Response to the Co-Editor #3

1) In-season trend: The text you added is sufficient to put your results into the context of other studies. However, there are some redundancies in this paragraph, and you should try to streamline it a bit.

Thank you for the suggestion. We have reduced the paragraph by removing the first two sentences and one of the sentences at the end that were redundant, and it now reads like this:

"Studies that have examined TC frequency overall have found increases in the number of most intense hurricanes [Wang and Lee, 2008; Knutson et al., 2010; Emanuel, 2013], yet no clear trend has been found when lower intensity TCs have been examined [Landsea, 2007]. The results of other studies show that there is no overall agreement on the relationship between SSTs and TC frequency (Landsea, 2005; Emanuel, 2005; Trenberth and Shea, 2006; Trenberth, 2007), yet some have found strong associations between TC variability and ENSO, AMO and IPO [Camargo and Sobel, 2005; Nogueira and Kim, 2007; Mahala et al., 2015; Zhao et al., 2018]. In this study we analyzed all off-season TCs and our results differ from those that have found no trend in overall TC frequency, since we found decadal increasing trends in the NA and EP basin in both the pre and post-satellite eras. The results presented here suggest that climate change trends like higher SSTs and more favorable moisture environments (CC) between the months of Dec to May in the NA and EP basins seem to be the major factors behind decadal increasing trends in off-season TCs."

2) Trends in climate indices: It is good that you added the table, but you now need to discuss these additional results in the text. Just stating that things are similar is not enough. What can we learn from the difference? I do not understand your explanation of the negative ENSO index in Figs. 3b and 5b and wonder whether there may be a mistake.

Thank you for the suggestion. We have now discussed similarities and differences between the original and detrended MLR models in the text and we have also added this sentence in the new version of the manuscript:

When the MLR results of the original and detrended series are compared (Table 3), we find that the models with the detrended series exhibit lower R squares than the MLR models with the original series, yet those models were still found to be statistically significant which suggests that the correlation between off-season TCs and climate change factors is strong even after decadal trends are removed.

Sorry for not elaborating a little bit more on why the ENSO seems to be mostly negative. Let me show you how the decadal average was calculated and that will help. We calculated the ENSO 3.4 decadal average by first calculating an average Dec-May ENSO 3.4 for each year and then we calculated the decadal average for each decade from that Dec-May ENSO 3.4 average for each year. Below we see the average Dec-May ENSO 3.4 for each year and when we calculate the decadal average, we get -0.327. We repeat that process with the other decades, and we get the ENSO 3.4 decadal average for each of the decades.

Year	ENSO 3.4
1900	0.4639176
1901	-0.4926524
1902	0.8448318

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1903 -0.5903594
1904 -0.6698176
1905 0.8433211
1906 -0.5974004
1907 -0.6810894
1908 -0.967404
1909 -1.426191
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3) Pre-satellite era: Again, it is good that you included the values in the table but now you need to discuss these findings in the text. You state in your reply that the MK test does not work for 60 years only. If that is the case, you should not present the results like this. Either use a different test or break down your timeseries in pentads instead. How is it possible that WP has negative trends in pre and post but positive in off? Please carefully check all numbers!

Sorry for the mistake in the WP off-season, the value was supposed to be a negative and not a positive. I also wanted to clarify something, it is not that the MK test doesn't work for 60 years it's that in the case of this study those 60 years are grouped into 6 decades, so when we do the test for the post-satellite era we are actually doing the test for 6 decades and not 60 years. In the new manuscript I have added a couple of sentences discussing the MK results for the post satellite era.

4) Data homogenization: The text you added creates at least some awareness of the problem. I think that should suffice for the moment.

Thank you.

5) Tropical depressions: I understand your point, but you should include this justification in the paper! You could state explicitly that it would be desirable to compare results including and excluding tropical depressions. Given the many uncertainties you are dealing with I suggest that you add 1-2 sentences to the conclusions with caveats (worse data quality in pre-satellite era, question of applying data homogenization for North Atlantic to other basins, lack of information about intensity).

I apologize for not including this in the previous manuscript. In the new version of the manuscript we have added a paragraph on limitations in the conclusion, here is the paragraph:

"One of the main limitations of this work was the inclusion of tropical depressions in the off-season TC analysis. If data on TC intensity were widely available for all off-season TCs, it would have been possible to exclude weaker tropical depressions from the analysis since the detection and classification of those storms was more difficult in the pre-satellite era. Other limitations of this study include the issues of worst data quality in the pre-satellite era, the problem of applying a universal missed TC adjustment method to all basins analyzed and the lack of information on TC intensity for many storms, especially in the pre-satellite era. "

6) TC tracks and SST change: Here you did not answer my question. Were there any changes over time in track or genesis location? Please check and include a respective statement in the text.

Sorry for that, based on spatial analysis looking at the mean center of the track and genesis location no major changes were detected. Statements regarding this have been added to the latest manuscript.

7) Relationship to cloud cover: The sentence you added is good, I think. There is, however, a misspelling in the reference (Klotzbach) here and elsewhere. Maybe just use "better" instead of "more optimum"?

Thank you for letting us know, we have corrected those errors.

8) ICOADS: What I meant is that you should explain how the ICOADS cloud estimates are produced and then discuss limitations (changes in data availability and methods to correct for that). The quote from Freeman et al. (2016) can help you to do this.

Thank you for being more specific as to what is it that you want to see in this section. We have added these sentences that explain how cloud estimates from ICOADS are produced and the limitations associated with the estimates:

"The CC estimates from ICOADS are obtained from voluntary observing ships that report CC in octas (eighths) ranging from 0 (completely clear sky) to 8 (completely overcast). These CC estimates are known to be temporally and spatially heterogenous with relatively high observational errors in some areas, yet increases in ocean coverage for clouds after the 1900 have been noted in the latest ICOADS 3.0 release [Eastman et al. 2011, Freeman et al. 2016; Aleksandrova et al. 2018]."

Increasing Frequency in Off-Season Tropical Cyclones and its

2 relation to Climate Variability and Change

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Abstract. This article analyzes the relationship between off-season tropical cyclone (TC) frequency and climate variability and change for the Pacific and Atlantic Ocean basins. TC track data was used to extract the off-season storms for the 1900-2019 period. TC counts were aggregated by decade and the number of storms for the first six decades (pre-satellite era) was adjusted. Mann-Kendall non-parametric tests were used to identify trends in decadal TC counts and multiple linear regression models (MRL) were used to test if climatic variability or climate change factors explained the trends in off-season storms. MRL stepwise procedures were implemented to identify the climate variability and change factors that explained most of the variability in off-season TC frequency. A total of 713 TCs were identified as occurring earlier or later than their peak seasons, most during the month of May and in the West Pacific and South Pacific basins. The East Pacific (EP), North Atlantic (NA) and West Pacific (WP) basins exhibit significant increasing trends in decadal off-season TC frequency. MRL results show that trends in sea surface temperature, global mean surface temperature, and cloud cover explain most of the increasing trend in decadal off-season TC counts in the EP, NA, and WP basins. Stepwise MLR results also identified climate change variables as the dominant forces behind increasing trends in off-season TC decadal counts, yet they also showed that climate variability factors like El Niño-Southern Oscillation, the Atlantic Multidecadal Oscillation, and the Interdecadal Pacific Oscillation also account for a portion of the variability.

