# Increasing Frequency in Off-Season Tropical Cyclones and its relation to Climate Variability and Change

# 3 José J. Hernández Ayala<sup>1</sup> and Rafael Méndez-Tejeda<sup>2</sup>

- <sup>1</sup> Department of Geography, Environment & Planning, Climate Research Center, Sonoma State
   University, California, USA. jose.hernandezayala@sonoma.edu
  - <sup>2</sup> Research Laboratory in Atmospheric Science, University of Puerto Rico at Carolina, Puerto Rico.
     P. O. Box 4800, 00984, Carolina, Puerto Rico. rafael.mendez@upr.edu
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Correspondence: José J. Hernández Ayala (jose.hernandezayala@sonoma.edu)

Abstract. This article analyzes the relationship between off-season tropical cyclone (TC) 11 frequency and climate variability and change for the Pacific and Atlantic Ocean basins. TC track 12 data was used to extract the off-season storms for the 1900-2019 period. TC counts were 13 aggregated by decade and the number of storms for the first six decades (pre-satellite era) was 14 adjusted. Mann-Kendall non-parametric tests were used to identify trends in decadal TC counts 15 and multiple linear regression models (MRL) were used to test if climatic variability or climate 16 17 change factors explained the trends in off-season storms. MRL stepwise procedures were implemented to identify the climate variability and change factors that explained most of the 18 variability in off-season TC frequency. A total of 713 TCs were identified as occurring earlier or 19 later than their peak seasons, most during the month of May and in the West Pacific and South 20 Pacific basins. The East Pacific (EP), North Atlantic (NA) and West Pacific (WP) basins exhibit 21 significant increasing trends in decadal off-season TC frequency. MRL results show that trends 22 in sea surface temperature, global mean surface temperature, and cloud cover explain most of 23 the increasing trend in decadal off-season TC counts in the EP, NA, and WP basins. Stepwise 24 MLR results also identified climate change variables as the dominant forces behind increasing 25 trends in off-season TC decadal counts, yet they also showed that climate variability factors like 26 El Niño-Southern Oscillation, the Atlantic Multidecadal Oscillation, and the Interdecadal Pacific 27 Oscillation also account for a portion of the variability. 28

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Keywords: Tropical Cyclones; Hurricane Season; Climate Variability; Climate Change

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# 34 1. Introduction

35 Increasingly, scientific evidence has shown a link between tropical cyclones (TC) and global

- 36 warming, especially following the dramatic rise in both the intensity and frequency of storms during
- the first two decades of the present century (Goldenberg et al., 2001; Holland and Webster, 2007).
- 38 Scientific studies (Landsea, 2005; Emanuel, 2005; Trenberth and Shea, 2006; Trenberth, 2007) are
- not in agreement as to whether sea surface temperatures have a measurable effect on the frequency
- 40 of tropical cyclones and other studies (Camargo and Sobel, 2005; Nogueira and Kim, 2007; Mahala
- 41 et al., 2015; Zhao et al., 2018] have evaluated cyclonic activity on a time scale longer than
- 42 interannual and have associated it with variability in the El Niño Southern Oscillation (ENSO), the
- 43 Atlantic Multidecadal Oscillation (AMO) and the Interdecadal Pacific Oscillation (IPO). However,
- 44 little is known about the changes in the frequency of off-season TCs, storms that occur before and
- 45 after the peak TC season months, and their connections to climate variability and change.

A number of recent papers (Wang and Lee, 2008; Knutson et al., 2010; Emanuel, 2013) have

47 documented global increases in the proportion of very intense cyclones as well as latitudinal trends

48 in maximum tropical cyclone (TC) intensity, which are consistent with future climate projections. A

- 49 detailed review of the behavior of TCs (Walsh et al., 2019) concluded that it remains uncertain
- 50 whether past changes in TC activity have exceeded the variability expected from natural causes,

