



1 **Increasing Frequency in Off-Season Tropical Cyclones and its** 2 **relation to Climate Variability and Change**

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11 **Abstract.** This article analyzes the relationship between off-season tropical cyclone (TC)
12 frequency and climate variability and change for the Pacific and Atlantic Ocean basins. TC track
13 data was used to extract the off-season storms for the 1900-2019 period. TC counts were
14 aggregated by decade and the number of storms for the first six decades (pre-satellite era) was
15 adjusted. Mann-Kendall non-parametric tests were used to identify trends in decadal TC counts
16 and multiple linear regression models (MRL) were used to test if climatic variability or climate
17 change factors explained the trends in off-season storms. MRL stepwise procedures were
18 implemented to identify the climate variability and change factors that explained most of the
19 variability in off-season TC frequency. A total of 713 TCs were identified as occurring earlier or
20 later than their peak seasons, most during the month of May and in the West Pacific and South
21 Pacific basins. The East Pacific (EP), North Atlantic (NA) and West Pacific (WP) basins exhibit
22 significant increasing trends in decadal off-season TC frequency. MRL results show that trends
23 in sea surface temperature, global mean surface temperature, and cloud cover explain most of
24 the increasing trend in decadal off-season TC counts in the EP, NA, and WP basins. Stepwise
25 MLR results also identified climate change variables as the dominant forces behind increasing
26 trends in off-season TC decadal counts, yet they also showed that climate variability factors like
27 El Niño-Southern Oscillation, the Atlantic Multidecadal Oscillation, and the Interdecadal Pacific
28 Oscillation also account for a portion of the variability.

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31 **Keywords:** Tropical Cyclones; Hurricane Season; Climate Variability; Climate Change

32 33 34 **1. Introduction**

35 Increasingly, scientific evidence has shown a link between tropical cyclones (TC) and global
36 warming, especially following the dramatic rise in both the intensity and frequency of storms during
37 the first two decades of the present century (Goldenberg et al., 2001; Holland and Webster, 2007).
38 Scientific studies (Landsea, 2005; Emanuel, 2005; Trenberth and Shea, 2006; Trenberth, 2007) are
39 not in agreement as to whether sea surface temperatures have a measurable effect on the frequency
40 of tropical cyclones and other studies (Camargo and Sobel, 2005; Nogueira and Kim, 2007; Mahala
41 et al., 2015; Zhao et al., 2018] have evaluated cyclonic activity on a time scale longer than
42 interannual and have associated it with variability in the El Niño Southern Oscillation (ENSO), the
43 Atlantic Multidecadal Oscillation (AMO) and the Interdecadal Pacific Oscillation (IPO). However,
44 little is known about the changes in the frequency of off-season TCs, storms that occur before and
45 after the peak TC season months, and their connections to climate variability and change.

46 A number of recent papers (Wang and Lee, 2008; Knutson et al., 2010; Emanuel, 2013) have
47 documented global increases in the proportion of very intense cyclones as well as latitudinal trends
48 in maximum tropical cyclone (TC) intensity, which are consistent with future climate projections. A
49 detailed review of the behavior of TCs (Walsh et al., 2016) concluded that it remains uncertain
50 whether past changes in TC activity have exceeded the variability expected from natural causes,



51 while concerns remain about the temporal homogeneity of the best record (Landsea et al., 2006;
52 Mann et al., 2007). Another study (Mann et al., 2009) found that recent increases in the frequency
53 of intense TCs in the North Atlantic (NA) were the product of reinforcing effects, such as La Niña-
54 like climate conditions and relative tropical Atlantic warming. Yet, no study has focused on
55 examining changing trends in off-season TC frequency and its relation to natural variability or
56 climate change.

57 A synthesis (Christensen et al., 2013) of the then-available regional projections of future TC
58 climatology for 2081–2100 in relation to 2000–2019, for a business as usual emissions scenario
59 similar to A1B, shows that worldwide the consensus projection was for decreases in TC numbers by
60 approximately 5–30%, increased frequency of Category 4 and 5 storms between 0 and 25%, an
61 increase by a small percentage in the typical maximum intensity of life, and an increase in TC
62 rainfall amounts by 5–20%. Nevertheless, it is clear that there is great uncertainty about these
63 projections. Such projections do not consider changes in off-season TC development in any of the
64 basins where TCs form

65 Several climatic reconstructions have been performed (Bradley et al., 2006; Mann et al., 2009)
66 using proxy data by collecting sediments from the impact of hurricanes in the period 500–1850 and
67 then calculated estimates from the statistical model of the activity of tropical cyclones based on
68 modern instrumental weather indexes for the period (1851–2006). In analyzing these results and
69 comparing them with the cyclone seasons fixed by the World Meteorological Organization, the
70 hurricane season (tropical depressions, tropical storms and hurricanes) in the Atlantic Ocean was
71 fixed as June 1 to November 30 in 1960, yet we observe a significant variability in off-season TC
72 occurrence before/after the hurricane season after the 1960s.

73 The formation of the extratropical storm Andrea on May 20, 2019 marks the decade of 2010 as that
74 with the greatest number of tropical cyclones in the Atlantic Ocean before or after the hurricane
75 season dates established by official bodies like the World Meteorological Organization (WMO) and
76 the National Oceanic and Atmospheric Administration (NOAA). The frequency of TCs in the North
77 Atlantic basin has been found to be influenced by fluctuations in teleconnections such as ENSO and
78 AMO (Trenberth et al., 2006; Nogueira et al., 2013). However, human-induced climate change
79 manifested as higher sea surface temperatures (SST) and increasing evaporation rates in the tropical
80 and sub-tropical North Atlantic basin could also be related to the higher frequency of off-season
81 tropical or extratropical cyclone occurrences in more recent decades. That increasing trend in SSTs
82 in the Atlantic and other ocean basins and its relation to out off-season TC occurrences during the
83 last century has not been thoroughly examined by the scientific community.

84 This study aims to determine if off-season TCs have increased in their frequency since the 1900 and
85 if that increment in the number pre and post off-season storms could be associated with normal
86 climatic variability or climate change. The total number of out off-season TCs per decade for the
87 North Atlantic (NA), West Pacific (WP), East Pacific (EP) and South Pacific (SP) ocean basins
88 where analyzed to determine if any of the basins experienced an increase in the number of off-
89 season tropical/extratropical cyclones over time that could be associated to climatic variability or
90 climate change. The Indian Ocean basins were not included in this analysis due to limited data
91 availability. Ocean basins that were found to have statistically significant trends were then analyzed
92 further with multiple liner regression models (MRL) and regression stepwise procedures to
93 determine if climate variability or change factors could explain increasing trends in off-season TC
94 frequency over time.

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98 **2. Data**

99 Six-hourly TC track data for all storms across all ocean basins were obtained from the International
 100 Best Track Archive for Climate Stewardship (IBTrACS) (Knapp et al., 2018) and all TCs that
 101 occurred at or after 1900 were extracted. The TC tracks were then then extracted for the northern
 102 hemisphere basins that include the East Pacific (EP), the North Atlantic (NA) and the West Pacific
 103 (WP) and for the southern hemisphere basin in the South Pacific (SP) (Fig 1a). The off-season TCs
 104 were then aggregated by decades in order to identify decadal variability or trends in total storm
 105 counts at the individual basin scales. Off-season TCs were defined as storms that occurred in the
 106 three months before and after the six-month period of peak cyclone activity in the basin.