Keywords: Tropical Cyclones; Hurricane Season; Climate Variability; Climate Change

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1. Introduction

Increasingly, scientific evidence has shown a link between tropical cyclones (TC) and global warming, especially following the dramatic rise in both the intensity and frequency of storms during the first two decades of the present century (Goldenberg et al., 2001; Holland and Webster, 2007). Scientific studies (Landsea, 2005; Emanuel, 2005; Trenberth and Shea, 2006; Trenberth, 2007) are not in agreement as to whether sea surface temperatures have a measurable effect on the frequency of tropical cyclones and other studies (Camargo and Sobel, 2005; Nogueira and Kim, 2007; Mahala et al., 2015; Zhao et al., 2018] have evaluated cyclonic activity on a time scale longer than interannual and have associated it with variability in the El Niño Southern Oscillation (ENSO), the Atlantic Multidecadal Oscillation (AMO) and the Interdecadal Pacific Oscillation (IPO). However, little is known about the changes in the frequency of off-season TCs, storms that occur before and after the peak TC season months, and their connections to climate variability and change.

A number of recent papers (Wang and Lee, 2008; Knutson et al., 2010; Emanuel, 2013) have

documented global increases in the proportion of very intense cyclones as well as latitudinal trends

in maximum tropical cyclone (TC) intensity, which are consistent with future climate projections. A

detailed review of the behavior of TCs (Walsh et al., 2019) concluded that it remains uncertain

50 whether past changes in TC activity have exceeded the variability expected from natural causes,

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- 51 while concerns remain about the temporal homogeneity of the best record (Landsea et al., 2006;
- 52 Mann et al., 2007). Another study (Mann et al., 2009) found that recent increases in the frequency
- 53 of intense TCs in the North Atlantic (NA) were the product of reinforcing effects, such as La Niña-
- 54 like climate conditions and relative tropical Atlantic warming. Yet, no study has focused on
- 55 examining changing trends in off-season TC frequency and its relation to natural variability or
- 56 climate change.

- 57 A synthesis (Christensen et al., 2013) of the then-available regional projections of future TC
 - climatology for 2081–2100 in relation to 2000–2019, for a business as usual emissions scenario
- 59 similar to A1B, shows that worldwide the consensus projection was for decreases in TC numbers by
- approximately 5–30%, increased frequency of Category 4 and 5 storms between 0 and 25%, an
- 61 increase by a small percentage in the typical maximum intensity of life, and an increase in TC
- rainfall amounts by 5–20%. Nevertheless, it is clear that there is great uncertainty about these
- 63 projections. Such projections do not consider changes in off-season TC development in any of the
- 64 basins where TCs form
- 65 Several climatic reconstructions have been performed (Bradley et al., 2006; Mann et al., 2009)
- using proxy data by collecting sediments from the impact of hurricanes in the period 500-1850 and
- 67 then calculated estimates from the statistical model of the activity of tropical cyclones based on
- 68 modern instrumental weather indexes for the period (1851–2006). In analyzing these results and
- 69 comparing them with the cyclone seasons fixed by the World Meteorological Organization, the
- 70 hurricane season (tropical depressions, tropical storms and hurricanes) in the Atlantic Ocean was
- 71 fixed as June 1 to November 30 in 1960, yet we observe a significant variability in off-season TC
- occurrence before/after the hurricane season after the 1960s.
- 73 Even though it not uncommon for TCs to develop outside of their peak TC season months, there is a
- 74 need to examine trends in the number of storms that are forming during low activity months. The
- 75 formation of the extratropical storm Andrea on May 20, 2019 marks the decade of 2010 as that with
- the greatest number of tropical cyclones in the Atlantic Ocean before or after the hurricane season
- dates established by official bodies like the World Meteorological Organization (WMO) and the
 National Oceanic and Atmospheric Administration (NOAA). The frequency of TCs in the North
- 79 Atlantic basin has been found to be influenced by fluctuations in teleconnections such as ENSO and
- 80 AMO (Trenberth et al., 2006; Nogueira et al., 2013). However, human-induced climate change
- 81 manifested as higher sea surface temperatures (SST) and increasing evaporation rates in the tropical
- and sub-tropical North Atlantic basin could also be related to the higher frequency of off-season
- 83 tropical or extratropical cyclone occurrences in more recent decades. That increasing trend in SSTs
- 84 in the Atlantic and other ocean basins and its relation to out off-season TC occurrences during the
- 85 last century has not been thoroughly examined by the scientific community.
- 86 This study aims to determine if off-season TCs have increased in their frequency since the 1900 and
- 87 if that increment in the number pre and post off-season storms could be associated with normal
- 88 climatic variability or climate change. The total number of out off-season TCs per decade for the
- 89 North Atlantic (NA), West Pacific (WP), East Pacific (EP) and South Pacific (SP) ocean basins
- 90 where analyzed to determine if any of the basins experienced an increase in the number of off-
- 91 season tropical/extratropical cyclones over time that could be associated to climatic variability or
- 92 climate change. The Indian Ocean basins were not included in this analysis due to limited data
- 93 availability. Ocean basins that were found to have statistically significant trends were then analyzed
- 94 further with multiple liner regression models (MRL) and regression stepwise procedures to
- 95 determine if climate variability or change factors could explain increasing trends in off-season TC
- 96 frequency over time.

2. Data

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99 Six-hourly TC track data for all storms across all ocean basins were obtained from the International 100 Best Track Archive for Climate Stewardship (IBTrACS) (Knapp et al., 2018) and all TCs that 101 occurred at or after 1900 were extracted. The TC tracks were then then extracted for the northern 102 hemisphere basins that include the East Pacific (EP), the North Atlantic (NA) and the West Pacific 103 (WP) and for the southern hemisphere basin in the South Pacific (SP) (Fig 1a). The off-season TCs 104 were then aggregated by decades in order to identify decadal variability or trends in total storm 105 counts at the individual basin scales. Off-season TCs were defined as storms that occurred in the 106 three months before and after the six-month period of peak cyclone activity in the basin.

107 The monthly frequency of TCs for each basin were analyzed for the entire period and based on that 108 analysis we determined that off-season TCs that occurred during the three months (Mar-Apr-May) before June 1st were pre-season and the three months (Dec-Jan-Feb) after November 30th were post-109 110 season in the northern hemisphere basins (NA, EP and WP) In the southern hemisphere, the three 111 months before (Aug-Sep-Oct) November 1st were classified as pre-season and the three months 112 (May-Jun-Jul) after April 30th were classified as post off-season in the southern hemisphere basins 113 (SP) (Fig 1b). Pre-season and post-season decadal time-series for the Northern/Southern hemisphere and individual basins were then constructed to calculate the total number of TCs per-decade from 114 115 1900 to 2019.