- 51 while concerns remain about the temporal homogeneity of the best record (Landsea et al., 2006;
- 52 Mann et al., 2007). Another study (Mann et al., 2009) found that recent increases in the frequency
- 53 of intense TCs in the North Atlantic (NA) were the product of reinforcing effects, such as La Niña-
- like climate conditions and relative tropical Atlantic warming. Yet, no study has focused on 54
- 55 examining changing trends in off-season TC frequency and its relation to natural variability or
- 56 climate change.
- 57 A synthesis (Christensen et al., 2013) of the then-available regional projections of future TC
- 58 climatology for 2081–2100 in relation to 2000–2019, for a business as usual emissions scenario
- 59 similar to A1B, shows that worldwide the consensus projection was for decreases in TC numbers by
- approximately 5-30%, increased frequency of Category 4 and 5 storms between 0 and 25%, an 60
- 61 increase by a small percentage in the typical maximum intensity of life, and an increase in TC
- rainfall amounts by 5–20%. Nevertheless, it is clear that there is great uncertainty about these 62
- 63 projections. Such projections do not consider changes in off-season TC development in any of the
- 64 basins where TCs form
- 65 Several climatic reconstructions have been performed (Bradley et al., 2006; Mann et al., 2009)
- 66 using proxy data by collecting sediments from the impact of hurricanes in the period 500-1850 and
- 67 then calculated estimates from the statistical model of the activity of tropical cyclones based on
- 68 modern instrumental weather indexes for the period (1851–2006). In analyzing these results and
- comparing them with the cyclone seasons fixed by the World Meteorological Organization, the 69
- 70 hurricane season (tropical depressions, tropical storms and hurricanes) in the Atlantic Ocean was
- 71 fixed as June 1 to November 30 in 1960, yet we observe a significant variability in off-season TC
- occurrence before/after the hurricane season after the 1960s. 72
- 73 Even though it not uncommon for TCs to develop outside of their peak TC season months, there is a
- 74 need to examine trends in the number of storms that are forming during low activity months. The
- 75 formation of the extratropical storm Andrea on May 20, 2019 marks the decade of 2010 as that with
- 76 the greatest number of tropical cyclones in the Atlantic Ocean before or after the hurricane season
- 77 dates established by official bodies like the World Meteorological Organization (WMO) and the
- 78 National Oceanic and Atmospheric Administration (NOAA). The frequency of TCs in the North
- 79 Atlantic basin has been found to be influenced by fluctuations in teleconnections such as ENSO and
- 80 AMO (Trenberth et al., 2006; Nogueira et al., 2013). However, human-induced climate change
- manifested as higher sea surface temperatures (SST) and increasing evaporation rates in the tropical 81 and sub-tropical North Atlantic basin could also be related to the higher frequency of off-season 82
- 83
- tropical or extratropical cyclone occurrences in more recent decades. That increasing trend in SSTs 84 in the Atlantic and other ocean basins and its relation to out off-season TC occurrences during the
- 85 last century has not been thoroughly examined by the scientific community.
- 86 This study aims to determine if off-season TCs have increased in their frequency since the 1900 and
- if that increment in the number pre and post off-season storms could be associated with normal 87
- 88 climatic variability or climate change. The total number of out off-season TCs per decade for the
- 89 North Atlantic (NA), West Pacific (WP), East Pacific (EP) and South Pacific (SP) ocean basins
- 90 where analyzed to determine if any of the basins experienced an increase in the number of off-
- 91 season tropical/extratropical cyclones over time that could be associated to climatic variability or
- 92 climate change. The Indian Ocean basins were not included in this analysis due to limited data
- 93 availability. Ocean basins that were found to have statistically significant trends were then analyzed 94
- further with multiple liner regression models (MRL) and regression stepwise procedures to
- 95 determine if climate variability or change factors could explain increasing trends in off-season TC
- 96 frequency over time.
- 97

# 98 2. Data

99 Six-hourly TC track data for all storms across all ocean basins were obtained from the International

100 Best Track Archive for Climate Stewardship (IBTrACS) (Knapp et al., 2018) and all TCs that

101 occurred at or after 1900 were extracted. The TC tracks were then then extracted for the northern

102 hemisphere basins that include the East Pacific (EP), the North Atlantic (NA) and the West Pacific

103 (WP) and for the southern hemisphere basin in the South Pacific (SP) (Fig 1a). The off-season TCs

104 were then aggregated by decades in order to identify decadal variability or trends in total storm

- 105 counts at the individual basin scales. Off-season TCs were defined as storms that occurred in the
- three months before and after the six-month period of peak cyclone activity in the basin.
- 107 The monthly frequency of TCs for each basin were analyzed for the entire period and based on that
- analysis we determined that off-season TCs that occurred during the three months (Mar-Apr-May)
- before June 1<sup>st</sup> were pre-season and the three months (Dec-Jan-Feb) after November 30<sup>th</sup> were post-
- season in the northern hemisphere basins (NA, EP and WP) In the southern hemisphere, the three
- 111 months before (Aug-Sep-Oct) November 1st were classified as pre-season and the three months

112 (May-Jun-Jul) after April 30<sup>th</sup> were classified as post off-season in the southern hemisphere basins

113 (SP) (Fig 1b). Pre-season and post-season decadal time-series for the Northern/Southern hemisphere

- and individual basins were then constructed to calculate the total number of TCs per-decade from
- 115 1900 to 2019.