107 The monthly frequency of TCs for each basin were analyzed for the entire period and based on that
 108 analysis we determined that off-season TCs that occurred during the three months (Mar-Apr-May)
 109 before June 1st were pre-season and the three months (Dec-Jan-Feb) after November 30th were post-
 110 season in the northern hemisphere basins (NA, EP and WP) In the southern hemisphere, the three
 111 months before (Aug-Sep-Oct) November 1st were classified as pre-season and the three months
 112 (May-Jun-Jul) after April 30th were classified as post off-season in the southern hemisphere basins
 113 (SP) (Fig 1b). Pre-season and post-season decadal time-series for the Northern/Southern hemisphere
 114 and individual basins were then constructed to calculate the total number of TCs per-decade from
 115 1900 to 2019.

116 The climate variability indexes of ENSO (Niño 3.4), AMO (Trenberth et al., 2019) and IOD were
 117 respectively obtained from the National Oceanic and Atmospheric Administration (NOAA)
 118 Physical Sciences Lab, the National Center for Atmospheric Research, and the Australian
 119 Government Bureau of Meteorology. The IPO index was obtained from the NOAA Physical
 120 Sciences Laboratory (Henley et al., 2015). The variables associated with anthropogenic climate
 121 change used in this study were sea surface temperature (SST), global mean surface temperature
 122 (GMST) and cloud cover (CC). SST data were obtained from the HadISST1 1° reconstruction,
 123 GMST data were accessed from the GISTEMP v4 and CC data was acquired from the ICOADS
 124 v2.5, all for the 1900-2019 period. A decadal average was calculated for all of the climate
 125 variability and change variables in order to use them as predictors of decadal TC total counts (Table
 126 1).

Table 1. Tropical cyclone, climate change and variability variables used in this study.

	Abbreviations	Units
Tropical Cyclone Counts	TCs	Decadal Total Counts
Climate Change Variables		
Sea Surface Temperature	SST	°
Global Mean Surface Temperature	GMST	°
Cloud Cover	CC	Oktas
Climate Variability Variables		
El Niño Southern Oscillation	ENSO 3.4	°SST anomalies index
Interdecadal Pacific Oscillation	IPO	°SST anomalies index
Atlantic Multi-decadal Oscillation	AMO	°SST anomalies index

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130 3. TC Adjustment Method

131 TC counts before 1966 (pre-satellite era) are incomplete (Mann et al., 2007; Landsea, 2007) since a
132 lot of storms that didn't make landfall weren't recorded, so in order to make any comparisons
133 between the earlier and later decades, the series for each basin need to be adjusted accordingly. The
134 average landfall percentage of TCs were calculated for the periods 1900-1965 (pre-satellite) and
135 1966-2019 (satellite era and new TC monitoring technologies available) in order to determine the
136 share of storms that made landfall in both periods. The percentage of landfalling TCs is expected to
137 be higher in the 1900-65 period since a higher number of storms that remained over the ocean were
138 not reported, so the landfall percentage of the pre-satellite period is then adjusted so that it matches
139 the 1966-2019 post-satellite period.

140 To obtain the estimated number of missing TCs for the 1900-65 period, the number of total storms
141 in the pre-satellite period is increased until its landfall percentage is equal to the one in the post
142 satellite era. The total number of additional TCs that resulted in the landfall percentages between the
143 two periods to be the same or near equal are then divided by the 7 decades of the pre-satellite era
144 and then the number of extra storms for each decade is multiplied by the percentage of off-season
145 storms for each basin and that resulting number is then added then to each of the individual decades
146 between 1900 and 1969. In a previous study (Landsea, 2007), this method was applied to adjust TC
147 counts in the North Atlantic to determine if the basin has experienced an increasing trend in annual
148 TC frequency since the 1900, and its results show that after adjusting the tropical storm counts no
149 trends were found.

150 Here we show how this TC series adjustment method was applied to the total TC count for the NA
151 basin for the 1900-2019 period. First, we calculate the landfall percentage for the pre-satellite period
152 1900-65 by dividing the number of landfalling TCs (LFTCs) with the total number of storms (TTCs)
153 and multiply by 100 to get the landfall percentage, check the equations below:
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$$(LFTCs / TTCs) * 100$$

Example: $(479/610) * 100 = 78.5\%$ (1)

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157 Then calculate landfall % for the period post-satellite period 1966-2019,
158

$$(LFTCs / TTCs) * 100$$

Example: $(583/844) * 100 = 69.1\%$ (2)

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161 Then artificially increase the number of TCs (+83 for the NA basin) until the landfall % of the 1900-
162 65 period is equal to landfall % of the 1966-2019:
163

$$LFTCs / (TTCs + AddTCs) * 100$$

Example: $479 / (610 + 83) * 100 = 69.1\%$ (3)

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166 Then calculate the percentage (OffP) of off-season TCs (OffTCs) by dividing it by total number of
167 TCs:
168

$$(OffTCs / TTCs) * 100$$

Example: $(67/1454) * 100 = 4.61\%$ (4)

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171 Then divide additional TCs (83) by the number of decades between 1900 and 1969 (7) and then
multiply by the off-season TC percentage (.0461)



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$$(AddTCs/Decades) * OffP$$

$$Example: (83/7) * .0461 = 0.54 \quad (5)$$

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In the case of the NA, we determined that by using the above TC series adjustment method the basin would get an additional 0.54 off-season TCs for each of the seven decades that go from the 1900 to 1969. Finally, the additional 0.54 TCs per decade will be divided between pre and post off-season TCs by multiplying the added storms with the respective percentage of pre/post off season cyclones:

$$DecOffTCs * Percentage/Post Season$$

$$Example: 0.54/0.62 = 0.33 \text{ and } 0.54/0.38 = 0.21 \quad (6)$$

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The pre off-season decades of the NA basin before 1970 will get an additional 0.33 TCs and the post off-season decades will get 0.21 more storms. This off-season TC adjustment method was applied to the other five basins.

