WP) and for the southern hemisphere basin in the South Pacific (SP) (Fig 1a). The off-season TCs were then aggregated by decades in order to identify decadal variability or trends in total storm counts at the individual basin scales. Off-season TCs were defined as storms that occurred in the three months before and after the six-month period of peak cyclone activity in the basin.

The monthly frequency of TCs for each basin were analyzed for the entire period and based on that analysis we determined that off-season TCs that occurred during the three months (Mar-Apr-May) before June 1st were pre-season and the three months (Dec-Jan-Feb) after November 30th were post-season in the northern hemisphere basins (NA, EP and WP). In the southern hemisphere, the three months before (Aug-Sep-Oct) November 1st were classified as pre-season and the three months (May-Jun-Jul) after April 30th were classified as post off-season in the southern hemisphere basins (SP) (Fig 1b). Pre-season and post-season decadal time-series for the Northern/Southern hemisphere and individual basins were then constructed to calculate the total number of TCs per-decade from 1900 to 2019.

The climate variability indexes of ENSO (Niño 3.4), AMO (Trenberth et al., 2019) were respectively obtained from the National Oceanic and Atmospheric Administration (NOAA) Physical Sciences Lab and the National Center for Atmospheric Research. The IPO index was obtained from the NOAA Physical Sciences Laboratory (Henley et al., 2015). The variables associated with anthropogenic climate change used in this study were sea surface temperature (SST), global mean surface temperature (GMST) and cloud cover (CC). SST data were obtained from the HadISST1 1° reconstruction, GMST data were accessed from the GISTEMP v4 and CC data was acquired from the International Comprehensive Ocean-Atmosphere Data Set (ICOADS) v2.5, all for the 1900-2019 period. It is important to note that the ICOADS data set has some key limitations, like data coverage been sparse and limited corrections that account for changes in observing practices. A decadal average was calculated for all of the climate variability and change variables in order to use them as predictors of decadal TC total counts (Table 1).

Table 1. Tropical cyclone, climate change and variability variables used in this study.

<table>
<thead>
<tr>
<th>Tropical Cyclone Counts</th>
<th>Abbreviations</th>
<th>Units</th>
</tr>
</thead>
<tbody>
<tr>
<td>TCs</td>
<td>Decadal Total Counts</td>
<td></td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th>Climate Change Variables</th>
<th>Abbreviations</th>
<th>Units</th>
</tr>
</thead>
<tbody>
<tr>
<td>Sea Surface Temperature</td>
<td>SST</td>
<td>°</td>
</tr>
<tr>
<td>Global Mean Surface Temperature</td>
<td>GMST</td>
<td>°</td>
</tr>
<tr>
<td>Cloud Cover</td>
<td>CC</td>
<td>Oktas</td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th>Climate Variability Variables</th>
<th>Abbreviations</th>
<th>Units</th>
</tr>
</thead>
<tbody>
<tr>
<td>El Niño Southern Oscillation</td>
<td>ENSO 3.4</td>
<td>°SST anomalies index</td>
</tr>
<tr>
<td>Interdecadal Pacific Oscillation</td>
<td>IPO</td>
<td>°SST anomalies index</td>
</tr>
<tr>
<td>Atlantic Multi-decadal Oscillation</td>
<td>AMO</td>
<td>°SST anomalies index</td>
</tr>
</tbody>
</table>
3. TC Adjustment Method

TC counts before 1966 (pre-satellite era) are incomplete (Mann et al., 2007; Landsea, 2007) since a lot of storms that didn’t make landfall weren’t recorded, so in order to make any comparisons between the earlier and later decades, the series for each basin need to be adjusted accordingly. Another issue related to pre-satellite era TC track data is the undercount of weaker tropical depressions, since the detection and classification of those weaker storms that showed poor organization was probably more difficult before 1966 (Moon et al., 2019). The average landfall percentage of TCs were calculated for the periods 1900-1965 (pre-satellite) and 1966-2019 (satellite era and new TC monitoring technologies available) in order to determine the share of storms that made landfall in both periods. The percentage of landfalling TCs is expected to be higher in the 1900-65 period since a higher number of storms that remained over the ocean were not reported, so the landfall percentage of the pre-satellite period is then adjusted so that it matches the 1966-2019 post-satellite period.