116 The climate variability indexes of ENSO (Niño 3.4), AMO (Trenberth et al., 2019) were

respectively obtained from the National Oceanic and Atmospheric Administration (NOAA)

- 118 Physical Sciences Lab and the National Center for Atmospheric Research. The IPO index was
- obtained from the NOAA Physical Sciences Laboratory (Henley et al., 2015). The variables
- associated with anthropogenic climate change used in this study were sea surface temperature
- 121 (SST), global mean surface temperature (GMST) and cloud cover (CC). SST data were obtained
- from the HadISST1 1° reconstruction, GMST data were accessed from the GISTEMP v4 and CC
- data was acquired from the International Comprehensive Ocean-Atmosphere Data Set (ICOADS)
- v2.5, all for the 1900-2019 period. The CC estimates from ICOADS are obtained from voluntary
- 125 observing ships that report CC in octas (eighths) ranging from 0 (completely clear sky) to 8
- 126 (completely overcast). These CC estimates are known to be temporally and spatially heterogenous
- with relatively high observational errors in some areas, yet increases in ocean coverage for clouds
- after the 1900 have been noted in the latest ICOADS 3.0 release [Eastman et al. 2011, Freeman et al. 2016; Aleksandrova et al. 2018]. A decadal average was calculated for all of the climate
- al. 2016; Aleksandrova et al. 2018]. A decadal average was calculated for all of the climate

variability and change variables to use them as predictors of decadal TC total counts (Table 1).

Table 1	Tronical cyclo	ne climate chanc	ne and variability	<i>i</i> variables use	d in this study
Table I.	TTOPICAL CYCIO	ne, chinale chang	je anu vanability	valiables use	a in this study.

	Abbreviations	Units
Tropical Cyclone Counts	TCs	Decadal Total Counts
Climate Change Variables		
Sea Surface Temperature	SST	0
Global Mean Surface Temperature	GMST	0
Cloud Cover	CC	Oktas
Climate Variability Variables		
El Niño Southern Oscillation	ENSO 3.4	°SST anomalies index
Interdecadal Pacific Oscillation	IPO	°SST anomalies index
Atlantic Multi-decadal Oscillation	AMO	°SST anomalies index

#### 131 3. TC Adjustment Method

TC counts before 1966 (pre-satellite era) are incomplete (Mann et al., 2007; Landsea, 2007) since a 132

133 lot of storms that didn't make landfall weren't recorded, so in order to make any comparisons

- between the earlier and later decades, the series for each basin need to be adjusted accordingly. 134
- 135 Another issue related to pre-satellite era TC track data is the undercount of weaker tropical
- 136 depressions, since the detection and classification of those weaker storms that showed poor
- 137 organization was probably more difficult before 1966 (Moon et al., 2019). The average landfall
- 138 percentage of TCs were calculated for the periods 1900-1965 (pre-satellite) and 1966-2019 (satellite
- 139 era and new TC monitoring technologies available) in order to determine the share of storms that 140
- made landfall in both periods. The percentage of landfalling TCs is expected to be higher in the 141 1900-65 period since a higher number of storms that remained over the ocean were not reported, so
- 142 the landfall percentage of the pre-satellite period is then adjusted so that it matches the 1966-2019
- 143 post-satellite period.

144 To obtain the estimated number of missing TCs for the 1900-65 period, the number of total storms

145 in the pre-satellite period is increased until its landfall percentage is equal to the one in the post

146 satellite era. The total number of additional TCs that resulted in the landfall percentages between the 147 two periods to be the same or near equal are then divided by the 7 decades of the pre-satellite era

148 and then the number of extra storms for each decade is multiplied by the percentage of off-season

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storms for each basin and that resulting number is then added then to each of the individual decades 150 between 1900 and 1969. In a previous study (Landsea, 2007), this method was applied to adjust TC

151 counts in the North Atlantic to determine if the basin has experienced an increasing trend in annual

152 TC frequency since the 1900, and its results show that after adjusting the tropical storm counts no

153 trends were found.

154 Here we show how this TC series adjustment method was applied to the total TC count for the NA basin for the 1900-2019 period. First, we calculate the landfall percentage for the pre-satellite period 155 1900-65 by dividing the number of landfalling TCs (LFTCs) with the total number of storms (TTCs) 156 157 and multiply by 100 to get the landfall percentage, check the equations below:

4 - 0

$$(LFTCs / TTCs) * 100$$
  
Example: (479/610) \* 100 = 78.5% (1)

159		
160		
161	Then calculate landfall % for the period post-satellite period 1966-2019,	
162		
	(LFTCs / TTCs) * 100	( <b>2</b> )
	<i>Example</i> : $(583/844) * 100 = 69.1\%$	(2)
163		
164		
165	Then artificially increase the number of TCs (+83 for the NA basin) until the landfall % of the 19	900-
166	65 period is equal to landfall % of the 1966-2019:	
167		
	LFTCs / (TTCs + AddTCs) * 100	(2)
	<i>Example</i> : $479/(610 + 83) * 100 = 69.1\%$	(3)
168		
169		
170	Then calculate the percentage (OffP) of off-season TCs (OffTCs) by dividing it by total number	of
171	TCs:	
172		

$$(OffTCs / TTCs) * 100$$
  
Example :  $(67/1454) * 100 = 4.61\%$  (4)

173

Then divide additional TCs (83) by the number of decades between 1900 and 1969 (7) and then
multiply by the off-season TC percentage (.0461)