185 4. Statistical Methods & Models

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Mann-Kendall (MK) tests for trends (Mann, 1945; McLeod, 2005) were applied to all the off-season TC decadal series for all basins in order to determine if the frequency of storms has increased or decreased over time. This test has the advantage of not assuming any special form for the distribution function of the data, while having a power nearly as high as their parametric equivalents and that is why its use is highly recommended by the World Meteorological Organization (Hipel and McLeod, 2005). The Mann-Kendall rank statistic t is calculated according to:

$$S = \sum_{k=1}^{n-1} \sum_{j=k+1}^n \text{sgn}(X_j - X_k) \quad (1)$$

194 with
 195

$$\text{gn}(x) = \begin{cases} 1 & \text{if } x > 0 \\ 0 & \text{if } x = 0 \\ -1 & \text{if } x < 0 \end{cases} \quad (2)$$

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The mean of S is $E[S] = 0$ and the variance σ^2 is

$$\sigma^2 = \{n(n-1)(2n+5) - \sum_{j=1}^p t_j(t_j-1)(2t_j+5)\}/18 \quad (3)$$

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where p is the number of the tied groups in the data set and t_j is the number of data points in the j th tied group. The statistic S is approximately normal distributed provided that the following Z -transformation is employed:

$$Z = \begin{cases} \frac{S-1}{\sigma} & \text{if } S > 0 \\ 0 & \text{if } S = 0 \\ \frac{S+1}{\sigma} & \text{if } S < 0 \end{cases} \quad (4)$$

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The statistic S is closely related to Kendall's τ as given by:



$$\tau = \frac{S}{D} \quad (5)$$

206
207 where
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$$D = \left[\frac{1}{2}n(n-1) - \frac{1}{2} \sum_{j=1}^p t_j(t_j-1) \right]^{1/2} \left[\frac{1}{2}n(n-1) \right]^{1/2} \quad (6)$$

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210 The decadal series that were then found to have a significant trend based on the MK results were
211 then furtherly analyzed by applying a series of multiple linear regression models (MLR). MLR were
212 used to model the decadal count of off-season TCs for basins that showed increasing or decreasing
213 trends in storm numbers to test if covariates associated with climatic variability and climate change
214 explained off-season TC frequency. MLR attempts to model the relationship between two or more
215 explanatory variables and a response variable by fitting a linear equation to observed data. The
216 notation for the model deviations is:

$$y_i = \beta_0 + \beta_1 x_{i1} + \beta_2 x_{i2} + \dots + \beta_p x_{ip} + \varepsilon_i \text{ for } i = 1, 2, \dots, n \quad (1)$$

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219 Three different MLR models were run for each off-season TC series that exhibited a statistically
220 significant trend, one MLR model with the climate change variables (SST, GMST & CC) as
221 predictors, another model with just the climate variability factors (ENSO, AMO & IOD) and a final
222 model with all of the variables included. Then the three MLR models (pre-season, post-season and
223 off-season) were run for each of the basins with increasing trends in off-season TCs, the best models
224 (highest adjusted R-squared and lowest p-value) were then selected for each of the series. The MLR
225 models were run in The R Project for Statistical Computing using the biglm package.

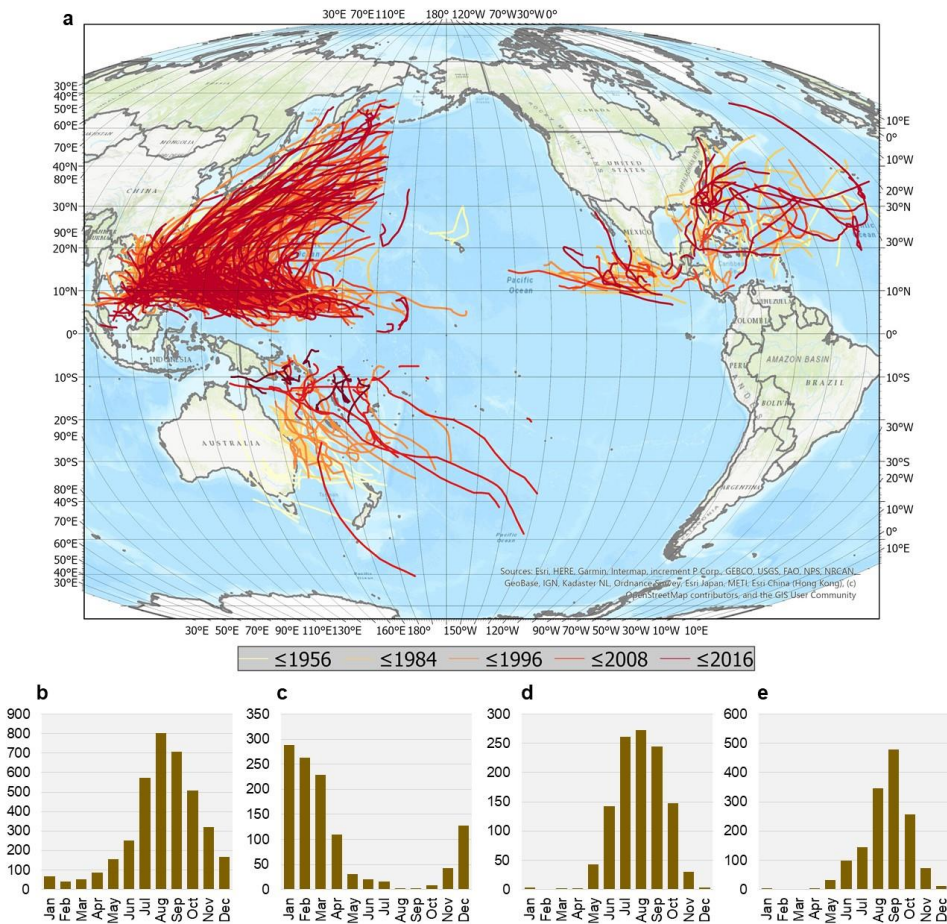
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227 Finally, stepwise selection MLR models were used to identify the climate variability or change factors
228 making the most statistically significant contributions to off-season increasing TC frequency. Here
229 we use stepwise selection which is a combination of the forward and backward procedures where you
230 start with no predictors, then sequentially add the most contributive predictors. After adding each new
231 variable, it removes the variables that no longer provide an improvement in the model fit (James et
232 al., 2014; Bruce and Bruce, 2017). The MLR and stepwise for the off-season TC count series for each
233 of the basins with significant increasing trends were run in The R Project for Statistical Computing
234 using the MASS package (Venables and Ripley, 2002).

235 5. Results & Discussion

236 When analyzing the number of TCs for all basins for the 1900-2019 period we found that 713
237 off-season storms occurred during that time, most during the months of May (NH pre-season and SH
238 post-season) with 430 and December (NH post-season) with 341 (Figure 1a, 1b). When looking at
239 the count of off-season TCs per basin we found that as expected the West Pacific (611) and South
240 Pacific (85) accounted for 81.3% of all off-season storm occurrences. When grouping the basins
241 between northern and southern hemispheres, we find that 89% of all off-season TCs occurred north
242 of the equator for the 1900-2019 period (Figure 1a, 1b). The North Atlantic and East Pacific basins
243 were found to be the ones with the lowest numbers of off-season TCs when compared to the other
244 two Pacific basins.