116 The climate variability indexes of ENSO (Niño 3.4), AMO (Trenberth et al., 2019) were 117 respectively obtained from the National Oceanic and Atmospheric Administration (NOAA) Physical Sciences Lab and the National Center for Atmospheric Research. The IPO index was 118 119 obtained from the NOAA Physical Sciences Laboratory (Henley et al., 2015). The variables 120 associated with anthropogenic climate change used in this study were sea surface temperature (SST), global mean surface temperature (GMST) and cloud cover (CC). SST data were obtained 121 122 from the HadISST1 1° reconstruction, GMST data were accessed from the GISTEMP v4 and CC 123 data was acquired from the International Comprehensive Ocean-Atmosphere Data Set (ICOADS) v2.5, all for the 1900-2019 period. The It is important to note that the ICOADS data set has some 124 125 key limitations, like data coverage been sparse and limited corrections that account for changes in 126 observing practices CC estimates from ICOADS are obtained from voluntary observing ships that 127 report CC in octas (eighths) ranging from 0 (completely clear sky) to 8 (completely overcast). These 128 CC estimates are known to be temporally and spatially heterogenous with relatively high 129 observational errors in some areas, yet increases in ocean coverage for clouds after the 1900 have 130 been noted in the latest ICOADS 3.0 release [Eastman et al. 2011, Freeman et al. 2016; 131 Aleksandrova et al. 2018]. A decadal average was calculated for all of the climate variability and

Table 1. Tropical cyclone, climate change and variability variables used in this study.

change variables in order to use them as predictors of decadal TC total counts (Table 1).

	Abbreviations	Units
Tropical Cyclone Counts	TCs	Decadal Total Counts
Climate Change Variables		
Sea Surface Temperature Global Mean Surface Temperature Cloud Cover	SST GMST CC	。 。 Oktas
Climate Variability Variables		
El Niño Southern Oscillation	ENSO 3.4	°SST anomalies index

	Interdecadal Pacific Oscillation	IPO	°SST anomalies index
422	Atlantic Multi-decadal Oscillation	AMO	°SST anomalies index
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135	3. TC Adjustment Method		
136 137 138 139 140 141 142 143 144 145 146	lot of storms that didn't make landfall very between the earlier and later decades, the Another issue related to pre-satellite era depressions, since the detection and classification was probably more difficult percentage of TCs were calculated for the era and new TC monitoring technologies made landfall in both periods. The percentage of TCs period since a higher number of	veren't recorde ne series for each a TC track data ssification of the lt before 1966 he periods 190 es available) in nentage of landf of storms that re	ch basin need to be adjusted accordingly. is the undercount of weaker tropical
148 149 150 151 152 153 154 155 156 157	in the pre-satellite period is increased us satellite era. The total number of additional two periods to be the same or near equal and then the number of extra storms for storms for each basin and that resulting between 1900 and 1969. In a previous storms in the North Atlantic to determine	ntil its landfall ponal TCs that re al are then divice each decade is number is ther tudy (Landsea, he if the basin h	e 1900-65 period, the number of total storms percentage is equal to the one in the post esulted in the landfall percentages between the ded by the 7 decades of the pre-satellite eras multiplied by the percentage of off-season added then to each of the individual decades, 2007), this method was applied to adjust TC has experienced an increasing trend in annual after adjusting the tropical storm counts no
158 159 160 161 162	basin for the 1900-2019 period . First, w	ve calculate the Ifalling TCs (L	was applied to the total TC count for the NA landfall percentage for the pre-satellite period FTCs) with the total number of storms (TTCs) eck the equations below:
		(LFTCs /TTCs	(1)
163	Examp	le: (479/610)	* 100 = 78.5%
164 165 166	Then calculate landfall % for the period	post-satellite	period 1966-2019,
100		(LFTCs /TTCs) * 100
	Examp	,	*100 = 69.1% (2)
167			
168	The and finially increase the number of	TC. (192.5	the NA hearn) motil the landfall 0/ af the 1000

Then artificially increase the number of TCs (+83 for the NA basin) until the landfall % of the 1900-

LFTCs / (TTCs + AddTCs) * 100

Example: 479/(610 + 83) * 100 = 69.1%

(3)

65 period is equal to landfall % of the 1966-2019:

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Then calculate the percentage (OffP) of off-season TCs (OffTCs) by dividing it by total number of

$$(OffTCs/TTCs) * 100$$

Example: $(67/1454) * 100 = 4.61\%$

Then divide additional TCs (83) by the number of decades between 1900 and 1969 (7) and then multiply by the off-season TC percentage (.0461)

$$(AddTCs/Decades) * OffP$$

 $Example: (83/7) * .0461 = 0.54$ (5)

 In the case of the NA, we determined that by using the above TC series adjustment method the basin would get an additional 0.54 off-season TCs for each of the seven decades that go from the 1900 to 1969. Finally, the additional 0.54 TCs per decade will be divided between pre and post off-season TCs by multiplying the added storms with the respective percentage of pre/post off season cyclones:

$$DecOffTCs*Percentage/Post Season \\ Example: 0.54/0.62 = 0.33 \ and \ 0.54/0.38 = 0.21 \\ \eqno(6)$$

The pre off-season decades of the NA basin before 1970 will get an additional 0.33 TCs and the post off-season decades will get 0.21 more storms. This off-season TC adjustment method was applied to the other five basins. It is important to note that this TC adjustment method has been only implemented in the NA basin and that this study is the first attempt to apply technique in other ocean basins. This adjustment method is in no way capable of detecting all TCs that formed before the satellite era, yet it offers us the opportunity to estimate missed storms by comparing the TC landfall percentage of the pre- and post-satellite eras.

4. Statistical Methods & Models

Mann-Kendall (MK) tests for trends (Mann, 1945; McLeod, 2005) were applied to all the off-season TC decadal series for all basins in order to determine if the frequency of storms has increased or decreased over time. This test has the advantage of not assuming any special form for the distribution function of the data, while having a power nearly as high as their parametric equivalents and that is why its use is highly recommended by the World Meteorological Organization (Hipel and McLeod, 2005).

The decadal series that were then found to have a significant trend based on the MK results were then furtherly analyzed by applying a series of multiple linear regression models (MLR). MLR were used to model the decadal count of off-season TCs for basins that showed increasing or decreasing trends in storm numbers to test if covariates associated with climatic variability and climate change explained off-season TC frequency. MLR attempts to model the relationship between two or more explanatory variables and a response variable by fitting a linear equation to observed data.

Three different MLR models were run for each off-season TC series that exhibited a statistically significant trend, one MLR model with the climate change variables (SST, GMST & CC) as predictors, another model with just the climate variability factors (ENSO, AMO & IOD) and a final model with all of the variables included. Then the three MLR models (pre-season, post-season and off-season) were run for each of the basins with increasing trends in off-season TCs, the best models

(highest adjusted R-squared and lowest p-value) were then selected for each of the series. The MLR models were run in The R Project for Statistical Computing using the biglm package.

Finally, stepwise selection MLR models were used to identify the climate variability or change factors making the most statistically significant contributions to off-season increasing TC frequency. Here we use stepwise selection which is a combination of the forward and backward procedures where you start with no predictors, then sequentially add the most contributive predictors. After adding each new variable, it removes the variables that no longer provide an improvement in the model fit (James et al., 2014; Bruce and Bruce, 2017). The MLR and stepwise for the off-season TC count series for each of the basins with significant increasing trends were run in The R Project for Statistical Computing using the MASS package (Venables and Ripley, 2002).

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5. Results & Discussion

When analyzing the number of TCs for all basins for the 1900-2019 period we found that 713 off-season storms occurred during that time, most during the months of May (NH pre-season and SH post-season) with 430 and December (NH post-season) with 341 (Figure 1a, 1b). When looking at the count of off-season TCs per basin we found that as expected the West Pacific (611) and South Pacific (85) accounted for 81.3% of all off-season storm occurrences. When grouping the basins between northern and southern hemispheres, we find that 89% of all off-season TCs occurred north of the equator for the 1900-2019 period (Figure 1a, 1b). The North Atlantic and East Pacific basins were found to be the ones with the lowest numbers of off-season TCs when compared to the other two Pacific basins.