To obtain the estimated number of missing TCs for the 1900-65 period, the number of total storms in the pre-satellite period is increased until its landfall percentage is equal to the one in the post-satellite era. The total number of additional TCs that resulted in the landfall percentages between the two periods to be the same or near equal are then divided by the 7 decades of the pre-satellite era and then the number of extra storms for each decade is multiplied by the percentage of off-season storms for each basin and that resulting number is then added then to each of the individual decades between 1900 and 1969. In a previous study (Landsea, 2007), this method was applied to adjust TC counts in the North Atlantic to determine if the basin has experienced an increasing trend in annual TC frequency since the 1900, and its results show that after adjusting the tropical storm counts no trends were found.

Here we show how this TC series adjustment method was applied to the total TC count for the NA basin for the 1900-2019 period. First, we calculate the landfall percentage for the pre-satellite period 1900-65 by dividing the number of landfalling TCs (LFTCs) with the total number of storms (TTCs) and multiply by 100 to get the landfall percentage, check the equations below:

\[
\frac{\text{LFTCs}}{\text{TTCs}} \times 100
\]

\[\text{Example: } \frac{479}{610} \times 100 = 78.5\% \quad (1)\]

Then calculate landfall % for the period post-satellite period 1966-2019,

\[
\frac{\text{LFTCs}}{\text{TTCs}} \times 100
\]

\[\text{Example: } \frac{583}{844} \times 100 = 69.1\% \quad (2)\]

Then artificially increase the number of TCs (+83 for the NA basin) until the landfall % of the 1900-65 period is equal to landfall % of the 1966-2019:

\[
\frac{\text{LFTCs}}{\text{TTCs} + \text{AddTCS}} \times 100
\]

\[\text{Example: } \frac{479}{610 + 83} \times 100 = 69.1\% \quad (3)\]

Then calculate the percentage (OffP) of off-season TCs (OffTCs) by dividing it by total number of TCs:

4
\[
\text{(OffTCs / TTCs) \times 100}
\]

*Example*: \((67/1454) \times 100 = 4.61\%

Then divide additional TCs (83) by the number of decades between 1900 and 1969 (7) and then multiply by the off-season TC percentage (.0461)

\[
\text{(AddTCs/Decades) \times OffP}
\]

*Example*: \((83/7) \times .0461 = 0.54

In the case of the NA, we determined that by using the above TC series adjustment method the basin would get an additional 0.54 off-season TCs for each of the seven decades that go from the 1900 to 1969. Finally, the additional 0.54 TCs per decade will be divided between pre and post off-season TCs by multiplying the added storms with the respective percentage of pre/post off season cyclones:

\[
\text{DecOffTCs} \times \text{Percentage/Post Season}
\]

*Example*: \(0.54/0.62 = 0.33 \text{ and } 0.54/0.38 = 0.21

The pre off-season decades of the NA basin before 1970 will get an additional 0.33 TCs and the post off-season decades will get 0.21 more storms. This off-season TC adjustment method was applied to the other five basins.

### 4. Statistical Methods & Models

Mann-Kendall (MK) tests for trends (Mann, 1945; McLeod, 2005) were applied to all the off-season TC decadal series for all basins in order to determine if the frequency of storms has increased or decreased over time. This test has the advantage of not assuming any special form for the distribution function of the data, while having a power nearly as high as their parametric equivalents and that is why its use is highly recommended by the World Meteorological Organization (Hipel and McLeod, 2005).

The decadal series that were then found to have a significant trend based on the MK results were then furtherly analyzed by applying a series of multiple linear regression models (MLR). MLR were used to model the decadal count of off-season TCs for basins that showed increasing or decreasing trends in storm numbers to test if covariates associated with climatic variability and climate change explained off-season TC frequency. MLR attempts to model the relationship between two or more explanatory variables and a response variable by fitting a linear equation to observed data.

Three different MLR models were run for each off-season TC series that exhibited a statistically significant trend, one MLR model with the climate change variables (SST, GMST & CC) as predictors, another model with just the climate variability factors (ENSO, AMO & IOD) and a final model with all of the variables included. Then the three MLR models (pre-season, post-season and off-season) were run for each of the basins with increasing trends in off-season TCs, the best models (highest adjusted R-squared and lowest p-value) were then selected for each of the series. The MLR models were run in The R Project for Statistical Computing using the biglm package.