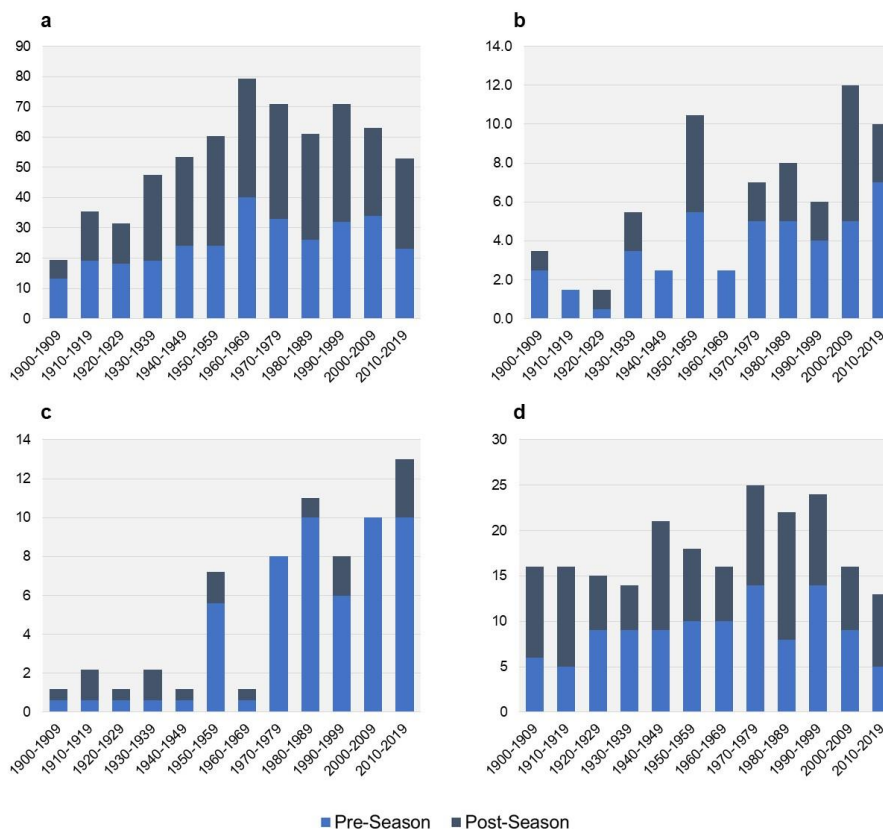
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247 **Figure 1.** Tracks of all out off-season TCs (a) and the number of storms per month for the WP (b),
 248 SP (c), EP (d) and NA (e) basins for the 1900-2019 period.

249 As shown in Figure 2, even after adding the estimated missing storms before the 1970 decade,
 250 most basins experienced their highest number of out off-season TCs (pre or post) in decades at or
 251 after 1960-69. The 1960-69 decade for the northern hemisphere basins (WP, NA and EP) was found to
 252 be the one with the highest number of pre off-season TCs with 69 and the 1950-1959 decade was
 253 identified as the one with the most post off-season storms with 68 (Figure 2a, 2b, 2c and 2d). When
 254 examining TC counts for all basins individually, we found that the NA and EP basins had their most
 255 active decades after 1970 and that the WP and SP basins experienced their highest storm count decade
 256 after 1960 (Figure 2c, 2d). It is important to note that these results already reflect the additional TCs
 257 that were added to the pre-satellite era.
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260 **Figure 2.** Adjusted decadal count of all observed and estimated off-season TCs for the WP (a), NA (b), EP (c)
 261 and SP (d) ocean basins for the 1900-2019 period.

262 The Mann-Kendall non-parametric tests for trends for all basins show that three basins exhibited
 263 statistically significant increasing trends in adjusted decadal off-season TC counts for the 1900-2019
 264 period (Table 2). The basins with statistically significant increasing trends were the EP (pre and off-
 265 season), NA (pre, post and off-season) and the WP (pre, post and off-season). The EP basin shows an
 266 increasing trend in pre- and off-season TCs that is more evident from the 1950s to the present (Figure
 267 2d), while the increasing trend in the NA basin can be observed from the 1970s to 2019.

268 The increasing trend in off-season TCs in the WP basin is more evident from the 1900 to 1969
 269 (Figure 2a), which was during the pre-satellite era where missing TCs were added to the series.
 270 However, no trend is found in the WP basin if the decadal counts are analyzed from the 1970s to the
 271 present. In the post-satellite era in the WP basin, the 1990-1999 decade was identified as the one with
 272 most off-season TCs, however the two following decades exhibited a decreasing trend. The EP and
 273 NA basins show significant increasing trends in off-season TC counts (Table 2). Opposites to the
 274 trends identified in the WP basin, the EP and NA also show increasing decadal counts after the 1960s
 275 and 1970s decades. The SP basin also exhibited a positive Tau coefficient, yet it was statistically
 276 insignificant (Figure 2d).

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Table 2. Results of Mann-Kendall trend tests for the 1900-2019 period for all ocean basins where TCs form.

Pre TCs	Tau S	P-value	Post TCs	Tau S	P-value	Off TCs	Tau S	P-value
EP	0.746	0.002	EP	0.098	0.723	EP	0.679	0.004
NA	0.572	0.015	NA	0.485	0.042	NA	0.554	0.016
SP	0.048	0.889	SP	0.015	1.000	SP	0.061	0.836
WP	0.554	0.016	WP	0.485	0.034	WP	0.534	0.019

Significant trends in bold.

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MLR models were run on the basins that exhibited statistically significant (< 0.05) increasing trends in decadal total off-season TC counts over time and here we report the best models for each of the series. The MLR results show that the statistically significant increasing trends in TC frequency for the EP (pre and off-season) and WP basins is best explained by climate change factors SST, GMST and CC at the 0.05 significance level (Table 3). Climate change factors accounted for 56% (pre-season) and 52% (off-season) of the increasing trend in TC counts for the EP basin. In the WP basin climate change factors explained 55% (pre-season), 64% (post-season) and 68% (off-season) of the trends in off-season TCs. Increasing trends in SSTs, GMST and moisture (CC) outside of the prime months of tropical storm development could promote more optimum conditions for higher off-season TC occurrences (Klozback, 2006; Hansen et al., 2010).

Table 3. Best multiple linear regression models (MLR) for basins with statistically significant increasing trends in off-season TCs.

Model	Multiple R-squared	Adjusted R ²	Factors	p-value
EP pre-season	0.682	0.563	SST, GMST & CC	0.021
EP off-season	0.653	0.522	SST, GMST & CC	0.030
NA pre-season	0.622	0.481	SST, GMST & CC	0.041
NA post-season	0.687	0.427	ENSO & AMO	0.130
NA off-season	0.552	0.384	SST, GMST & CC	0.070
WP pre-season	0.673	0.551	SST, GMST & CC	0.020
WP post-season	0.742	0.645	SST, GMST & CC	0.000
WP off-season	0.774	0.689	SST, GMST & CC	0.005

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The climate variability factors (ENSO, AMO & IOD) did not exhibit statistically significant relationships with increasing off-season TC counts, which shows that natural variability does not explain the incrementing number of storms in the EP and WP basins. MLR model results for the NA basin also showed the climate change variables accounting for 48% (pre-season) and 38% (off-season) of the increasing trend in TCs (Table 3). However, the MLR model results for the post-season months in the NA basin showed that the climate variability variables (ENSO & AMO) accounted for 42% of the increasing trend in TCs, yet the model was not found to be statistically significant. It is well known that cold phases of ENSO (La Niña) and warm phases of AMO tend to be associated with higher TC frequency in the North Atlantic ocean (Tang and Neelin, 2004; Briggs, 2008) and this could explain why those teleconnections were found to have the most significant influence on post-season TC frequency in the NA basin. Yet, it is important to note that in most basin series, including the NA, climate change variables explained more of the off-season TC increasing trend than the climate variability factors.