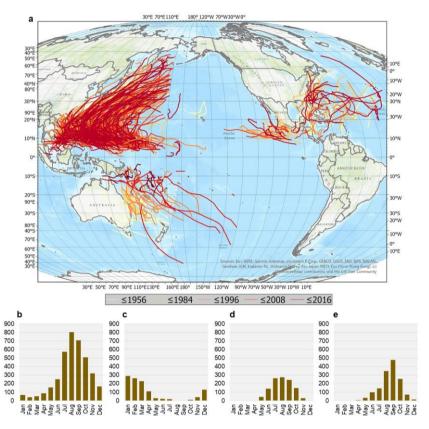
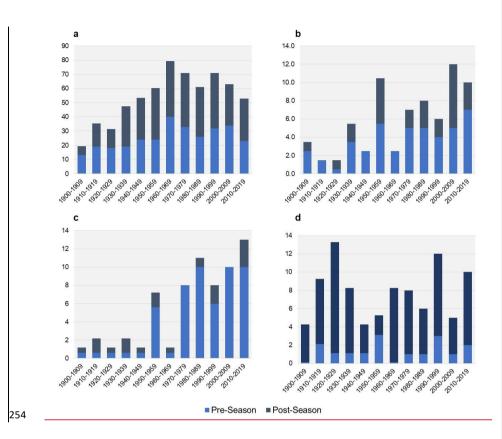


Figure 1. Tracks of all out off-season TCs (a) and the number of storms per month for the WP (b), SP (c), EP (d) and NA (e) basins for the 1900-2019 period.

As shown in Figure 2, even after adding the estimated missing storms before the 1970 decade, most basins experienced their highest number of out off-season TCs (pre or post) in decades at or after 1960-69. The 1960-69 decade for the northern hemisphere basins (WP, NA and EP) was found to be the one with the highest number of pre off-season TCs with 69 and the 1950-1959 decade was identified as the one with the most post off-season storms with 68 (Figure 2a, 2b, 2c and 2d). When examining TC counts for all basins individually, we found that the NA and EP basins had their most active decades after 1970 and that the WP and SP basins experienced their highest storm count decade after 1960 (Figure 2c, 2d). It is important to note that these results already reflect the additional TCs that were added to the pre-satellite era.



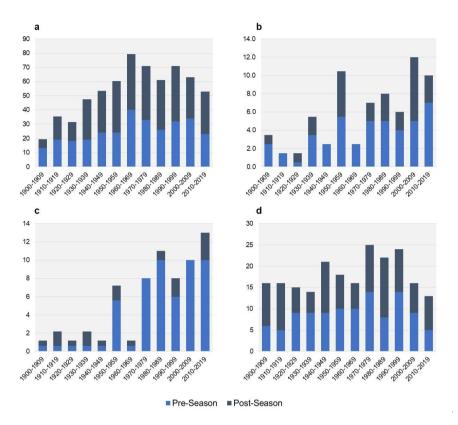


Figure 2. Adjusted decadal count of all observed and estimated off-season TCs for the WP (a), NA (b), EP (c) and SP (d) ocean basins for the 1900-2019 period.

The Mann-Kendall non-parametric tests for trends for all basins show that three basins exhibited statistically significant increasing trends in adjusted decadal off-season TC counts for the 1900-2019 period (Table 2). The basins with statistically significant increasing trends for the entire time period were the EP (pre and off-season), NA (pre, post and off-season) and the WP (pre, post and off-season). The EP basin shows an increasing trend in pre- and off-season TCs that is more evident from the 1950s to the present (Figure 2d), while the increasing trend in the NA basin can be observed from the 1970s to 2019. The NA and EP basins also exhibit positive trends when the analysis is done only considering the post satellite era, yet the results are not statistically significant (Table 2)

The increasing trend in off-season TCs in the WP basin is more evident from the 1900 to 1969 (Figure 2a), which was during the pre-satellite era where missing TCs were added to the series. However, no trenda negative trend is found in the WP basin if the decadal counts are analyzed from the 19670s to the present, yet those results were not statistically significant. In the post-satellite era in the WP basin, the 1990-1999 decade was identified as the one with most off-season TCs, however the two following decades exhibited a decreasing trend. The EP and NA basins show significant increasing trends in off-season TC counts (Table 2). Opposites to the trends identified in the WP basin, the EP and NA also show increasing decadal counts after the 1960s and 1970s__decades. The SP basin also exhibited a positive Tau coefficient, yet it was statistically insignificant for the entire period and the post-satellite era (Figure 2d).

Table 2. Results of Mann-Kendall trend tests for the 1900-2019 period for all ocean basins where TCs form.

Trends for the 1900-2019 period								
Pre TCs	Tau S	P-value	Post TCs	Tau S	P-value	Off TCs	Tau S	P-value
EP	0.746	0.002	EP	0.098	0.723	EP	0.679	0.004
NA	0.572	0.015	NA	0.485	0.042	NA	0.554	0.016
SP	0.048	0.889	SP	0.015	1.000	SP	0.061	0.836
WP	0.554	0.016	WP	0.485	0.034	WP	0.534	0.019
			Trends for the	he 1960-2	019 period			4
<u>EP</u>	0.596	<u>0.158</u>	<u>EP</u>	0.414	0.338	<u>EP</u>	<u>0.69</u>	0.085
<u>NA</u>	0.596	<u>0.158</u>	<u>NA</u>	0.645	<u>0.119</u>	<u>NA</u>	0.6	0.132
<u>SP</u>	0.596	<u>0.158</u>	<u>SP</u>	<u>-0.2</u>	0.707	<u>SP</u>	<u>-0.06</u>	<u>1</u> \
<u>WP</u>	-0.467	0.259	<u>WP</u>	-0.6	0.132	WP	-0.69	0.085

Significant trends in bold.

MLR models were run on the basins that exhibited statistically significant (< 0.05) increasing trends in decadal total off-season TC counts over time and here we report the best models for each of the series. The MLR results show that the statistically significant increasing trends in TC frequency for the EP (pre and off-season) and WP basins is best explained by climate change factors SST, GMST and CC in both the normal and detrended series at the 0.05 significance level (Table 3). Climate change factors accounted for 56% (pre-season) and 52% (off-season) of the increasing trend in TC counts for the EP basin. In the WP basin climate change factors explained 55% (pre-season), 64% (post-season) and 68% (off-season) of the trends in off-season TCs, yet lower R squares were found when the analysis was done with the detrended series. Increasing trends in SSTs, GMST and moisture (CC) outside of the prime months of tropical storm development could promote more optimumbetter conditions for higher off-season TC occurrences (Klozback, 2006; Hansen et al., 2010). The MLR models were also done with detrended decadal series of off season TCs, and detrended decadal series of the different climate variability and change factors, and similar results were found.