- multiply by the off-season TC percentage (.046)
- 177

1

$$(AddTCs/Decades) * OffP$$
  
Example: (83/7) \* .0461 = 0.54 (5)

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In the case of the NA, we determined that by using the above TC series adjustment method the basin
would get an additional 0.54 off-season TCs for each of the seven decades that go from the 1900 to
1969. Finally, the additional 0.54 TCs per decade will be divided between pre and post off-season
TCs by multiplying the added storms with the respective percentage of pre/post off season cyclones:

DecOffTCs \* Percentage/Post Season Example: 0.54/0.62 = 0.33 and 0.54/0.38 = 0.21(6)

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The pre off-season decades of the NA basin before 1970 will get an additional 0.33 TCs and the post off-season decades will get 0.21 more storms. This off-season TC adjustment method was applied to the other five basins. It is important to note that this TC adjustment method has been only implemented in the NA basin and that this study is the first attempt to apply technique in other ocean basins. This adjustment method is in no way capable of detecting all TCs that formed before the satellite era, yet it offers us the opportunity to estimate missed storms by comparing the TC landfall percentage of the pre- and post-satellite eras.

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# 193 4. Statistical Methods & Models

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Mann-Kendall (MK) tests for trends (Mann, 1945; McLeod, 2005) were applied to all the offseason TC decadal series for all basins in order to determine if the frequency of storms has increased or decreased over time. This test has the advantage of not assuming any special form for the distribution function of the data, while having a power nearly as high as their parametric equivalents and that is why its use is highly recommended by the World Meteorological Organization (Hipel and McLeod, 2005).

201

The decadal series that were then found to have a significant trend based on the MK results were then furtherly analyzed by applying a series of multiple linear regression models (MLR). MLR were used to model the decadal count of off-season TCs for basins that showed increasing or decreasing trends in storm numbers to test if covariates associated with climatic variability and climate change explained off-season TC frequency. MLR attempts to model the relationship between two or more explanatory variables and a response variable by fitting a linear equation to observed data.

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Three different MLR models were run for each off-season TC series that exhibited a statistically significant trend, one MLR model with the climate change variables (SST, GMST & CC) as predictors, another model with just the climate variability factors (ENSO, AMO & IOD) and a final model with all of the variables included. Then the three MLR models (pre-season, post-season and off-season) were run for each of the basins with increasing trends in off-season TCs, the best models (highest adjusted R-squared and lowest p-value) were then selected for each of the series. The MLR models were run in The R Project for Statistical Computing using the biglm package.

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Finally, stepwise selection MLR models were used to identify the climate variability or change factors making the most statistically significant contributions to off-season increasing TC frequency. Here we use stepwise selection which is a combination of the forward and backward procedures where you start with no predictors, then sequentially add the most contributive predictors. After adding each new variable, it removes the variables that no longer provide an improvement in the model fit (James et al., 2014; Bruce and Bruce, 2017). The MLR and stepwise for the off-season TC count series for each of the basins with significant increasing trends were run in The R Project for Statistical Computing using the MASS package (Venables and Ripley, 2002).

### 225 5. Results & Discussion

226 When analyzing the number of TCs for all basins for the 1900-2019 period we found that 713 227 off-season storms occurred during that time, most during the months of May (NH pre-season and SH post-season) with 430 and December (NH post-season) with 341 (Figure 1a, 1b). When looking at 228 the count of off-season TCs per basin we found that as expected the West Pacific (611) and South 229 230 Pacific (85) accounted for 81.3% of all off-season storm occurrences. When grouping the basins 231 between northern and southern hemispheres, we find that 89% of all off -season TCs occurred north 232 of the equator for the 1900-2019 period (Figure 1a, 1b). The North Atlantic and East Pacific basins 233 were found to be the ones with the lowest numbers of off-season TCs when compared to the other 234 two Pacific basins.





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Figure 1. Tracks of all out off-season TCs (a) and the number of storms per month for the WP (b),
SP (c), EP (d) and NA (e) basins for the 1900-2019 period.

239 As shown in Figure 2, even after adding the estimated missing storms before the 1970 decade, 240 most basins experienced their highest number of out off-season TCs (pre or post) in decades at or 241 after 1960-69. The 1960-69 decade for the northern hemisphere basins (WP, NA and EP) was found to be the one with the highest number of pre off-season TCs with 69 and the 1950-1959 decade was 242 243 identified as the one with the most post off-season storms with 68 (Figure 2a, 2b, 2c and 2d). When 244 examining TC counts for all basins individually, we found that the NA and EP basins had their most 245 active decades after 1970 and that the WP and SP basins experienced their highest storm count decade 246 after 1960 (Figure 2c, 2d). It is important to note that these results already reflect the additional TCs 247 that were added to the pre-satellite era.