307 Stepwise MLR model results showed that climate change factors (SST, GMST & CC) were among
 308 the selected variables that explained most of the increasing trend in off-season TCs for all basins
 309 analyzed (Table 4). In the EP basin, SST, ENSO, and CC, accounted for 69% (pre-season) and 65%
 310 (off-season) of the increasing trend in TCs. In the NA basin, the stepwise procedure selected CC as
 311 the sole climate change factor that explained 52% (pre-season) and 40% (off-season) of the rising
 312 frequency in TC counts. However, CC & AMO were selected as the variables that explained (43%)
 313 most of the variability in TC frequency during the post-season months in the NA basin. Stepwise
 314 procedure results for the WP basin show that climate change and variability factors were selected as
 315 the best predictors of TC frequency, with GMST and CC accounting for 57% (pre-season), CC,
 316 GMST, ENSO and IPO explaining 72% (post-season) and 74% (off-season) of the variability of
 317 TCs.

Table 4. Stepwise multiple linear regression models (MLR) for basins with statistically significant increasing trends in off-season TCs.

Model	R-squared	Adjusted R ²	Factors	p-value
EP pre-season	0.777	0.694	SST, ENSO & CC	0.005
EP off-season	0.747	0.652	SST, ENSO & CC	0.008
NA pre-season	0.569	0.526	CC	0.004
NA post-season	0.687	0.427	CC & AMO	0.098
NA off-season	0.460	0.406	CC	0.015
WP pre-season	0.655	0.578	GMST & CC	0.008
WP post-season	0.826	0.726	CC, GMST, ENSO & IPO	0.008
WP off-season	0.839	0.747	CC, GMST, ENSO & IPO	0.006

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319 The EP experienced a steady increase in off-season TC total counts from 1900 to 2019 at a rate
 320 of 1.1 additional storms per decade. The decadal off-season total TC count series for the EP basin
 321 closely resembles the increasing trend in average SSTs and CCs (Fig 3a, 3c). The correlation between
 322 off-season TCs in the EP basin and ENSO is not as clear as the one between SST and CC, with some
 323 mostly warm ENSO decades like the 1990-1999 exhibiting lower storm counts and other periods with
 324 cooler phases dominating showing a higher number of cyclones. When SST patterns for areas in the
 325 EP basin where TCs develop are examined over time, we find that most tropical/sub-tropical ocean
 326 waters have experienced a statistically significant increasing trend in ocean surface temperatures from
 327 1900 to 2019 (Fig 3d). Similar to other studies (Hansen et al., 2010), we find that the EP tropical
 328 ocean surfaces have increased by 0.051 degrees C° per decade. When CC patterns are examined, we
 329 find that it has also experienced a statistically significant increasing trend in some areas in the EP
 330 basin (Fig 3e).

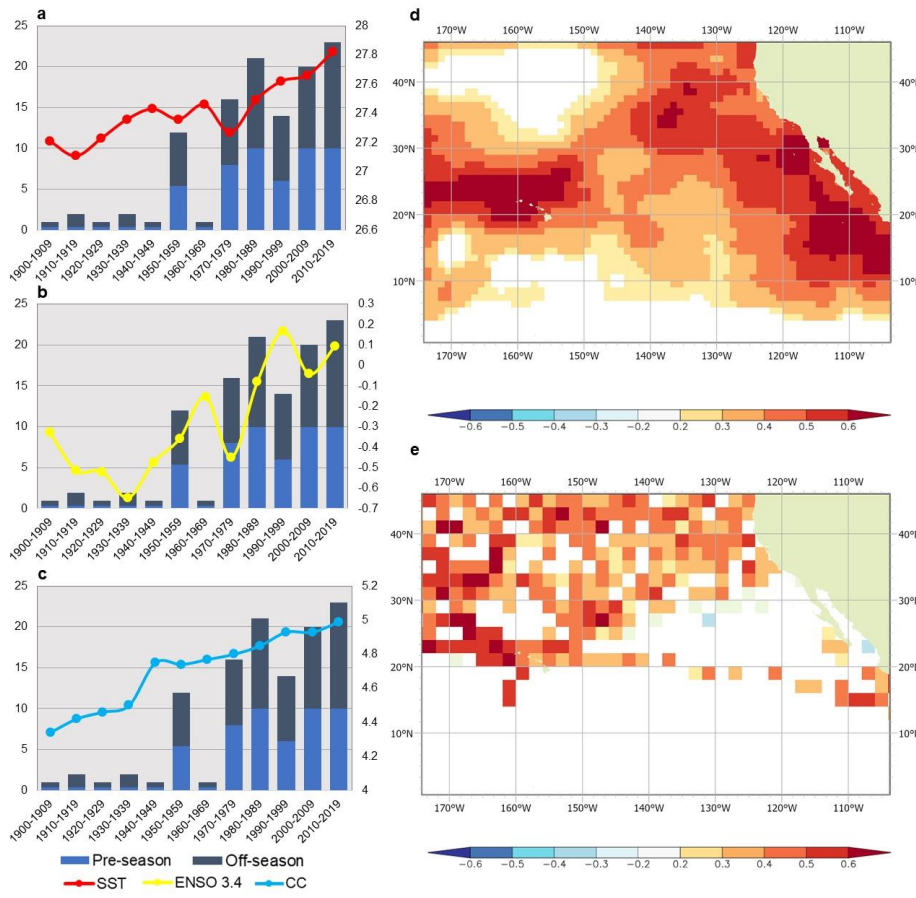
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332 The decadal off-season total TC count series for the NA basin closely resembles the increasing
 333 trend in average SSTs and CCs (Fig 4a, 4c). The NA decadal series shows a steady increase in off-
 334 season TC total counts from 1900 to 2019 at a rate of 0.7 additional storms per decade and an SST
 335 increasing trend of 0.055 C° per decade. Both the average decadal SST and CC series coincide with
 336 the peaks and valleys in off-season TC counts for the NA basin, with the 1950-1959 showing a high
 337 number of storms associated with high mean SSTs and CCs while the drop in storm counts in the
 338 1960-1969 decade matches a drastic drop in ocean surface temperatures (Figures 4a, 4c). Even though
 339 average SSTs increase to 0.135 C° per decade from 1970 to 2019, off-season TC total counts went
 340 down in the 1990-1999 and 2010-2019 decades, with the decade in between (2000-2009) exhibiting
 341 the highest number of off-season TCs (14) of all decades examined. However, it is important to note
 342 that 5 out of the 6 decades with the most off-season TCs in the NA basin occurred after the 1970s.