Table 3. Best multiple linear regression models (MLR) for basins with statistically significant increasing trends in off-season TCs with detrended climate indices

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significant increasing trends in on-season ros with detrended climate indices.					<u>. </u>
<u>Model</u>	Adj. R ²	Adj. R ² Det.	<u>Factors</u>	<u>p-val.</u>	p-val Det.
EP pre-season	0.563	0.444	SST, GMST & CC	0.021	0.038
EP off-season	0.522	0.472	SST, GMST & CC	0.030	0.024
NA pre-season	0.481	0.496	SST, GMST & CC	0.041	0.022
NA post-season	0.427	0.247	ENSO & AMO	0.130	<u>0.057</u>
NA off-season	0.384	0.406	SST, GMST & CC	0.070	<u>0.010</u>
WP pre-season	<u>0.551</u>	0.462	SST, GMST & CC	0.020	0.000
WP post-season	0.645	0.478	SST, GMST & CC	0.000	0.023
,WP off-season	0.689	0.481	SST, GMST & CC	0.005	0.017

Table 3. Best multiple linear regression models (MLR) for basins with statistically significant increasing trends in off-season TCs.

Model	Multiple R-squared	Adjusted R ²	Factors	p-value
EP pre-season	0.682	0.563	SST, GMST & CC	0.021

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EP off-season	0.653	0.522 SST, GMST & CC 0.030
NA pre-season	0.622	0.481 SST, GMST & CC 0.041
NA post-season	0.687	0.427 ENSO & AMO 0.130
NA off-season	0.552	0.384 SST, GMST & CC 0.070
WP pre-season	0.673	0.551 SST, GMST & CC 0.020
WP post-season	0.742	0.645 SST, GMST & CC 0.000
WP off-season	0.774	0.689 SST, GMST & CC 0.005

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The climate variability factors (ENSO, AMO & IOD) did not exhibit statistically significant relationships with increasing off-season TC counts, which shows that natural variability does not explain the incrementing number of storms in the EP and WP basins. MLR model results for the NA basin also showed the climate change variables accounting for 48% (pre-season) and 38% (offseason) of the increasing trend in TCs (Table 3). However, the MLR model results for the postseason months in the NA basin showed that the climate variability variables (ENSO & AMO) accounted for 42% (25% in the detrended) of the increasing trend in TCs, yet the model was not found to be statistically significant in both the normal and detrended series. It is well known that cold phases of ENSO (La Niña) and warm phases of AMO tend to be associated with higher TC frequency in the North Atlantic ocean (Tang and Neelin, 2004; Briggs, 2008) and this could explain why those teleconnections were found to have the most significant influence on post-season TC frequency in the NA basin. When the MLR results of the original and detrended series are compared (Table 3), we find that the models with the detrended series exhibit lower R squares than the MLR models with the original series, yet those models were still found to be statistically significant which suggests that the correlation between off-season TCs and climate change factors is strong even after decadal trends are removed. . Yet, it is important to note that in most basin series, including the NA, climate change variables explained more of the off season TC increasing trend than the climate variability factors.

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Stepwise MLR model results showed that climate change factors (SST, GMST & CC) were among the selected variables that explained most of the increasing trend in off-season TCs for all basins analyzed (Table 4). In the EP basin, SST, ENSO, and CC, accounted for 69% (pre-season) and 65% (off-season) of the increasing trend in TCs. In the NA basin, the stepwise procedure selected CC as the sole climate change factor that explained 52% (pre-season) and 40% (off-season) of the rising frequency in TC counts. However, CC & AMO were selected as the variables that explained (43%) most of the variability in TC frequency during the post-season months in the NA basin. Stepwise procedure results for the WP basin show that climate change and variability factors were selected as the best predictors of TC frequency, with GMST and CC accounting for 57% (pre-season), CC, GMST, ENSO and IPO explaining 72% (post-season) and 74% (off-season) of the variability of TCs.

Table 4. Stepwise multiple linear regression models (MLR) for basins with statistically significant increasing trends in off-season TCs.

Model	R-squared	Adjusted R ²	Factors	p-value
EP pre-season	0.777	0.694	SST, ENSO & CC	0.005
EP off-season	0.747	0.652	SST, ENSO & CC	0.008
NA pre-season	0.569	0.526	CC	0.004
NA post-season	0.687	0.427	CC & AMO	0.098
NA off-season	0.460	0.406	CC	0.015

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WP pre-season	0.655	0.578	GMST & CC	0.008
WP post-season	0.826	0.726	CC, GMST, ENSO & IPO	0.008
WP off-season	0.839	0.747	CC, GMST, ENSO & IPO	0.006

The EP experienced a steady increase in off-season TC total counts from 1900 to 2019 at a rate of 1.1 additional storms per decade. The decadal off-season total TC count series for the EP basin closely resembles the increasing trend in average SSTs and CCs (Fig 3a, 3c). When the EP off-season TC tracks are examined, it shows that most storms have formed in areas that have experienced statistically significant increasing trends in SST and CC, yet no changes in track or genesis location were detected over time. (Fig 3d, 3e). The correlation between off-season TCs in the EP basin and ENSO is not as clear as the one between SST and CC, with some mostly warm ENSO decades like the 1990-1999 exhibiting lower storm counts and other periods with cooler phases dominating showing a higher number of cyclones. When SST patterns for areas in the EP basin where TCs develop are examined over time, we find that most tropical/sub-tropical ocean waters have experienced a statistically significant increasing trend in ocean surface temperatures from 1900 to 2019 (Fig 3d). Similar to other studies (Hansen et al., 2010), we find that the EP tropical ocean surfaces have increased by 0.051 degrees C° per decade. When CC patterns are examined, we find that it has also experienced a statistically significant increasing trend in some areas in the EP basin (Fig 3e).

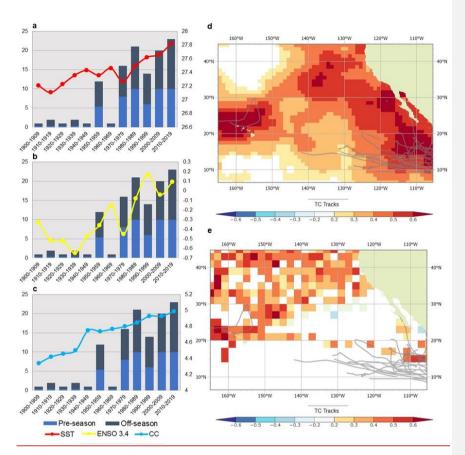


Figure 3. Decadal TC counts for the EP off-seasons and decadal average SSTs (a), decadal TC counts for the EP off-seasons and decadal average ENSO 3.4 (b), decadal TC counts for the EP off-seasons and decadal average Correlation between Time and Dec-May averaged CC (c), correlation between Time and Dec-May averaged SST (C°) and the off-season TC tracks for the 1900-2019 period (d) and correlation between Time and Dec-May averaged CC (oktas) and the off-season TC tracks for the 1900-2019 period.