Finally, stepwise selection MLR models were used to identify the climate variability or change factors making the most statistically significant contributions to off-season increasing TC frequency. Here we use stepwise selection which is a combination of the forward and backward procedures where you start with no predictors, then sequentially add the most contributive predictors. After adding each new variable, it removes the variables that no longer provide an improvement in the model fit (James et al., 2014; Bruce and Bruce, 2017). The MLR and stepwise for the off-season TC count series for
each of the basins with significant increasing trends were run in The R Project for Statistical Computing using the MASS package (Venables and Ripley, 2002).

5. Results & Discussion

When analyzing the number of TCs for all basins for the 1900-2019 period we found that 713 off-season storms occurred during that time, most during the months of May (NH pre-season and SH post-season) with 430 and December (NH post-season) with 341 (Figure 1a, 1b). When looking at the count of off-season TCs per basin we found that as expected the West Pacific (611) and South Pacific (85) accounted for 81.3% of all off-season storm occurrences. When grouping the basins between northern and southern hemispheres, we find that 89% of all off-season TCs occurred north of the equator for the 1900-2019 period (Figure 1a, 1b). The North Atlantic and East Pacific basins were found to be the ones with the lowest numbers of off-season TCs when compared to the other two Pacific basins.

Figure 1. Tracks of all out off-season TCs (a) and the number of storms per month for the WP (b), SP (c), EP (d) and NA (e) basins for the 1900-2019 period.
As shown in Figure 2, even after adding the estimated missing storms before the 1970 decade, most basins experienced their highest number of out of season TCs (pre or post) in decades at or after 1960-69. The 1960-69 decade for the northern hemisphere basins (WP, NA and EP) was found to be the one with the highest number of pre off-season TCs with 69 and the 1950-1959 decade was identified as the one with the most post off-season storms with 68 (Figure 2a, 2b, 2c and 2d). When examining TC counts for all basins individually, we found that the NA and EP basins had their most active decades after 1970 and that the WP and SP basins experienced their highest storm count decade after 1960 (Figure 2c, 2d). It is important to note that these results already reflect the additional TCs that were added to the pre-satellite era.

**Figure 2.** Adjusted decadal count of all observed and estimated off-season TCs for the WP (a), NA (b), EP (c) and SP (d) ocean basins for the 1900-2019 period.

The Mann-Kendall non-parametric tests for trends for all basins show that three basins exhibited statistically significant increasing trends in adjusted decadal off-season TC counts for the 1900-2019 period (Table 2). The basins with statistically significant increasing trends were the EP (pre and off-season), NA (pre, post and off-season) and the WP (pre, post and off-season). The EP basin shows an increasing trend in pre- and off-season TCs that is more evident from the 1950s to the present (Figure 2d), while the increasing trend in the NA basin can be observed from the 1970s to 2019.

The increasing trend in off-season TCs in the WP basin is more evident from the 1900 to 1969 (Figure 2a), which was during the pre-satellite era where missing TCs were added to the series. However, no trend is found in the WP basin if the decadal counts are analyzed from the 1970s to the
present. In the post-satellite era in the WP basin, the 1990-1999 decade was identified as the one with most off-season TCs, however the two following decades exhibited a decreasing trend. The EP and NA basins show significant increasing trends in off-season TC counts (Table 2). Opposite to the trends identified in the WP basin, the EP and NA also show increasing decadal counts after the 1960s and 1970s decades. The SP basin also exhibited a positive Tau coefficient, yet it was statistically insignificant (Figure 2d).

Table 2. Results of Mann-Kendall trend tests for the 1900-2019 period for all ocean basins where TCs form.