343



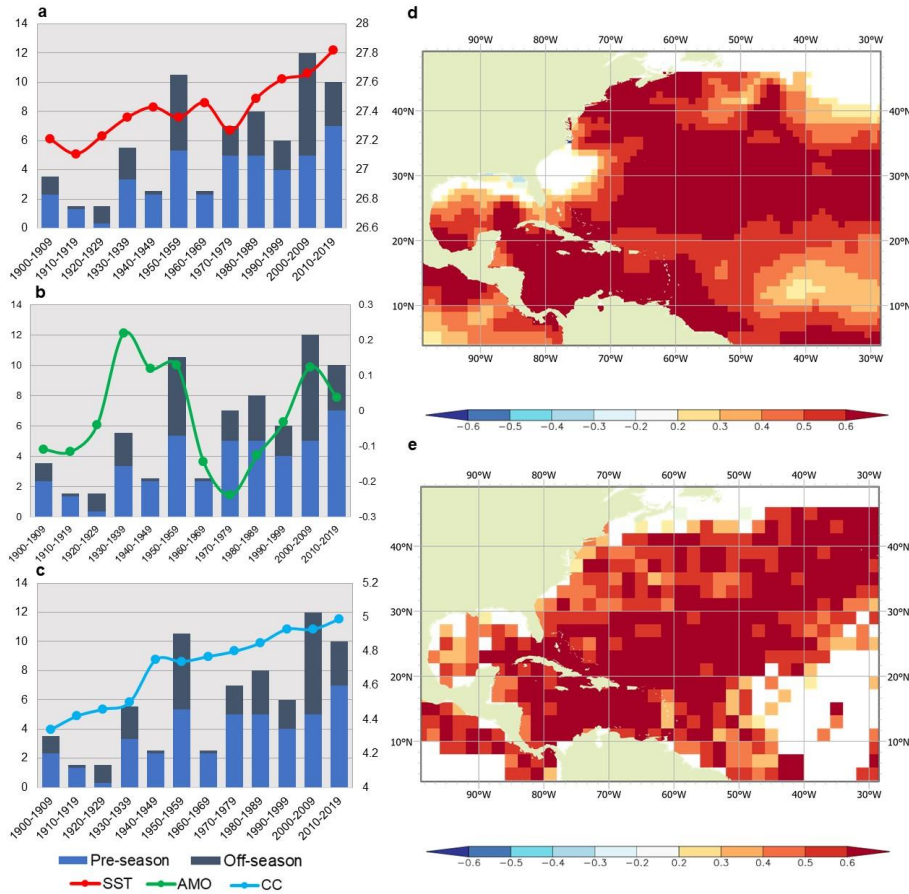
344 When North Atlantic SSTs are examined in areas where TCs form, we found that ocean surface
 345 temperatures have increased at a rate of 0.055 degrees C° per decade for the off-season months of
 346 Dec-March (Fig 4d). When CC patterns are examined, we find that it has also experienced a
 347 statistically significant increasing trend of 0.06 oktas (eighths of the sky that are covered in clouds)
 348 per decade in the North Atlantic basin since the 1900 (Figure 4e). If the NA pre/post off-season series
 349 is modified to begin in the 1960s, we find that SSTs have increased at a decadal rate of 0.082 C° per
 350 decade at a rate of 1.2 additional storms per decade. Overall, these results suggest that increasing
 351 trends in SSTs, which also drive increasing trends in evaporation rates associated with high CCs, are
 352 the physical mechanisms behind most of the recent increase in the total number of out of season TCs
 353 in the NA basin. The correlation between off-season TCs in the NA basin and AMO is not as clear as
 354 the one between SST and CC, with some warm AMO phases between 1930-1959 exhibiting lower
 355 storm counts while some cooler phases (1970-89) showing a higher number of cyclones.
 356



357

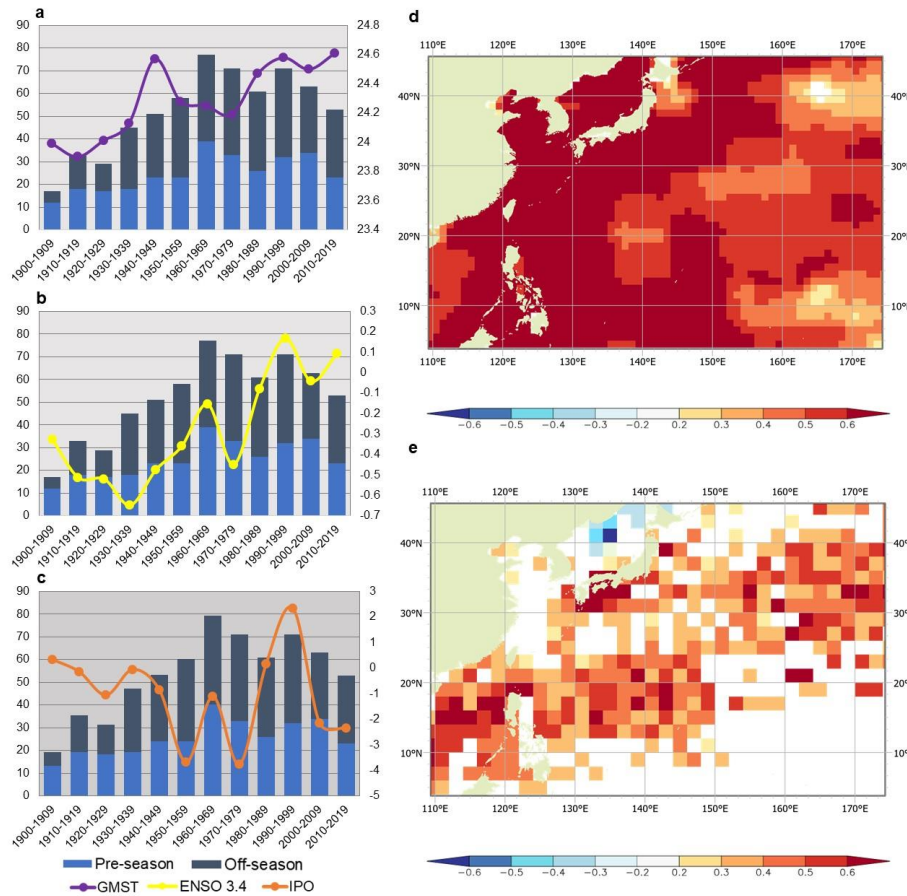
358 **Figure 3.** Decadal TC counts for the EP off-seasons and decadal average SSTs (a), decadal TC counts for
 359 the EP off-seasons and decadal average ENSO 3.4 (b), decadal TC counts for the EP off-seasons and
 360 decadal average Correlation between Time and Dec-May averaged CC (c), correlation between Time and
 361 Dec-May averaged SST (C°) for the 1900-2019 period (d) and correlation between Time and Dec-May
 362 averaged CC (oktas) for the 1900-2019 period (e).

363



364
 365 **Figure 4.** Decadal TC counts for the NA off-seasons and decadal average SSTs (a), decadal TC counts for the
 366 NA off-seasons and decadal average AMO (b), decadal TC counts for the NA off-seasons and decadal average
 367 Correlation between Time and Dec-May averaged CC (c), correlation between Time and Dec-May averaged
 368 SST (C°) for the 1900-2019 period (d) and correlation between Time and Dec-May averaged CC (oktas) for the
 369 1900-2019 period (e).