The decadal off-season total TC count series for the NA basin closely resembles the increasing trend in average SSTs and CCs (Fig 4a, 4c). The NA decadal series shows a steady increase in off-season TC total counts from 1900 to 2019 at a rate of 0.7 additional storms per decade and an SST increasing trend of 0.055 C° per decade. Both the average decadal SST and CC series coincide with the peaks and valleys in off-season TC counts for the NA basin, with the 1950-1959 showing a high number of storms associated with high mean SSTs and CCs while the drop in storm counts in the 1960-1969 decade matches a drastic drop in ocean surface temperatures (Figures 4a, 4c). The off-season TC tracks in the NA basin also formed in areas that exhibited increasing trends in SST and CC, however no changes in track or genesis location were detected over time -(Figure 4d, 4e). Even though average SSTs increase to 0.135 C° per decade from 1970 to 2019, off-season TC total counts went down in the 1990-1999 and 2010-2019 decades, with the decade in between (2000-2009) exhibiting the highest number of off-season TCs (14) of all decades examined. However, it is important to note that 5 out of the 6 decades with the most off-season TCs in the NA basin occurred after the 1970s.

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When North Atlantic SSTs are examined in areas where TCs form, we found that ocean surface temperatures have increased at a rate of 0.055 degrees C° per decade for the off-season months of Dec-March (Fig 4d). When CC patterns are examined, we find that it has also experienced a statistically significant increasing trend of 0.06 oktas (eighths of the sky that are covered in clouds) per decade in the North Atlantic basin since the 1900 (Figure 4e). If the NA pre/post off-season series is modified to begin in the 1960s, we find that SSTs have increased at a decadal rate of 0.082 C° per decade at a rate of 1.2 additional storms per decade. Overall, these results suggest that increasing trends in SSTs, which also drive increasing trends in evaporation rates associated with high CCs, are the physical mechanisms behind most of the recent increase in the total number of out of season TCs in the NA basin. The correlation between off-season TCs in the NA basin and AMO is not as clear as the one between SST and CC, with some warm AMO phases between 1930-1959 exhibiting lower storm counts while some cooler phases (1970-89) showing a higher number of cyclones.

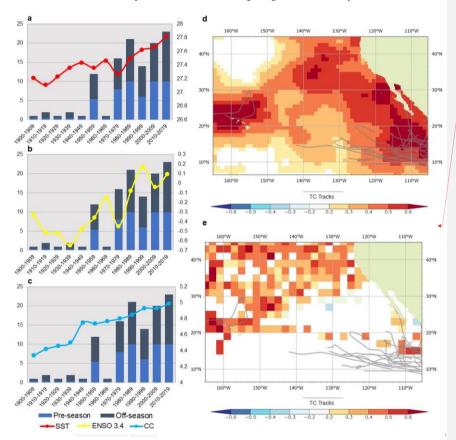


Figure 3. Decadal TC counts for the EP off-seasons and decadal average SSTs (a), decadal TC counts for the EP off-seasons and decadal average ENSO 3.4 (b), decadal TC counts for the EP off-seasons and decadal average Correlation between Time and Dec May averaged CC (c), correlation between Time and Dec May averaged SST (C°) and the off-season TC tracks for the 1900-2019 period (d) and correlation between Time and Dec May averaged CC (oktas) and the off-season TC tracks for the 1900-2019 period

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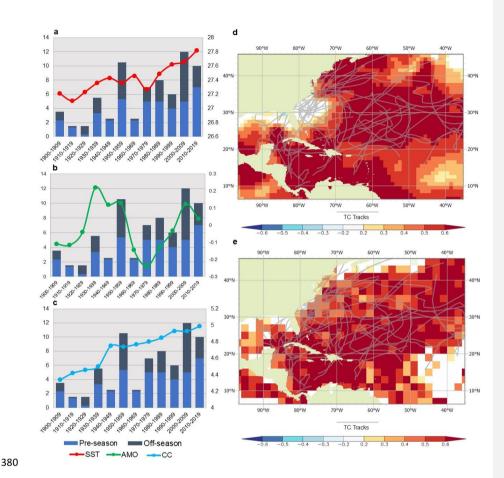


Figure 4. Decadal TC counts for the NA off-seasons and decadal average SSTs (a), decadal TC counts for the NA off-seasons and decadal average AMO (b), decadal TC counts for the NA off-seasons and decadal average Correlation between Time and Dec-May averaged CC (c), correlation between Time and Dec-May averaged SST (C°) and the off-season TC tracks for the 1900-2019 period (d) and correlation between Time and Dec-May averaged CC (oktas) and the off-season TC tracks for the 1900-2019 period (e).

The decadal off-season total TC count series for the WP basin closely resembles the increasing trend in GMST (Fig 5a). However, the WP basin experienced the highest count of off-season TCs in the 1960-69 decade, not in the more recent decades like the EP and NA basins. More importantly, if trend analysis for off-season TC counts is done from 1960-2019 in the WP basin, we find find no statistically significant increasing a none statistically significant or decreasing trend. However, it is important to note that four out of the five decades with most off-season TCs in the WP basin occurred after 1960. However, the 2010-2019 decade was identified as the period with the lowest total number of off-season TCs even though increasing trends in mean SST, GMST and CC continued (Fig 5a, 5d and 5e). Off-season TC tracks in the WP basin also correlate spatially with areas that show increasing trends in SST and CC (Fig 5d, 5e), yet no changes in TC genesis or tracks were identified over time. The decreasing number of off-season TCs in the last two decades coincided with a negative phase of the IPO, which suggests that TC frequency in the WP basin is influenced by fluctuations in the IPO (Fig 5c), whose recent negative phase since 1998 resembles La Niña-like SST anomaly patterns

(Zhao et.al, 2018). Even though most of the variability in off-season TC frequency in the WP basin can be explained by climate change trends in GMST, SST and CC, the rest of the variance in TCs is account by fluctuations in the IPO and ENSO teleconnections.

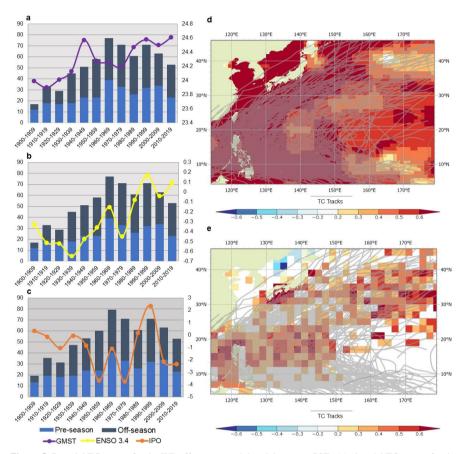


Figure 5. Decadal TC counts for the WP off-seasons and decadal average SSTs (a), decadal TC counts for the WP off-seasons and decadal average AMO (b), decadal TC counts for the WP off-seasons and decadal average Correlation between Time and Dec-May averaged CC (c), correlation between Time and Dec-May averaged SST (C°) and the off-season TC tracks for the 1900-2019 period (d) and correlation between Time and Dec-May averaged CC (oktas) and the off-season TC tracks for the 1900-2019 period (e).

The results of the MLR and stepwise procedures have shown that decadal off season TC increasing trends in the NA, EP and WP basins is mostly associated with climate change (SST and CC) and climate variability factors (ENSO and IPO). Since no previous studies have focused on analyzing trends in off season TC trends, our results can only be compared to analyses that have consider in season storms. Studies that have examined TC frequency overall have found increases in the number of most intense hurricanes [¿Wang and Lee, 2008; Knutson et al., 2010; Emanuel, 2013]; yet no clear trend has been found when lower intensity TCs have been examined [Landsea, 2007]. The results of other studies show that there is no overall agreement on the relationship between SSTs and TC frequency (Landsea, 2005; Emanuel, 2005; Trenberth and Shea, 2006; Trenberth, 2007), yet some have found strong associations between TC variability and ENSO, AMO and IPO [Camargo

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and Sobel, 2005; Nogueira and Kim, 2007; Mahala et al., 2015; Zhao et al., 2018]. In this study we analyzed off-season TCs and our results differ from those that have found no trend in overall TC frequency, since we found decadal increasing trends in the NA and EP basin in both the pre and post-satellite eras. The results presented here suggest that climate change trends like increasing SSTs and more favorable moisture environments (CC) between the months of Dec to May in the NA and EP basins seem to be the major factors behind decadal increasing trends in off-season TCs.