<table>
<thead>
<tr>
<th>Pre TCs</th>
<th>Tau S</th>
<th>P-value</th>
<th>Post TCs</th>
<th>Tau S</th>
<th>P-value</th>
<th>Off TCs</th>
<th>Tau S</th>
<th>P-value</th>
</tr>
</thead>
<tbody>
<tr>
<td>EP</td>
<td>0.746</td>
<td>0.002</td>
<td>EP</td>
<td>0.098</td>
<td>0.723</td>
<td>EP</td>
<td>0.679</td>
<td>0.004</td>
</tr>
<tr>
<td>NA</td>
<td>0.572</td>
<td>0.015</td>
<td>NA</td>
<td>0.485</td>
<td>0.042</td>
<td>NA</td>
<td>0.554</td>
<td>0.016</td>
</tr>
<tr>
<td>SP</td>
<td>0.048</td>
<td>0.889</td>
<td>SP</td>
<td>0.015</td>
<td>1.000</td>
<td>SP</td>
<td>0.061</td>
<td>0.836</td>
</tr>
<tr>
<td>WP</td>
<td>0.554</td>
<td>0.016</td>
<td>WP</td>
<td>0.485</td>
<td>0.034</td>
<td>WP</td>
<td>0.534</td>
<td>0.019</td>
</tr>
</tbody>
</table>

Significant trends in bold.

MLR models were run on the basins that exhibited statistically significant (< 0.05) increasing trends in decadal total off-season TC counts over time and here we report the best models for each of the series. The MLR results show that the statistically significant increasing trends in TC frequency for the EP (pre and off-season) and WP basins is best explained by climate change factors SST, GMST and CC at the 0.05 significance level (Table 3). Climate change factors accounted for 56% (pre-season) and 52% (off-season) of the increasing trend in TC counts for the EP basin. In the WP basin climate change factors explained 55% (pre-season), 64% (post-season) and 68% (off-season) of the trends in off-season TCs. Increasing trends in SSTs, GMST and moisture (CC) outside of the prime months of tropical storm development could promote more optimum conditions for higher off-season TC occurrences (Klozback, 2006; Hansen et al., 2010). The MLR models were also done with detrended decadal series of off-season TCs, and detrended decadal series of the different climate variability and change factors, and similar results were found.

Table 3. Best multiple linear regression models (MLR) for basins with statistically significant increasing trends in off-season TCs.

<table>
<thead>
<tr>
<th>Model</th>
<th>Multiple R-squared</th>
<th>Adjusted R²</th>
<th>Factors</th>
<th>p-value</th>
</tr>
</thead>
<tbody>
<tr>
<td>EP pre-season</td>
<td>0.682</td>
<td>0.563</td>
<td>SST, GMST &amp; CC</td>
<td>0.021</td>
</tr>
<tr>
<td>EP off-season</td>
<td>0.653</td>
<td>0.522</td>
<td>SST, GMST &amp; CC</td>
<td>0.030</td>
</tr>
<tr>
<td>NA pre-season</td>
<td>0.622</td>
<td>0.481</td>
<td>SST, GMST &amp; CC</td>
<td>0.041</td>
</tr>
<tr>
<td>NA post-season</td>
<td>0.687</td>
<td>0.427</td>
<td>ENSO &amp; AMO</td>
<td>0.130</td>
</tr>
<tr>
<td>NA off-season</td>
<td>0.552</td>
<td>0.384</td>
<td>SST, GMST &amp; CC</td>
<td>0.070</td>
</tr>
<tr>
<td>WP pre-season</td>
<td>0.673</td>
<td>0.551</td>
<td>SST, GMST &amp; CC</td>
<td>0.020</td>
</tr>
<tr>
<td>WP post-season</td>
<td>0.742</td>
<td>0.645</td>
<td>SST, GMST &amp; CC</td>
<td>0.000</td>
</tr>
<tr>
<td>WP off-season</td>
<td>0.774</td>
<td>0.689</td>
<td>SST, GMST &amp; CC</td>
<td>0.005</td>
</tr>
</tbody>
</table>

The climate variability factors (ENSO, AMO & IOD) did not exhibit statistically significant relationships with increasing off-season TC counts, which shows that natural variability does not explain the incrementing number of storms in the EP and WP basins. MLR model results for the NA basin also showed the climate change variables accounting for 48% (pre-season) and 38% (off-season) of the increasing trend in TCs (Table 3). However, the MLR model results for the post-
season months in the NA basin showed that the climate variability variables (ENSO & AMO) accounted for 42% of the increasing trend in TCs, yet the model was not found to be statistically significant. It is well known that cold phases of ENSO (La Niña) and warm phases of AMO tend to be associated with higher TC frequency in the North Atlantic ocean (Tang and Neelin, 2004; Briggs, 2008) and this could explain why those teleconnections were found to have the most significant influence on post-season TC frequency in the NA basin. Yet, it is important to note that in most basin series, including the NA, climate change variables explained more of the off-season TC increasing trend than the climate variability factors.