248





Pre-Season Post-Season



252 The Mann-Kendall non-parametric tests for trends for all basins show that three basins exhibited 253 statistically significant increasing trends in adjusted decadal off-season TC counts for the 1900-2019 254 period (Table 2). The basins with statistically significant increasing trends for the entire time period 255 were the EP (pre and off-season), NA (pre, post and off-season) and the WP (pre, post and off-season). 256 The EP basin shows an increasing trend in pre- and off-season TCs that is more evident from the 257 1950s to the present (Figure 2d), while the increasing trend in the NA basin can be observed from the 258 1970s to 2019. The NA and EP basins also exhibit positive trends when the analysis is done only 259 considering the post satellite era, yet the results are not statistically significant (Table 2)

260 The increasing trend in off-season TCs in the WP basin is more evident from the 1900 to 1969 261 (Figure 2a), which was during the pre-satellite era where missing TCs were added to the series. However, a negative trend is found in the WP basin if the decadal counts are analyzed from the 1960s 262 263 to the present, yet those results were not statistically significant. In the post-satellite era in the WP 264 basin, the 1990-1999 decade was identified as the one with most off-season TCs, however the two 265 following decades exhibited a decreasing trend. The EP and NA basins show significant increasing trends in off-season TC counts (Table 2). Opposites to the trends identified in the WP basin, the EP 266 267 and NA also show increasing decadal counts after the 1960s and 1970s. The SP basin also exhibited 268 a positive Tau coefficient, yet it was statistically insignificant for the entire period and the post-269 satellite era (Figure 2d).

270

Table 2. Results of Mann-Kendall trend tests for the 1900-2019 period for all ocean bas	sins
where TCs form.	

Trends for the 1900-2019 period								
Pre TCs	Tau S	P-value	Post TCs	Tau S	P-value	Off TCs	Tau S	P-value
EP	0.746	0.002	EP	0.098	0.723	EP	0.679	0.004
NA	0.572	0.015	NA	0.485	0.042	NA	0.554	0.016
SP	0.048	0.889	SP	0.015	1.000	SP	0.061	0.836
WP	0.554	0.016	WP	0.485	0.034	WP	0.534	0.019
Trends for the 1960-2019 period								
EP	0.596	0.158	EP	0.414	0.338	EP	0.69	0.085
NA	0.596	0.158	NA	0.645	0.119	NA	0.6	0.132
SP	0.596	0.158	SP	-0.2	0.707	SP	-0.06	1
WP	-0.467	0.259	WP	-0.6	0.132	WP	-0.69	0.085

Significant trends in bold.

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272 MLR models were run on the basins that exhibited statistically significant (< 0.05) increasing 273 trends in decadal total off-season TC counts over time and here we report the best models for each of the series. The MLR results show that the statistically significant increasing trends in TC frequency 274 for the EP (pre and off-season) and WP basins is best explained by climate change factors SST, GMST 275 and CC in both the normal and detrended series at the 0.05 significance level (Table 3). Climate 276 277 change factors accounted for 56% (pre-season) and 52% (off-season) of the increasing trend in TC 278 counts for the EP basin. In the WP basin climate change factors explained 55% (pre-season), 64% 279 (post-season) and 68% (off-season) of the trends in off-season TCs, yet lower R squares were found 280 when the analysis was done with the detrended series. Increasing trends in SSTs, GMST and moisture 281 (CC) outside of the prime months of tropical storm development could promote better conditions for 282 higher off-season TC occurrences (Klozback, 2006; Hansen et al., 2010).

Table 3. Best multiple linear regression models (MLR) for basins with statistically significant increasing trends in off-season TCs with detrended climate indices.

Model	Adj. R²	Adj. R² Det.	Factors	p-val.	p-val Det.		
EP pre-season	0.563	0.444	SST, GMST & CC	0.021	0.038		
EP off-season	0.522	0.472	SST, GMST & CC	0.030	0.024		
NA pre-season	0.481	0.496	SST, GMST & CC	0.041	0.022		
NA post-season	0.427	0.247	ENSO & AMO	0.130	0.057		
NA off-season	0.384	0.406	SST, GMST & CC	0.070	0.010		
WP pre-season	0.551	0.462	SST, GMST & CC	0.020	0.000		

WP post-season	0.645	0.478	SST, GMST & CC	0.000	0.023
WP off-season	0.689	0.481	SST, GMST & CC	0.005	0.017

### 283

284 The climate variability factors (ENSO, AMO & IOD) did not exhibit statistically significant 285 relationships with increasing off-season TC counts, which shows that natural variability does not explain the incrementing number of storms in the EP and WP basins. MLR model results for the 286 287 NA basin also showed the climate change variables accounting for 48% (pre-season) and 38% (offseason) of the increasing trend in TCs (Table 3). However, the MLR model results for the post-288 289 season months in the NA basin showed that the climate variability variables (ENSO & AMO) 290 accounted for 42% (25% in the detrended) of the increasing trend in TCs, yet the model was not 291 found to be statistically significant in both the normal and detrended series. It is well known that 292 cold phases of ENSO (La Niña) and warm phases of AMO tend to be associated with higher TC 293 frequency in the North Atlantic ocean (Tang and Neelin, 2004; Briggs, 2008) and this could explain 294 why those teleconnections were found to have the most significant influence on post-season TC frequency in the NA basin. When the MLR results of the original and detrended series are compared 295 296 (Table 3), we find that the models with the detrended series exhibit lower R squares than the MLR 297 models with the original series, yet those models were still found to be statistically significant 298 which suggests that the correlation between off-season TCs and climate change factors is strong 299 even after decadal trends are removed.