The results of the MLR and stepwise procedures have shown that decadal off-season TC increasing trends in the NA, EP and WP basins is mostly associated with climate change (SST and CC) and climate variability factors (ENSO and IPO). Since no previous studies have focused on analyzing trends in off-season TC trends, our results can only be compared to analyses that have consider inseason storms. Studies that have examined TC frequency overall have found increases in the number of most intense hurricanes (Wang and Lee, 2008; Knutson et al., 2010; Emanuel, 2013).

6. Summary and concluding remarks

The frequency of TCs that developed outside of their prime season months were analyzed to determine if trends in higher storm totals in the Pacific and Atlantic Ocean basins were associated with natural variability, climate change or both. Adjusted off-season decadal TC total counts for six ocean basins were analyzed for the 1900-2019 period in order to determine if the number of storms have been increasing over time. Mann-Kendall tests for trends were done and the basins that exhibited statistically significant increasing trends were then furtherly analyzed using multiple linear regression models and stepwise procedures to determine if those trends could be explained by fluctuations associated with climate variability, climate change trends or a combination of both.

The main results of this study suggest that decadal total off-season (pre/post) TC counts have significantly increase over time since the 1900 in the East Pacific (EP), North Atlantic (NA) and West Pacific (WP) basins. The EP and NA basins exhibited statistically significant increasing trends even if the analysis was done from the 1960s instead of the 1900. The WP basin showed an overall increasing trend in the total number of off-season TCs per decade, yet if the analysis is done from the 1960s to the present, no statistically significant increasing trend is found. However, the three basins that reflected an overall increase in decadal off-season TC frequency had their most active decades after the 1970s.

Results from the best multiple linear regression (MLR) models show that the increasing decadal count of off-season TCs has been found to be strongly associated with climate change trends in sea surface temperature (SST), global mean surface temperature (GMST) and cloud cover (CC) in all three basins (EP, NA and WP). The MLR model where climate variability variables (ENSO and AMO) explained most of the variance in off-season TC counts was in the storm decadal counts for the post-season months of the NA basin.

Results of the MLR stepwise procedures showed that the selected variables that accounted for most of the variability in off-season TCs for the EP basin were SST, CC and ENSO, while CC (preseason and off-season) and AMO (post-season) were chosen as the best variables for the NA basin. The stepwise procedure identified the climate change trends in GMST and CC, and fluctuations in ENSO and IPO as the variables that accounted for most of the variability in decadal off-season total TC counts in the WP basin,

The findings of this study suggest that trends in SST, GMST and CC associated with climate change are not only altering the frequency (Klotzbachk, 2006; Saunders and Lea, 2006; Hansen et al., 2010) and intensity of TCs that develop during the peak months of the season, they are also altering the total number of storms that form in the off-season months (Dec-May), especially in the EP and NA basins. The results of this study have important implications for the NA and EP basins, if off-season TCs have been increasing in frequency since the 1900 we can expect that this trend associated

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with climate change would continue in future decades. This increasing number of off-season TCs could potentially impact societies in their path during times of the year when storms are least expected, possibly increasing environmental and economic impacts in areas that are already experiencing the effects of climate change exacerbated phenomena.

One of the main limitations of this work was the inclusion of tropical depressions in the off-season TC analysis. If data on TC intensity were widely available for all off-season TCs, it would have been possible to exclude weaker tropical depressions from the analysis since the detection and classification of those storms was more difficult in the pre-satellite era. Other limitations of this study include the issues of worst data quality in the pre-satellite era, the problem of applying a universal missed TC adjustment method to all basins analyzed and the lack of information on TC intensity for many storms, especially in the pre-satellite era.

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508 Atmospheric Administration Physical Sciences Lab (PSL; https://psl.noaa.gov/data/timeseries/IPOTPI/).

References

Aleksandrova, M., Gulev, S. K., and Belyaev, K.: Probability distribution for the visually observed
 fractional cloud cover over the ocean. *Journal of Climate*, *31*(8), 3207-3232.
 https://doi.org/10.1175/JCLI-D-17-0317.1, 2018.

National Center for Atmospheric Research (NCAR; https://climatedataguide.ucar.edu/climate-

data/atlantic-multi-decadal-oscillation-amo). The Interdecadal Pacific Oscillation data supporting this article are based on publicly available measurements from the National Oceanic and

- Bradley, R.S., M.K.-Hughes M.K., and M.E.-Mann, M.E.: (2006) Authors were clear about hockeystick uncertainties₂- Nature, 442 (7103): 627. https://doi.org/10.1038/442627b, 2006.
- Bruce, P.Peter, and Andrew Bruce, A.: 2017. Practical Statistics for Data Scientists. O'Reilly Media, 2017.
- Camargo, S. J., and & Sobel, A. H. (2005): —Western North Pacific tropical cyclone intensity and
 ENSO. *Journal of Climate*, 18(15), 2996-3006. https://doi.org/10.1175/JCLI3457.1, 2005.