Stepwise MLR model results showed that climate change factors (SST, GMST & CC) were among the selected variables that explained most of the increasing trend in off-season TCs for all basins analyzed (Table 4). In the EP basin, SST, ENSO, and CC, accounted for 69% (pre-season) and 65% (off-season) of the increasing trend in TCs. In the NA basin, the stepwise procedure selected CC as the sole climate change factor that explained 52% (pre-season) and 40% (off-season) of the rising frequency in TC counts. However, CC & AMO were selected as the variables that explained (43%) most of the variability in TC frequency during the post-season months in the NA basin. Stepwise procedure results for the WP basin show that climate change and variability factors were selected as the best predictors of TC frequency, with GMST and CC accounting for 57% (pre-season), CC, GMST, ENSO and IPO explaining 72% (post-season) and 74% (off-season) of the variability of TCs.

Table 4. Stepwise multiple linear regression models (MLR) for basins with statistically significant increasing trends in off-season TCs.

<table>
<thead>
<tr>
<th>Model</th>
<th>R-squared</th>
<th>Adjusted R²</th>
<th>Factors</th>
<th>p-value</th>
</tr>
</thead>
<tbody>
<tr>
<td>EP pre-season</td>
<td>0.777</td>
<td>0.694</td>
<td>SST, ENSO &amp; CC</td>
<td>0.005</td>
</tr>
<tr>
<td>EP off-season</td>
<td>0.747</td>
<td>0.652</td>
<td>SST, ENSO &amp; CC</td>
<td>0.008</td>
</tr>
<tr>
<td>NA pre-season</td>
<td>0.569</td>
<td>0.526</td>
<td>CC</td>
<td>0.004</td>
</tr>
<tr>
<td>NA post-season</td>
<td>0.687</td>
<td>0.427</td>
<td>CC &amp; AMO</td>
<td>0.098</td>
</tr>
<tr>
<td>NA off-season</td>
<td>0.460</td>
<td>0.406</td>
<td>CC</td>
<td>0.015</td>
</tr>
<tr>
<td>WP pre-season</td>
<td>0.655</td>
<td>0.578</td>
<td>GMST &amp; CC</td>
<td>0.008</td>
</tr>
<tr>
<td>WP post-season</td>
<td>0.826</td>
<td>0.726</td>
<td>CC, GMST, ENSO &amp; IPO</td>
<td>0.008</td>
</tr>
<tr>
<td>WP off-season</td>
<td>0.839</td>
<td>0.747</td>
<td>CC, GMST, ENSO &amp; IPO</td>
<td>0.006</td>
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</tbody>
</table>

The EP experienced a steady increase in off-season TC total counts from 1900 to 2019 at a rate of 1.1 additional storms per decade. The decadal off-season total TC count series for the EP basin closely resembles the increasing trend in average SSTs and CCs (Fig 3a, 3c). When the EP off-season TC tracks are examined, it shows that most storms have formed in areas that have experienced statistically significant increasing trends in SST and CC (Fig 3d, 3e). The correlation between off-season TCs in the EP basin and ENSO is not as clear as the one between SST and CC, with some mostly warm ENSO decades like the 1990-1999 exhibiting lower storm counts and other periods with cooler phases dominating showing a higher number of cyclones. When SST patterns for areas in the EP basin where TCs develop are examined over time, we find that most tropical/sub-tropical ocean waters have experienced a statistically significant increasing trend in ocean surface temperatures from 1900 to 2019 (Fig 3d). Similar to other studies (Hansen et al., 2010), we find that the EP tropical ocean surfaces have increased by 0.051 degrees C° per decade. When CC patterns are examined, we find that it has also experienced a statistically significant increasing trend in some areas in the EP basin (Fig 3e).
The decadal off-season total TC count series for the NA basin closely resembles the increasing trend in average SSTs and CCs (Fig 4a, 4c). The NA decadal series shows a steady increase in off-season TC total counts from 1900 to 2019 at a rate of 0.7 additional storms per decade and an SST increasing trend of 0.055 C° per decade. Both the average decadal SST and CC series coincide with the peaks and valleys in off-season TC counts for the NA basin, with the 1950-1959 showing a high number of storms associated with high mean SSTs and CCs while the drop in storm counts in the 1960-1969 decade matches a drastic drop in ocean surface temperatures (Figures 4a, 4c). The off-season TC tracks in the NA basin also formed in areas that exhibited increasing trends in SST and CC (Figure 4d, 4e). Even though average SSTs increase to 0.135 C° per decade from 1970 to 2019, off-season TC total counts went down in the 1990-1999 and 2010-2019 decades, with the decade in between (2000-2009) exhibiting the highest number of off-season TCs (14) of all decades examined. However, it is important to note that 5 out of the 6 decades with the most off-season TCs in the NA basin occurred after the 1970s.