300 Stepwise MLR model results showed that climate change factors (SST, GMST & CC) were among 301 the selected variables that explained most of the increasing trend in off-season TCs for all basins 302 analyzed (Table 4). In the EP basin, SST, ENSO, and CC, accounted for 69% (pre-season) and 65% 303 (off-season) of the increasing trend in TCs. In the NA basin, the stepwise procedure selected CC as the sole climate change factor that explained 52% (pre-season) and 40% (off-season) of the rising 304 305 frequency in TC counts. However, CC & AMO were selected as the variables that explained (43%) 306 most of the variability in TC frequency during the post-season months in the NA basin. Stepwise procedure results for the WP basin show that climate change and variability factors were selected as 307 308 the best predictors of TC frequency, with GMST and CC accounting for 57% (pre-season), CC, GMST, ENSO and IPO explaining 72% (post-season) and 74% (off-season) of the variability of 309 310 TCs.

Model	R-squared	Adjusted R <sup>2</sup>	Factors	p-value
EP pre-season	0.777	0.694	SST, ENSO & CC	0.005
EP off-season	0.747	0.652	SST, ENSO & CC	0.008
NA pre-season	0.569	0.526	CC	0.004
NA post-season	0.687	0.427	CC & AMO	0.098
NA off-season	0.460	0.406	CC	0.015
WP pre-season	0.655	0.578	GMST & CC	0.008
WP post-season	0.826	0.726	CC, GMST, ENSO & IPO	0.008
WP off-season	0.839	0.747	CC, GMST, ENSO & IPO	0.006

Table 4. Stepwise multiple linear regression models (MLR) for basins with statistically significant increasing trends in off-season TCs.

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The EP experienced a steady increase in off-season TC total counts from 1900 to 2019 at a rate of 1.1 additional storms per decade. The decadal off-season total TC count series for the EP basin closely resembles the increasing trend in average SSTs and CCs (Fig 3a, 3c). When the EP off-season 315 TC tracks are examined, it shows that most storms have formed in areas that have experienced 316 statistically significant increasing trends in SST and CC, yet no changes in track or genesis location were detected over time. (Fig 3d, 3e). The correlation between off-season TCs in the EP basin and 317 ENSO is not as clear as the one between SST and CC, with some mostly warm ENSO decades like 318 319 the 1990-1999 exhibiting lower storm counts and other periods with cooler phases dominating 320 showing a higher number of cyclones. When SST patterns for areas in the EP basin where TCs develop are examined over time, we find that most tropical/sub-tropical ocean waters have 321 322 experienced a statistically significant increasing trend in ocean surface temperatures from 1900 to 2019 (Fig 3d). Similar to other studies (Hansen et al., 2010), we find that the EP tropical ocean 323 surfaces have increased by 0.051 degrees C° per decade. When CC patterns are examined, we find 324 that it has also experienced a statistically significant increasing trend in some areas in the EP basin 325 326 (Fig 3e).

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Figure 3. Decadal TC counts for the EP off-seasons and decadal average SSTs (a), decadal TC counts for
the EP off-seasons and decadal average ENSO 3.4 (b), decadal TC counts for the EP off-seasons and
decadal average Correlation between Time and Dec-May averaged CC (c), correlation between Time and
Dec-May averaged SST (C°) and the off-season TC tracks for the 1900-2019 period (d) and correlation
between Time and Dec-May averaged CC (oktas) and the off-season TC tracks for the 1900-2019 period.

The decadal off-season total TC count series for the NA basin closely resembles the increasing trend in average SSTs and CCs (Fig 4a, 4c). The NA decadal series shows a steady increase 336 in off-season TC total counts from 1900 to 2019 at a rate of 0.7 additional storms per decade and an 337 SST increasing trend of 0.055 C° per decade. Both the average decadal SST and CC series coincide with the peaks and valleys in off-season TC counts for the NA basin, with the 1950-1959 showing a 338 339 high number of storms associated with high mean SSTs and CCs while the drop in storm counts in 340 the 1960-1969 decade matches a drastic drop in ocean surface temperatures (Figures 4a, 4c). The off-341 season TC tracks in the NA basin also formed in areas that exhibited increasing trends in SST and 342 CC, however no changes in track or genesis location were detected over time (Figure 4d, 4e). Even though average SSTs increase to 0.135 C° per decade from 1970 to 2019, off-season TC total counts 343 went down in the 1990-1999 and 2010-2019 decades, with the decade in between (2000-2009) 344 exhibiting the highest number of off-season TCs (14) of all decades examined. However, it is 345 important to note that 5 out of the 6 decades with the most off-season TCs in the NA basin occurred 346 347 after the 1970s.