370
 371 The decadal off-season total TC count series for the WP basin closely resembles the increasing
 372 trend in GMST (Fig 5a). However, the WP basin experienced the highest count of off-season TCs in
 373 the 1960-69 decade, not in the more recent decades like the EP and NA basins. More importantly, if
 374 trend analysis for off-season TC counts is done from 1960-2019 in the WP basin, we find no
 375 statistically significant increasing or decreasing trend. However, it is important to note that four out
 376 of the five decades with most off-season TCs in the WP basin occurred after 1960. However, the
 377 2010-2019 decade was identified as the period with the lowest total number of off-season TCs even
 378 though increasing trends in mean SST, GMST and CC continued (Fig 5a, 5d and 5e). The decreasing
 379 number of off-season TCs in the last two decades coincided with a negative phase of the IPO, which
 380 suggests that TC frequency in the WP basin is influenced by fluctuations in the IPO (Fig 5c), whose
 381 recent negative phase since 1998 resembles La Niña-like SST anomaly patterns (Zhao et.al, 2018).
 382 Even though most of the variability in off-season TC frequency in the WP basin can be explained by
 383 climate change trends in GMST, SST and CC, the rest of the variance in TCs is account by fluctuations
 384 in the IPO and ENSO teleconnections.



385
 386

387 **Figure 5.** Decadal TC counts for the WP off-seasons and decadal average SSTs (a), decadal TC counts for the
 388 WP off-seasons and decadal average AMO (b), decadal TC counts for the WP off-seasons and decadal average
 389 Correlation between Time and Dec-May averaged CC (c), correlation between Time and Dec-May averaged
 390 SST (C°) for the 1900-2019 period (d) and correlation between Time and Dec-May averaged CC (oktas) for the
 391 1900-2019 period (e).

392 **6. Summary and concluding remarks**

393 The frequency of TCs that developed outside of their prime season months were analyzed to
 394 determine if trends in higher storm totals in the Pacific and Atlantic Ocean basins were associated
 395 with natural variability, climate change or both. Adjusted off-season decadal TC total counts for six
 396 ocean basins were analyzed for the 1900-2019 period in order to determine if the number of storms
 397 have been increasing over time. Mann-Kendall tests for trends were done and the basins that exhibited
 398 statistically significant increasing trends were then furtherly analyzed using multiple linear regression
 399 models and stepwise procedures to determine if those trends could be explained by fluctuations
 400 associated with climate variability, climate change trends or a combination of both.

401

402 The main results of this study suggest that decadal total off-season (pre/post) TC counts have
 403 significantly increase over time since the 1900 in the East Pacific (EP), North Atlantic (NA) and West
 404 Pacific (WP) basins. The EP and NA basins exhibited statistically significant increasing trends even
 405 if the analysis was done from the 1960s instead of the 1900. The WP basin showed an overall



406 increasing trend in the total number of off-season TCs per decade, yet if the analysis is done from the
407 1960s to the present, no statistically significant increasing trend is found. However, the three basins
408 that reflected an overall increase in decadal off-season TC frequency had their most active decades
409 after the 1970s.

410
411 Results from the best multiple linear regression (MLR) models show that the increasing decadal
412 count of off-season TCs has been found to be strongly associated with climate change trends in sea
413 surface temperature (SST), global mean surface temperature (GMST) and cloud cover (CC) in all
414 three basins (EP, NA and WP). The MLR model where climate variability variables (ENSO and
415 AMO) explained most of the variance in off-season TC counts was in the storm decadal counts for
416 the post-season months of the NA basin.

417
418 Results of the MLR stepwise procedures showed that the selected variables that accounted for
419 most of the variability in off-season TCs for the EP basin were SST, CC and ENSO, while CC (pre-
420 season and off-season) and AMO (post-season) were chosen as the best variables for the NA basin.
421 The stepwise procedure identified the climate change trends in GMST and CC, and fluctuations in
422 ENSO and IPO as the variables that accounted for most of the variability in decadal off-season total
423 TC counts in the WP basin,

424
425 The findings of this study suggest that trends in SST, GMST and CC associated with climate
426 change are not only altering the frequency (Klotzback, 2006; Saunders and Lea, 2006; Hansen et al.,
427 2010) and intensity of TCs that develop during the peak months of the season, they are also altering
428 the total number of storms that form in the off-season months (Dec-May), especially in the EP and
429 NA basins. The results of this study have important implications for the NA and EP basins, if off-
430 season TCs have been increasing in frequency since the 1900 we can expect that this trend associated
431 with climate change would continue in future decades. This increasing number of off-season TCs
432 could potentially impact societies in their path during times of the year when storms are least
433 expected, possibly increasing environmental and economic impacts in areas that are already
434 experiencing the effects of climate change exacerbated phenomena.

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437
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444 measurements that can be accessed from the Kingdom of the Netherlands Meteorology Institute
445 (KNMI; <https://climexp.knmi.nl/start.cgi>). The global mean surface temperature (GISTEMP v4)
446 data supporting this article are based on publicly available measurements that from the NASA
447 Goddard Institute for Space Studies (GISS; <https://climatedataguide.ucar.edu/climate-data/global-surface-temperature-data-gistemp-nasa-goddard-institute-space-studies-giss>). The El Niño Southern
448 Oscillation (ENSO 3.4) data supporting this article are based on publicly available measurements
449 from the National Oceanic and Atmospheric Administration Physical Sciences Lab (PSL;
450 https://psl.noaa.gov/gcos_wgsp/Timeseries/Data/nino34.long.data). The Atlantic Multidecadal
451 Oscillation data supporting this article are based on publicly available measurements from the
452 National Center for Atmospheric Research (NCAR; <https://climatedataguide.ucar.edu/climate-data/atlantic-multi-decadal-oscillation-amo>). The Interdecadal Pacific Oscillation data supporting
453 this article are based on publicly available measurements from the National Oceanic and
454 Atmospheric Administration Physical Sciences Lab (PSL;
455 <https://psl.noaa.gov/data/timeseries/IPOTPI/>).