When North Atlantic SSTs are examined in areas where TCs form, we found that ocean surface temperatures have increased at a rate of 0.055 degrees C° per decade for the off-season months of Dec-March (Fig 4d). When CC patterns are examined, we find that it has also experienced a statistically significant increasing trend of 0.06 oktas (eighths of the sky that are covered in clouds) per decade in the North Atlantic basin since the 1900 (Figure 4e). If the NA pre/post off-season series is modified to begin in the 1960s, we find that SSTs have increased at a decadal rate of 0.082 C° per decade at a rate of 1.2 additional storms per decade. Overall, these results suggest that increasing trends in SSTs, which also drive increasing trends in evaporation rates associated with high CCs, are the physical mechanisms behind most of the recent increase in the total number of out of season TCs in the NA basin. The correlation between off-season TCs in the NA basin and AMO is not as clear as the one between SST and CC, with some warm AMO phases between 1930-1959 exhibiting lower storm counts while some cooler phases (1970-89) showing a higher number of cyclones.
Figure 3. Decadal TC counts for the EP off-seasons and decadal average SSTs (a), decadal TC counts for the EP off-seasons and decadal average ENSO 3.4 (b), decadal TC counts for the EP off-seasons and decadal average Correlation between Time and Dec-May averaged CC (c), correlation between Time and Dec-May averaged SST and the off-season TC tracks (d), and correlation between Time and Dec-May averaged CC (oktas) and the off-season TC tracks for the 1900-2019 period (e).
Figure 4. Decadal TC counts for the NA off-seasons and decadal average SSTs (a), decadal TC counts for the NA off-seasons and decadal average AMO (b), decadal TC counts for the NA off-seasons and decadal average SST (C°) and the off-season TC tracks for the 1900-2019 period (d) and correlation between Time and Dec-May averaged CC (oktas) and the off-season TC tracks for the 1900-2019 period (e).

The decadal off-season total TC count series for the WP basin closely resembles the increasing trend in GMST (Fig 5a). However, the WP basin experienced the highest count of off-season TCs in the 1960-69 decade, not in the more recent decades like the EP and NA basins. More importantly, if trend analysis for off-season TC counts is done from 1960-2019 in the WP basin, we find no statistically significant increasing or decreasing trend. However, it is important to note that four out of the five decades with most off-season TCs in the WP basin occurred after 1960. However, the 2010-2019 decade was identified as the period with the lowest total number of off-season TCs even though increasing trends in mean SST, GMST and CC continued (Fig 5a, 5d and 5e). Off-season TC tracks in the WP basin also correlate spatially with areas that show increasing trends in SST and CC (Fig 5d, 5e). The decreasing number of off-season TCs in the last two decades coincided with a negative phase of the IPO, which suggests that TC frequency in the WP basin is influenced by fluctuations in the IPO (Fig 5c), whose recent negative phase since 1998 resembles La Niña–like SST anomaly patterns (Zhao et al., 2018). Even though most of the variability in off-season TC frequency...
in the WP basin can be explained by climate change trends in GMST, SST and CC, the rest of the variance in TCs is accounted by fluctuations in the IPO and ENSO teleconnections.

Figure 5. Decadal TC counts for the WP off-seasons and decadal average SSTs (a), decadal TC counts for the WP off-seasons and decadal average AMO (b), decadal TC counts for the WP off-seasons and decadal average SSt (C°) and the TC tracks for the 1900-2019 period (d) and correlation between Time and Dec-May averaged CC (c), correlation between Time and Dec-May averaged SST (C°) and the off-season TC tracks for the 1900-2019 period (e).