348

349 When North Atlantic SSTs are examined in areas where TCs form, we found that ocean surface temperatures have increased at a rate of 0.055 degrees C° per decade for the off-season months of 350 Dec-March (Fig 4d). When CC patterns are examined, we find that it has also experienced a 351 352 statistically significant increasing trend of 0.06 oktas (eighths of the sky that are covered in clouds) 353 per decade in the North Atlantic basin since the 1900 (Figure 4e). If the NA pre/post off-season series 354 is modified to begin in the 1960s, we find that SSTs have increased at a decadal rate of 0.082 C° per 355 decade at a rate of 1.2 additional storms per decade. Overall, these results suggest that increasing 356 trends in SSTs, which also drive increasing trends in evaporation rates associated with high CCs, are 357 the physical mechanisms behind most of the recent increase in the total number of out of season TCs 358 in the NA basin. The correlation between off-season TCs in the NA basin and AMO is not as clear as 359 the one between SST and CC, with some warm AMO phases between 1930-1959 exhibiting lower 360 storm counts while some cooler phases (1970-89) showing a higher number of cyclones.



361

Figure 4. Decadal TC counts for the NA off-seasons and decadal average SSTs (a), decadal TC counts for the
NA off-seasons and decadal average AMO (b), decadal TC counts for the NA off-seasons and decadal average
Correlation between Time and Dec-May averaged CC (c), correlation between Time and Dec-May averaged
SST (C°) and the off-season TC tracks for the 1900-2019 period (d) and correlation between Time and Dec-May averaged
May averaged CC (oktas) and the off-season TC tracks for the 1900-2019 period (e).

368 The decadal off-season total TC count series for the WP basin closely resembles the increasing 369 trend in GMST (Fig 5a). However, the WP basin experienced the highest count of off-season TCs in 370 the 1960-69 decade, not in the more recent decades like the EP and NA basins. More importantly, if trend analysis for off-season TC counts is done from 1960-2019 in the WP basin, we find a none 371 372 statistically significant decreasing trend. However, it is important to note that four out of the five decades with most off-season TCs in the WP basin occurred after 1960. However, the 2010-2019 373 374 decade was identified as the period with the lowest total number of off-season TCs even though 375 increasing trends in mean SST, GMST and CC continued (Fig 5a, 5d and 5e). Off-season TC tracks 376 in the WP basin also correlate spatially with areas that show increasing trends in SST and CC (Fig 5d, 5e), yet no changes in TC genesis or tracks were identified over time. The decreasing number of 377 378 off-season TCs in the last two decades coincided with a negative phase of the IPO, which suggests 379 that TC frequency in the WP basin is influenced by fluctuations in the IPO (Fig 5c), whose recent negative phase since 1998 resembles La Niña-like SST anomaly patterns (Zhao et.al, 2018). Even 380

though most of the variability in off-season TC frequency in the WP basin can be explained by climate

change trends in GMST, SST and CC, the rest of the variance in TCs is account by fluctuations in the

383 IPO and ENSO teleconnections.

384



Figure 5. Decadal TC counts for the WP off-seasons and decadal average SSTs (a), decadal TC counts for the
WP off-seasons and decadal average AMO (b), decadal TC counts for the WP off-seasons and decadal average
Correlation between Time and Dec-May averaged CC (c), correlation between Time and Dec-May averaged
SST (C°) and the off-season TC tracks for the 1900-2019 period (d) and correlation between Time and DecMay averaged CC (oktas) and the off-season TC tracks for the 1900-2019 period (e).

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385

392 Studies that have examined TC frequency overall have found increases in the number of most 393 intense hurricanes [Wang and Lee, 2008; Knutson et al., 2010; Emanuel, 2013], yet no clear trend has 394 been found when lower intensity TCs have been examined [Landsea, 2007]. The results of other 395 studies show that there is no overall agreement on the relationship between SSTs and TC frequency 396 (Landsea, 2005; Emanuel, 2005; Trenberth and Shea, 2006; Trenberth, 2007), yet some have found 397 strong associations between TC variability and ENSO, AMO and IPO [Camargo and Sobel, 2005; 398 Nogueira and Kim, 2007; Mahala et al., 2015; Zhao et al., 2018]. In this study we analyzed off-season 399 TCs and our results differ from those that have found no trend in overall TC frequency, since we 400 found decadal increasing trends in the NA and EP basin in both the pre and post-satellite eras. The 401 results presented here suggest that climate change trends like increasing SSTs and more favorable

402 moisture environments (CC) between the months of Dec to May in the NA and EP basins seem to be403 the major factors behind decadal increasing trends in off-season TCs.