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459 **References**

- 460 Bradley, R.S., M.K. Hughes and M.E. Mann (2006) Authors were clear about hockey-stick
461 uncertainties. *Nature*, 442 (7103): 627. <https://doi.org/10.1038/442627b>
- 462 Bruce, Peter, and Andrew Bruce. 2017. *Practical Statistics for Data Scientists*. O'Reilly Media.
- 463 Camargo, S. J., & Sobel, A. H. (2005). Western North Pacific tropical cyclone intensity and
464 ENSO. *Journal of Climate*, 18(15), 2996-3006. <https://doi.org/10.1175/JCLI3457.1>
- 465 Christensen J.H. et al. in *Climate Change (2013) The Physical Science Basis. Contribution of*
466 *Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate*
467 *Change*, (eds). Stocker T.F., Qin D., Plattner G.-K., Tignor M., Allen S.K., Boschung J., Nauels
468 A., Xia Y., V.B. & P.M.M.) Cambridge; Cambridge University Press.
- 469 Emanuel, K.A. (2005) Increasing destructiveness of tropical cyclones over the past 30 years.
470 *Nature*, 436: 686–688. <https://doi.org/10.1038/nature03906>
- 471 Emanuel, K.A. (2013) Downscaling CMIP5 climate models shows increased tropical cyclone
472 activity over the 21st century. *PNAS July 23, 2013* 110 (30) 12219–12224. Accessed from:
473 <https://doi.org/10.1073/pnas.1301293110>
- 474 Goldenberg, S.B., C.W. Landsea, A.M. Mestas-Nuñez and W.M. Gray (2001) The recent increase
475 in Atlantic hurricane activity: Causes and implications. *Science*, 293: 474–478. DOI:
476 10.1126/science.1060040
- 477 Hansen, J., Ruedy, R., Sato, M., & Lo, K. (2010). Global surface temperature change. *Reviews of*
478 *Geophysics*, 48(4). <https://doi.org/10.1029/2010RG000345>
- 479 Henley, B.J., Gergis, J., Karoly, D.J., Power, S.B., Kennedy, J., & Folland, C.K. (2015). A Tripole
480 Index for the Interdecadal Pacific Oscillation. *Climate Dynamics*, 45(11-12), 3077-
481 3090. <http://dx.doi.org/10.1007/s00382-015-2525-1>. Accessed on MM DD YYYY at
482 "/data/timeseries/IPOTPI".
- 483 Hipel, K.W. and McLeod, A.I., (2005). *Time Series Modelling of Water Resources and*
484 *Environmental Systems*. Electronic reprint of our book originally published in 1994.
485 <http://www.stats.uwo.ca/faculty/aim/1994Book/>
- 486 Holland, G.J. and P.J. Webster (2007) Heightened tropical cyclone activity in the North Atlantic:
487 Natural variability or climate trend? *Philos. Trans. Roy. Soc.*, 365: 2695–2716.
488 <https://doi.org/10.1098/rsta.2007.2083>
- 489 James, Gareth, Daniela Witten, Trevor Hastie, and Robert Tibshirani. 2014. *An Introduction to*
490 *Statistical Learning: With Applications in R*. Springer Publishing Company, Incorporated.
- 491 Knapp, K. R., H. J. Diamond, J. P. Kossin, M. C. Kruk, C. J. Schreck, 2018: International Best
492 Track Archive for Climate Stewardship (IBTrACS) Project, Version 4. [indicate subset used].
493 NOAA National Centers for Environmental Information. non-government domain
494 <https://doi.org/10.25921/82ty-9e16>
- 495 Knutson, T.R., J.L. McBride, J. Chan, K. Emanuel, G. Holland, C. Landsea, I. Held, J.P. Kossin,
496 A.K. Srivastava and M. Sugi (2010) Tropical cyclones and climate change. *Nature Geoscience*,
497 3: 157–163; doi:10.1038/ngeo779
- 498 Klotzbach, P.J. (2006) Trends in global tropical cyclone activity over the past twenty years (1986–
499 2005) *Geophys. Res. Lett.*, 33: L10805, pp 1–4. <https://doi.org/10.1029/2006GL025881>
- 500 Landsea, C.W. (2005) Meteorology: Hurricanes and global warming. *Nature*, 438: E11–E12.
501 <https://doi.org/10.1038/nature04477>



- 502 Landsea, C.W., B.A. Harper, K. Hoarau, J.A. Knaff (2006) Can we detect trends in extreme tropical
503 cyclones? *Science* 2006, **313**:452–454. DOI: 10.1126/science.1128448
- 504 Landsea, C. (2007). Counting Atlantic tropical cyclones back to 1900. *Eos, Transactions American*
505 *Geophysical Union*, 88(18), 197-202. <https://doi.org/10.1029/2007EO180001>
- 506 Mahala, B. K., Nayak, B. K., & Mohanty, P. K. (2015). Impacts of ENSO and IOD on tropical
507 cyclone activity in the Bay of Bengal. *Natural Hazards*, 75(2), 1105-1125.
508 <https://doi.org/10.1007/s11069-014-1360-8>
- 509 Mann, H.B. (1945), Nonparametric tests against trend, *Econometrica*, 13, 245-259.
- 510 Mann, M.E., T.A. Sabbatelli and U. Neu (2007) Evidence for a modest undercount bias in early
511 historical Atlantic tropical cyclone counts. *Geophys. Res. Lett.*, 34: L22707;
512 doi:10.1029/2007GL031781
- 513 Mann, M.E., J.D. Woodruff, J. Donnelly and Z. Zhihua (2009) Atlantic hurricanes and climate over
514 the past 1,500 years. Vol. 460 (13 August 2009); doi:10.1038/nature0821.
- 515 McLeod, A. I. (2005). Kendall rank correlation and Mann-Kendall trend test. R Package Kendall.
- 516 Nogueira, R. C., Keim, B. D., Brown, D. P., & Robbins, K. D. (2013). Variability of rainfall from
517 tropical cyclones in the eastern USA and its association to the AMO and ENSO. *Theoretical*
518 *and applied climatology*, 112(1-2), 273-283. <https://doi.org/10.1007/s00704-012-0722-y>
- 519 Saunders, M., Lea, A. Large contribution of sea surface warming to recent increase in Atlantic
520 hurricane activity. *Nature* 451, 557–560 (2008) doi:10.1038/nature06422
- 521 Trenberth, K.E. and D.J. Shea (2006) Atlantic hurricanes and natural variability in 2005. *Geophys.*
522 *Res. Lett.*, 33: L12704, doi:10.1029/2006GL026894
- 523 Trenberth, K. E. (2007). Warmer oceans, stronger hurricanes. *Scientific American*, 297(1), 44-51.
524 <https://www.jstor.org/stable/26069374>
- 525 Trenberth, Kevin, Zhang, Rong & National Center for Atmospheric Research Staff (Eds). Last
526 modified 10 Jan 2019. "The Climate Data Guide: Atlantic Multi-decadal Oscillation (AMO)."
527 Retrieved from [https://climatedataguide.ucar.edu/climate-data/atlantic-multi-decadal-](https://climatedataguide.ucar.edu/climate-data/atlantic-multi-decadal-oscillation-amo)
528 [oscillation-amo](https://climatedataguide.ucar.edu/climate-data/atlantic-multi-decadal-oscillation-amo).
- 529 Venables, W. N. and Ripley, B. D. (2002) *Modern Applied Statistics with S*. Fourth edition.
530 Springer.
- 531 Walsh, K. J., McBride, J. L., Klotzbach, P. J., Balachandran, S., Camargo, S. J., Holland, G., ... &
532 Sugi, M. (2016). Tropical cyclones and climate change. *Wiley Interdisciplinary Reviews:*
533 *Climate Change*, 7(1), 65-89.L doi: 10.1002/wcc.371
- 534 Wang, C., & Lee, S. K. (2008). Global warming and United States landfalling hurricanes.
535 *Geophysical Research Letters*, 35(2). <https://doi.org/10.1029/2007GL032396>
- 536 Zhao, J., Zhan, R., Wang, Y., & Xu, H. (2018). Contribution of the interdecadal Pacific oscillation
537 to the recent abrupt decrease in tropical cyclone genesis frequency over the western North
538 Pacific since 1998. *Journal of Climate*, 31(20), 8211-8224. [https://doi.org/10.1175/JCLI-D-18-](https://doi.org/10.1175/JCLI-D-18-0202.1)
539 [0202.1](https://doi.org/10.1175/JCLI-D-18-0202.1)