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```
521
            V., Bex and P.M. Midgley (eds.) Christensen J.H. et al. in Climate Change.: (2013) The
            Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the
522
523
            Intergovernmental Panel on Climate Change, (eds). Stocker T.F., Qin D., Plattner G.-K., Tignor
524
            M., Allen S.K., Boschung J., Nauels A., Xia Y., V.B. & P.M.M.) Cambridge; Cambridge
525
            University Press, 2013.
526
        Emanuel, K.A. (200: 5)-Increasing destructiveness of tropical cyclones over the past 30 years.
            Nature, 436: 686-688. https://doi.org/10.1038/nature03906, -2005.
527
        Emanuel, K.A.: (2013)-Downscaling CMIP5 climate models shows increased tropical cyclone
528
529
            activity over the 21st century. PNAS July 23, 2013 110 (30) 12219-12224. Accessed from:
530
            https://doi.org/10.1073/pnas.1301293110, 2013.
531
        Freeman, E., Woodruff, S. D., Worley, S. J., Lubker, S. J., Kent, E. C., Angel, W. E., ... & Gloeden,
            W.: ICOADS Release 3.0: a major update to the historical marine climate record. International
532
533
            Journal of Climatology, 37(5), 2211-2232, 2017.
534
        Goldenberg, S.B., C.W. Landsea Landseer, C.W., A.M. Mestas-Nuñez, A.M., -and W.M. Gray,
535
        W.M.: (2001) The recent increase in Atlantic hurricane activity: Causes and implications. Science,
        293: 474-478. DOI: 10.1126/science.1060040, 2001.
536
        Hansen, J., Ruedy, R., Sato, M., & Lo, K.: (2010). Global surface temperature change. Reviews of
537
            Geophysics, 48(4). https://doi.org/10.1029/2010RG000345, 2010.
538
539
        Henley, B.J., Gergis, J., Karoly, D.J., Power, S.B., Kennedy, J., & Folland, C.K.: (2015). A Tripole
           Index for the Interdecadal Pacific Oscillation. Climate Dynamics, 45(11-12), 3077-
540
            3090.http://dx.doi.org/10.1007/s00382-015-2525-1 . Accessed on MM DD YYYY at
541
542
            "/data/timeseries/IPOTPI", 2015 .-
        Hipel, K.W. and McLeod, A.I.: (2005). Time Series Modelling of Water Resources and
543
           Environmental Systems. Electronic reprint of our book originally published in 1994.
544
545
            http://www.stats.uwo.ca/faculty/aim/1994Book/, 2005.
546
        Holland, G.J. and P.J. Webster.: (2007) Heightened tropical cyclone activity in the North Atlantic:
547
            Natural variability or climate trend? Philos. Trans. Roy. Soc., 365: 2695-2716.
            https://doi.org/10.1098/rsta.2007.2083, 2007.
548
549
        James, Gareth, J., Daniela Witten, D., Trevor Hastie, T. -and Robert Tibshirani, R.: 2014. An
        Introduction to Statistical Learning: With Applications in R. Springer Publishing Company,
550
551
        Incorporated, 2014. -
552
        Knapp, K. R., H. J. Diamond, H. J., J. P. Kossin, J. P., M. C. Kruk, M. C. and C. J. Schreck, C. J.:
553
            2018: International Best Track Archive for Climate Stewardship (IBTrACS) Project, Version 4.
            [indicate subset used]. NOAA National Centers for Environmental Information. non-
554
555
            gonvernment domain <a href="https://doi.org/10.25921/82ty-9e16">https://doi.org/10.25921/82ty-9e16</a>, 2018.
556
        Knutson, T.R., J.L. McBride, J.L., J. Chan, J., K. Emanuel, K., G. Holland, G., C. Landsea, C., I.
557
            Held, I., J.P. Kossin, J.P., A.K. Srivastava, A.K., and M. Sugi, M.: (2010) Tropical cyclones
            and climate change. Nature Geoscience, 3: 157-163; doi:10.1038/ngeo779, 2010.
558
559
        Klotzbach, P.J.: (2006) Trends in global tropical cyclone activity over the past twenty years
            (1986-2005) Geophys. Res. Lett., 33: L10805, pp 1-4. https://doi.org/10.1029/2006GL025881
560
```

Stocker, T.F., D. Qin, G.-K., Plattner, M., Tignor, S.K., Allen, J., Boschung, A., Nauels, Y., Xia,

520

561 562

563

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Landsea, C.W.: (2005) Meteorology: Hurricanes and global warming. Nature, 438: E11–E12.

https://doi.org/10.1038/nature04477, 2005.

```
565
           trends in extreme tropical cyclones? Science 2006, 313:452-454. DOI:
566
           10.1126/science.1128448, 2006.
567
       Landsea, C.W.: (2007). Counting Atlantic tropical cyclones back to 1900. Eos, Transactions
568
           American Geophysical Union, 88(18), 197-202. https://doi.org/10.1029/2007E0180001, 2007.
569
       Mahala, B. K., Nayak, B. K., and Mohanty, P. K.: (2015). Impacts of ENSO and IOD on tropical
570
           cyclone activity in the Bay of Bengal. Natural Hazards, 75(2), 1105-1125.
           https://doi.org/10.1007/s11069-014-1360-8, 2015.
571
572
       Mann, H.B.: (1945), Nonparametric tests against trend, Econometrica, 13, 245-259, 1945.
573
       Mann, M.E., T.A. Sabbatelli, T.A., -and U. Neu, U.: (2007) Evidence for a modest undercount bias
           in early historical Atlantic tropical cyclone counts. Geophys. Res. Lett., 34: L22707;
574
575
           doi:10.1029/2007GL031781, 2007.
576
       Mann, M.E., J.D. Woodruff, J.D., J. Donnelly J., and Z. Zhihua, Z.: (2009) Atlantic hurricanes and
577
           climate over the past 1,500 years. Vol. 460 (13 August 2009); doi:10.1038/nature0821, 2009. -
578
       McLeod, A. I.: (2005). Kendall rank correlation and Mann-Kendall trend test. R Package Kendall,
579
580
       Moon, I. J., Kim, S. H., and Chan, J. C.: (2019). Climate change and tropical cyclone
581
           trend. Nature, 570(7759), E3-E5, 2019. -
582
       Nogueira, R. C., Keim, B. D., Brown, D. P., and Robbins, K. D.: (2013). Variability of rainfall
           from tropical cyclones in the eastern USA and its association to the AMO and
583
584
           ENSO. Theoretical and applied climatology, 112(1-2), 273-283.
585
           https://doi.org/10.1007/s00704-012-0722-y, 2013.
586
       Saunders, M., Lea, A.: Large contribution of sea surface warming to recent increase in Atlantic
587
           hurricane activity. Nature 451, 557–560, (2008) doi:10.1038/nature06422, 2008.
588
       Trenberth, K.E. and D.J. Shea, D.J.: (2006) Atlantic hurricanes and natural variability in 2005.
589
           Geophys. Res. Lett., 33: L12704, doi:10.1029/2006GL026894, 2006
590
       Trenberth, K. E.: (2007). Warmer oceans, stronger hurricanes. Scientific American, 297(1), 44-51.
           https://www.jstor.org/stable/26069374, 2007.
591
592
       Trenberth, K.E.evin, Zhang, R., ong. & National Center for Atmospheric Research Staff (Eds). Last
593
           modified 10 Jan 2019. "The Climate Data Guide: Atlantic Multi-decadal Oscillation (AMO)."
```

Retrieved from https://climatedataguide.ucar.edu/climate-data/atlantic-multi-decadal-

Venables, W. N. and Ripley, B. D. (2002) Modern Applied Statistics with S. Fourth edition.

250, https://www.sciencedirect.com/science/article/pii/S2225603220300047, 2019.

Geophysical Research Letters, 35(2). https://doi.org/10.1029/2007GL032396, 2008.

archive. International journal of climatology, 31(7), 951-967, 2011.

Wang, C. and, & Lee, S. K.: (2008). Global warming and United States landfalling hurricanes.

Woodruff, S. D., Worley, S. J., Lubber, S. J., Ji, Z., Eric Freeman, J., Berry, D. I., ... & Wilkinson, C.: ICOADS Release 2.5: extensions and enhancements to the surface marine meteorological

Walsh, K. J., Camargo, S. J., Knutson, T. R., Kossin, J., Lee, T. C., Murakami, H., Particular, C.,

2019: Tropical cyclones and climate change. Tropical Cyclone Research and Review, 8(4), 240-

Landsea, C.W., B.A. Harper, B.A., K. Hoarau, K., and J.A. Knaff, J.A.: (2006) Can we detect

564

594

595

596

597

598

599

600 601

602

603

604 605 oscillation-amo, 2019. -

Springer, 2002. -

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Zhao, J., Zhan, R., Wang, Y., and & Xu, H.: (2018). Contribution of the interdecadal Pacific oscillation to the recent abrupt decrease in tropical cyclone genesis frequency over the western North Pacific since 1998. Journal of Climate, 31(20), 8211-8224. https://doi.org/10.1175/JCLI-D-18-0202.1, 2018.