404

## 405 6. Summary and concluding remarks

406 The frequency of TCs that developed outside of their prime season months were analyzed to 407 determine if trends in higher storm totals in the Pacific and Atlantic Ocean basins were associated 408 with natural variability, climate change or both. Adjusted off-season decadal TC total counts for six 409 ocean basins were analyzed for the 1900-2019 period in order to determine if the number of storms 410 have been increasing over time. Mann-Kendall tests for trends were done and the basins that exhibited 411 statistically significant increasing trends were then furtherly analyzed using multiple linear regression 412 models and stepwise procedures to determine if those trends could be explained by fluctuations 413 associated with climate variability, climate change trends or a combination of both.

414

415 The main results of this study suggest that decadal total off-season (pre/post) TC counts have significantly increase over time since the 1900 in the East Pacific (EP). North Atlantic (NA) and West 416 Pacific (WP) basins. The EP and NA basins exhibited statistically significant increasing trends even 417 418 if the analysis was done from the 1960s instead of the 1900. The WP basin showed an overall increasing trend in the total number of off-season TCs per decade, yet if the analysis is done from the 419 420 1960s to the present, no statistically significant increasing trend is found. However, the three basins 421 that reflected an overall increase in decadal off-season TC frequency had their most active decades 422 after the 1970s.

423

Results from the best multiple linear regression (MLR) models show that the increasing decadal count of off-season TCs has been found to be strongly associated with climate change trends in sea surface temperature (SST), global mean surface temperature (GMST) and cloud cover (CC) in all three basins (EP, NA and WP). The MLR model where climate variability variables (ENSO and AMO) explained most of the variance in off-season TC counts was in the storm decadal counts for the post-season months of the NA basin.

430

Results of the MLR stepwise procedures showed that the selected variables that accounted for most of the variability in off-season TCs for the EP basin were SST, CC and ENSO, while CC (preseason and off-season) and AMO (post-season) were chosen as the best variables for the NA basin. The stepwise procedure identified the climate change trends in GMST and CC, and fluctuations in ENSO and IPO as the variables that accounted for most of the variability in decadal off-season total TC counts in the WP basin,

437

438 The findings of this study suggest that trends in SST, GMST and CC associated with climate 439 change are not only altering the frequency (Klotzbach, 2006; Saunders and Lea, 2006; Hansen et al., 440 2010) and intensity of TCs that develop during the peak months of the season, they are also altering 441 the total number of storms that form in the off-season months (Dec-May), especially in the EP and 442 NA basins. The results of this study have important implications for the NA and EP basins, if off-443 season TCs have been increasing in frequency since the 1900 we can expect that this trend associated 444 with climate change would continue in future decades. This increasing number of off-season TCs 445 could potentially impact societies in their path during times of the year when storms are least 446 expected, possibly increasing environmental and economic impacts in areas that are already 447 experiencing the effects of climate change exacerbated phenomena.

448

One of the main limitations of this work was the inclusion of tropical depressions in the offseason TC analysis. If data on TC intensity were widely available for all off-season TCs, it would have been possible to exclude weaker tropical depressions from the analysis since the detection and classification of those storms was more difficult in the pre-satellite era. Other limitations of this study include the issues of worst data quality in the pre-satellite era, the problem of applying a universal missed TC adjustment method to all basins analyzed and the lack of information on TC intensity formany storms, especially in the pre-satellite era.

456 457

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- surface temperature (HadISST1 1° reconstruction), and cloud cover (ICOADS v2.5 1°) datasets
- supporting this article are based on publicly available measurements that can be accessed from the
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- 468 available measurements that from the NASA Goddard Institute for Space Studies (GISS:
- 469 https://climatedataguide.ucar.edu/climate-data/global-surface-temperature-data-gistemp-nasa-
- 470 goddard-institute-space-studies-giss.) The El Niño Southern Oscillation (ENSO 3.4) data supporting
- this article are based on publicly available measurements from the National Oceanic and
- 472 Atmospheric Administration Physical Sciences Lab (PSL;
- 473 <u>https://psl.noaa.gov/gcos\_wgsp/Timeseries/Data/nino34.long.data</u>). The Atlantic Multidecadal
- 474 Oscillation data supporting this article are based on publicly available measurements from the
- 475 National Center for Atmospheric Research (NCAR; <u>https://climatedataguide.ucar.edu/climate-</u>
- 476 <u>data/atlantic-multi-decadal-oscillation-amo</u>). The Interdecadal Pacific Oscillation data supporting
- this article are based on publicly available measurements from the National Oceanic and
- 478 Atmospheric Administration Physical Sciences Lab (PSL;
- 479 <u>https://psl.noaa.gov/data/timeseries/IPOTPI/</u>).

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