6. Summary and concluding remarks

The frequency of TCs that developed outside of their prime season months were analyzed to determine if trends in higher storm totals in the Pacific and Atlantic Ocean basins were associated with natural variability, climate change or both. Adjusted off-season decadal TC total counts for six ocean basins were analyzed for the 1900-2019 period in order to determine if the number of storms have been increasing over time. Mann-Kendall tests for trends were done and the basins that exhibited statistically significant increasing trends were then furtherly analyzed using multiple linear regression models and stepwise procedures to determine if those trends could be explained by fluctuations associated with climate variability, climate change trends or a combination of both.
The main results of this study suggest that decadal total off-season (pre/post) TC counts have significantly increased over time since the 1900 in the East Pacific (EP), North Atlantic (NA) and West Pacific (WP) basins. The EP and NA basins exhibited statistically significant increasing trends even if the analysis was done from the 1960s instead of the 1900. The WP basin showed an overall increasing trend in the total number of off-season TCs per decade, yet if the analysis is done from the 1960s to the present, no statistically significant increasing trend is found. However, the three basins that reflected an overall increase in decadal off-season TC frequency had their most active decades after the 1970s.

Results from the best multiple linear regression (MLR) models show that the increasing decadal count of off-season TCs has been found to be strongly associated with climate change trends in sea surface temperature (SST), global mean surface temperature (GMST) and cloud cover (CC) in all three basins (EP, NA and WP). The MLR model where climate variability variables (ENSO and AMO) explained most of the variance in off-season TC counts was in the storm decadal counts for the post-season months of the NA basin.

Results of the MLR stepwise procedures showed that the selected variables that accounted for most of the variability in off-season TCs for the EP basin were SST, CC and ENSO, while CC (pre-season and off-season) and AMO (post-season) were chosen as the best variables for the NA basin. The stepwise procedure identified the climate change trends in GMST and CC, and fluctuations in ENSO and IPO as the variables that accounted for most of the variability in decadal off-season total TC counts in the WP basin.

The findings of this study suggest that trends in SST, GMST and CC associated with climate change are not only altering the frequency (Klotzbach, 2006; Saunders and Lea, 2006; Hansen et al., 2010) and intensity of TCs that develop during the peak months of the season, they are also altering the total number of storms that form in the off-season months (Dec-May), especially in the EP and NA basins. The results of this study have important implications for the NA and EP basins, if off-season TCs have been increasing in frequency since the 1900 we can expect that this trend associated with climate change would continue in future decades. This increasing number of off-season TCs could potentially impact societies in their path during times of the year when storms are least expected, possibly increasing environmental and economic impacts in areas that are already experiencing the effects of climate change exacerbated phenomena.

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this article are based on publicly available measurements from the National Oceanic and
Atmospheric Administration Physical Sciences Lab (PSL; https://psl.noaa.gov/data/timeseries/IPOTPI/).

References

Bradley, R.S., M.K. Hughes and M.E. Mann (2006) Authors were clear about hockey-stick

Bruce, Peter, and Andrew Bruce. 2017. Practical Statistics for Data Scientists. O'Reilly Media.

Camargo, S. J., & Sobel, A. H. (2005). Western North Pacific tropical cyclone intensity and
ENSO. Journal of Climate, 18(15), 2996-3006. https://doi.org/10.1175/JCLI3457.1

Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate
A., Xia Y., V.B. & P.M.M.) Cambridge; Cambridge University Press.


Emanuel, K.A. (2013) Downscaling CMIP5 climate models shows increased tropical cyclone
https://doi.org/10.1073/pnas.1301293110

in Atlantic hurricane activity: Causes and implications. Science, 293: 474–478. DOI:
10.1126/science.1060040


http://www. stats.uwo.ca/faculty/aim/1994Book/

Holland, G.J. and P.J. Webster (2007) Heightened tropical cyclone activity in the North Atlantic:
https://doi.org/10.1098/rsta.2007.2083

James, Gareth, Daniela Witten, Trevor Hastie, and Robert Tibshirani. 2014. An Introduction to

Track Archive for Climate Stewardship (IBTrACS) Project, Version 4. [indicate subset used].
NOAA National Centers for Environmental Information. non-governement domain
https://doi.org/10.25921/82vy-9e16

3: 157–163; doi:10.1038/ngeo779

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